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1 **Oil palm 'slash-and-burn' practice increases post-fire greenhouse gas emissions and**  
2 **nutrient concentrations in burnt regions of an agricultural tropical peatland**

3

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20

21 **Abstract**

22 Fire is one of the major issues facing Southeast Asian peatlands causing socio-economic,  
23 human health and climate crises. Many of these fires in the region are associated with land  
24 clearing or management practices for oil palm plantations. Here we study the direct post-fire  
25 impacts of slash-and-burn oil palm agriculture on greenhouse gas emissions, peat physico-  
26 chemical properties and nutrient concentrations. Greenhouse gas (GHG) emissions were  
27 measured using Los Gatos ultraportable greenhouse gas analyser one month after a fire in dry  
28 season and five months after the fire event, in wet season. Surface soil samples were collected  
29 from each individual GHG measurement points, along with 50 cm cores from both burnt and  
30 non-burnt control areas for lab analyses. As an immediate post-fire impact, carbon dioxide  
31 (CO<sub>2</sub>) and methane (CH<sub>4</sub>) emissions, pH, electrical conductivity, and all macronutrient  
32 concentrations except nitrogen (N) were increased multi-fold, while the redox potential, carbon  
33 (C) and N content were greatly reduced in the burnt region. While some of the properties such  
34 as CO<sub>2</sub> emissions, and electrical conductivity reverted to normal after five months, other  
35 properties such as CH<sub>4</sub> emissions, pH and nutrient concentrations remained high in the burnt  
36 region. This study also found very high loss of surface peat C content in the burnt region post  
37 fire, which is irreversible. The results also show that surface peat layers up to 20 cm depth were  
38 affected the most by slash-and-burn activity in oil palm agriculture, however the intensity of  
39 fire can vary widely between different oil palm management and needs further research to fully  
40 understand the long term and regional impacts of such slash-and-burn activity in tropical  
41 peatlands.

42 **Keywords:** tropical peat fire; oil palm; carbon dioxide; methane; land use change; burnt  
43 peatlands.

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## 46        **1. Introduction**

47        One of the most fundamental challenges of the 21<sup>st</sup> century is to meet the global food and  
48        energy needs of the growing population, whilst conserving nature and soil carbon (C) storage  
49        to mitigate climate change (Godfray et al., 2010). Oil palm is considered as a crop for the future  
50        to serve these global needs, owing to its high productivity and versatility in usage (Abdullah et  
51        al., 2009; Corley, 2009; Tan et al., 2009). Oil palm has become the most consumed vegetable  
52        oil in the world over the past decade, and already represents 30% of the global usage (Lam et  
53        al., 2019). However, as oil palm is an equatorial crop, this has come at the expense of some of  
54        the world's most biodiverse and C rich ecosystems (Corley, 2009; Koh and Wilcove, 2008;  
55        Murdiyarso et al., 2010; Murdiyarso et al., 2019). For the same reason, oil palm has a higher C  
56        footprint than most other vegetable oil crops (Schmidt, 2015), with its expansion into tropical  
57        peatlands further increasing its C footprint (Reijnders and Huijbregts, 2008).

58        Oil palm is native to West Africa, and was first introduced to Malaysia as an ornamental plant  
59        in 1875 (Abdullah et al., 2009) and then commercially as a crop of choice for agricultural  
60        diversification in 1917 (Rashid et al., 2013). The initial expansion of oil palm plantations was  
61        in mineral soil at the expense of other agricultural plantations such as cocoa, rubber and coconut  
62        (Basiron, 2007; Dhandapani, 2015). Since then, oil palm has rapidly expanded in Southeast  
63        Asia, greatly expanding into peatlands from early 1990s (Miettinen et al., 2012; Shevade and  
64        Loboda, 2019). As a consequence, Malaysia had the highest 21<sup>st</sup> century deforestation rate  
65        globally (Hansen et al., 2013). A forested land should be completely cleared of vegetation,  
66        drained and levelled to establish an oil palm plantation (Luskin and Potts, 2011). The lowering  
67        of water table in peatlands make them highly susceptible for fire, as the dry peat itself is highly  
68        flammable (Posa et al., 2011). Unlike contexts involving mineral soils, peat fires can sustain  
69        both above and below ground, profoundly affecting surface and subsurface biota and seed bank,  
70        damaging the vegetation structure (Posa et al., 2011; Turetsky et al., 2015; Uda et al., 2017).

71 The forest fires in Southeast Asia are often associated with land clearing for oil palm  
72 plantations (Chazdon, 1998; Dennis et al., 2005). An estimated 60,000 peatland fires have  
73 occurred in Indonesia between 1997 and 2007, causing harmful haze (Tan et al., 2009). About  
74 2.6 million hectares of land was burned in Indonesia between June and October of 2015, which  
75 led Indonesian president to cancel existing concession on agricultural expansion and make  
76 conversion of peatlands to agricultural land illegal (World Bank, 2016). However, despite this,  
77 fires and the use of fire in peatland agriculture persists in both Malaysia and Indonesia.  
78 Furthermore, the use of fire is not just confined to the initial conversion and establishment  
79 stage, it is also commonly used in 'slash-and-burn' to burn the waste from previous generation  
80 oil palms to clear land for the next generation of crops. Slash-and-burn practice has been used  
81 all round the world as a quick short term activity to shift agriculture or for land clearing  
82 (Myllyntaus et al., 2002). Even though under certain circumstances in some ecosystems, slash-  
83 and-burn practice can be sustainable (Kleinman et al., 1995; Myllyntaus et al., 2002), it is  
84 certainly not sustainable in tropical peatlands in any circumstance, as dried peat itself is highly  
85 flammable, and fundamentally any agriculture involving drainage itself in tropical peatlands is  
86 not sustainable (Evers et al., 2017).

87 Fires in such tropical peatlands can convert large amount of C stored in peat for millennia into  
88 carbon dioxide (CO<sub>2</sub>) and cause socio-economic, human health, pollution and climate crises  
89 (Cheong et al., 2019; Page et al., 2002; Turetsky et al., 2015; Wiggins et al., 2018). Tropical  
90 peatlands in their natural undisturbed state are fire resistant, and historically fire had not played  
91 any significant role in tropical peatland ecology (Turetsky et al., 2015). Most of these fires  
92 have anthropogenic origin such as clearing of forest land for oil palm or similar agricultural  
93 use, and slash-and-burn agricultural practices (Page et al., 2002). As tropical peatlands are  
94 naturally acidic, this deliberate use of fire for agricultural practices brings short-term  
95 advantages such as increase in pH and reduced cost for land conversion (Islam et al., 2016).

96 Though there were some research and estimates of direct emission from fire as smoke, haze or  
97 CO<sub>2</sub> (Cheong et al., 2019; Hu et al., 2019; Hu et al., 2018; Islam et al., 2016; Smith et al., 2018;  
98 Varkkey, 2013; Varkkey, 2016; Wiggins et al., 2018), the subsequent impact of fire on peat  
99 physico-chemical properties and long term peat surface GHG emissions are not well  
100 documented.

101 Peat fires can continue for a long time, lasting for months in the form of smouldering, persistent  
102 flameless low-heat combustion of organic matter (Ohlemiller, 1985; Rein, 2013). Smouldering  
103 in peatlands can be a complex process with interactions between aerobic and anaerobic  
104 conditions, where pyrolysis of porous fuel occurs in high temperature at anaerobic conditions,  
105 producing char (Ohlemiller, 1985; Rein, 2013) and the produced char is then oxidised when it  
106 comes in contact with atmospheric oxygen, simultaneously producing and burning char during  
107 the process (Rein, 2013). These complex interactions of aerobic and anaerobic conditions  
108 along with other variations in intensity of fire and peat moisture levels can greatly impact peat  
109 physico-chemical properties such as pH, electrical conductivity and redox potential, yet these  
110 properties are not well documented for fire affected tropical peatlands.

111 The limited number of studies that do exist, have shown that peat fire significantly reduces the  
112 organic matter content of the peat soil (Sazawa et al., 2018), as that very organic matter is lost  
113 as C gases during the fire (Wiggins et al., 2018). Sazawa et al. (2018) also found that fire and  
114 resultant heat dehydrates and denatures organic matter, which leaves fire affected peat more  
115 susceptible to greater C losses, either through oxidation or repeated fire. Fire is generally found  
116 to increase both total and available nutrient concentrations in mineral soils due to disintegration  
117 of complex forms held in plants, and ash addition (Giardina et al., 2000; Van Reuler and  
118 Janssen, 1993). This increase in available nutrients make soils more fertile, which is another  
119 motive for the practice of slash-and-burn agriculture (Giardina et al., 2000). There are also  
120 instances when total nutrient concentrations increased immediately after fire and then levelled

121 off after few months (Kutiel and Naveh, 1987). The decrease in nutrient concentrations after  
122 few months may be because some of the nutrients are highly prone to leaching (Beliveau et al.,  
123 2015; Ulery et al., 1993). Similar increase immediately after a fire was also observed for other  
124 physico-chemical characteristics such as pH and electrical conductivity because of the addition  
125 of ash (Bang-Andreasen et al., 2017; Gay-Des-Combes et al., 2017), before reverting back to  
126 normal over time (Kutiel and Naveh, 1987). Fire may also effect increased mineralisation  
127 indirectly through soil microbial processes (Gay-Des-Combes et al., 2017). This is especially  
128 important in nutrient poor and acidic tropical peatlands, where fire is seen as a short-term and  
129 cost-effective solution to make peat more cultivable. Naturally, tropical peat is a very complex  
130 soil system and many of the environmental and microbiological interactions are understudied  
131 and not well known (Dhandapani et al., 2019c). The interactions of fire in such a complex soil  
132 system that is, in itself flammable (Rein, 2013; Uda et al., 2017), needs a more intensive and  
133 greater number of research with varying environmental parameters to comprehend and infer a  
134 cause and effect pattern.

135 In this study, we aim to evaluate the direct post-fire impact of slash-and-burn agriculture on  
136 peat physico-chemical properties, nutrient concentrations and GHG emissions. We hypothesise  
137 that fire increases both CO<sub>2</sub> and methane (CH<sub>4</sub>) emission as an immediate post-fire effect. We  
138 also hypothesise that fire significantly affects peat physico-chemical properties and nutrient  
139 concentrations at surface layers, driven by high heat denaturing organic matter and the addition  
140 of ash increasing pH and nutrient concentrations. We postulate that changes in peat physico-  
141 chemical properties exhibit significant correlation with GHG emissions.

## 142 **2. Materials and Methods**

### 143 **2.1. Study site and sampling strategy**

144 The study site ( $3^{\circ}25'17.7''\text{N}$   $101^{\circ}18'44.4''\text{E}$ ) is located at Kampong Raja Musa village in North  
145 Selangor peatlands, Malaysia (Fig 1). The climatic conditions of the site is tropical, and the  
146 soil classification comes under Histosols. The North Selangor peatland complex is the second  
147 largest area of peatland in Peninsular Malaysia, and is located adjacent to Thennamaram region,  
148 where oil palm was first commercially planted in Malaysia. The site is bordered with oil palm  
149 monocropping, polyculture consisting of oil palm, yam and pineapple, pineapple  
150 monocropping under an electric pylon trail, and a gravel road on the fourth side. The study site  
151 and all the neighbouring land-use type blocks are roughly 2 ha in size each, consistent with  
152 other small-holdings in the village (Dhandapani et al., 2019a,b). Two drainage ditches run as  
153 two parallel borders with the oil palm monoculture, and the oil palm, yam and pineapple  
154 polyculture. There is no drainage ditch within the site itself (Fig 2). The previous generation of  
155 oil palm monoculture was cleared and the waste such as dead wood and fronds were stacked  
156 in two parallel lines within the site. The stacked piles were burnt on June 2018, but with the  
157 fire front extending into the peat itself, and smoke still visible from smouldering of the peat  
158 surface for few weeks. The depth of burn was uneven throughout the burnt region, with surface  
159 reaching mineral layer in some parts, and some other parts still containing peat to a depth of  
160 at least 50 cm from surface. New young oil palm rows were planted in non-burnt area in July  
161 2018. Alongside this, a pineapple crop was planted in all open area in the non-burnt region and  
162 a banana crop was planted in rows in the burnt area, in between dry season (July 2018) and wet  
163 season (December 2018) sampling. Complete random sampling was carried out with 20 GHG  
164 measurements and surface peat (0-5 cm) collection each for burnt and non-burnt region during  
165 August 2018 dry season.

166 All sampling points were at least 1 metre away from each other. The same method was used  
167 for December 2018 wet season sampling with reduced number of measurements to 10 each for  
168 burnt and non-burnt region. The dry season measurement points were not the exact

169 measurement points used for the wet season, however they are in the same general area of burnt  
170 and non-burnt locations used for the wet season measurements. The measurement points from  
171 the first sampling were not marked with collars in order to minimise the disturbances from  
172 disturbance from our field measurements, and hence a different set of random measurement  
173 points in the same region were used for wet season sampling. Weather information for each  
174 sampling period is given in Table 1. During the dry season sampling, 3 peat cores to a depth of  
175 50 cm were collected for each burnt and non-burnt region. In some areas of burnt area peat was  
176 completely burnt and mineral layers was reached at the surface. Selective sampling was carried  
177 out for core collection, unlike complete random sampling for GHG measurements and  
178 associated surface peat collection. The 3 peat cores for burnt region were collected from burnt  
179 areas with peat remaining and not the areas with surface mineral exposure, to characterise the  
180 impact on leftover peat through the peat depth profile.

## 181 **2.2. Greenhouse gas emissions**

182 CO<sub>2</sub> and CH<sub>4</sub> emissions from the peat surface were measured using a Los Gatos (San Jose,  
183 California, USA) ultraportable greenhouse gas analyser as described in Dhandapani et al.  
184 (2019a,b,c). The gas measurements were made using closed dynamic chamber method with an  
185 inlet and outlet connecting to the gas analyser. The dynamic chamber was inserted up to 1 cm  
186 into the peat during each measurement, to seal the chamber to the surface. The chamber was  
187 15 cm high with 27 cm diameter. The Los Gatos gas analyser was set to record the gas flux  
188 changes within the chamber every 20 seconds, and measurement was made at each sampling  
189 point for 3 minutes, where the measurements from the first minute were ignored allowing the  
190 gas concentrations inside the chamber to settle down after initial disturbance. The gas  
191 concentrations were then converted to mg CO<sub>2</sub> m<sup>-2</sup> hr<sup>-1</sup> and µg CH<sub>4</sub> m<sup>-2</sup> hr<sup>-1</sup> for CO<sub>2</sub> and CH<sub>4</sub>  
192 respectively, as described in Dhandapani et al. (2019a,b,c).

### 193 **2.3. Peat analyses**

194 The procedure used for peat analyses were based on Dhandapani et al., (2019a,b,c). Peat  
195 temperature was measured *in situ*, using a digital thermometer Cosmark PDT300 (Norwich,  
196 UK). Peat samples were collected for measuring gravimetric moisture. For this, fresh peat was  
197 dried in an oven at 105°C for 48 hours. The gravimetric moisture was calculated as follows:

198 Gravimetric moisture (%) = Mass of the water lost in oven drying/ mass of oven dried peat

199 For pH, redox and electric conductivity measurements, 5 mL volume of peat sample was  
200 diluted in 10 mL deionised water in a centrifuge tube and shaken on a table shaker for 2 hours.  
201 The pH of the supernatant was then measured using a Eutech pH700 pH meter supplied by  
202 Thermo scientific (Loughborough, UK). The redox potential and electrical conductivity were  
203 measured using Eutech Ion 2100 (Thermo scientific, Loughborough, UK) and Groline  
204 HI98331 probe (Hanna, Leighton Buzzard, UK), respectively.

205 For analysing total C and nitrogen (N) content, all samples were oven dried (105°C for 48 h)  
206 and finely ground using a Fritsch mortar grinder pulveristte 2 (Brackley, UK). Approximately  
207 70 mg of sample was weighed into a Skalar ceramic crucible and the exact weight was  
208 recorded. The samples were then transferred to an auto sampler in Skalar primacs series  
209 SNC100 TC TN analyser (Breda, The Netherlands) and analysed for C and N content.

### 210 **2.4. Nutrient analyses**

211 Total nutrient concentrations of phosphorus (P), calcium (Ca), magnesium (Mg), sulphur (S)  
212 and potassium (K) in peat were analysed using inductively coupled plasma mass spectroscopy  
213 (ICP-MS). For this, approximately 0.15 g of oven dried (105°C for 48 h) and ground peat were  
214 weighed in microwave digestion tubes (MARSXpress vessels, CEM Microwave Technology  
215 Ltd., Buckingham, UK.). The digestion tubes are sealed with a stopper and a screw lid, after  
216 adding 10 mL of nitric acid to each sample. The digestion tubes were then placed in a

217 MARSXpress microwave (CEM Microwave Technology Ltd., Buckingham, UK.) and run at  
218 1600 W & 100% power with a ramp for 20 minutes and held for 20 minutes at 170°C. The  
219 tubes were left overnight in the microwave to cool down. The digested samples were then  
220 filtered and made up to 30 mL using milliQ water. Then, 1 mL of each sample were transferred  
221 in to 10 mL tube and further diluted with 9 mL of milliQ water. The samples were then analysed  
222 using 'Agilent Technologies (Milton Keynes, UK) 7900 ICP-MS' fitted with 'SPS 4'  
223 autosampler.

## 224 **2.5. Statistical analyses**

225 All the statistical analyses were carried out using Genstat® 17th edition (VSN international,  
226 2017). The significance of differences between sites for greenhouse gas emissions, nutrient  
227 concentrations and other physico-chemical properties were evaluated using linear mixed  
228 models with restricted maximum likelihood (REML) incorporating conditions (burnt or non-  
229 burnt) and season (August and December 2018) as fixed affects. Similar REML was also  
230 performed incorporating condition and depth as fixed effects, to identify the changes with  
231 depth. For the data sets that were not normally distributed, the data were log transformed. For  
232 data that did not meet normality assumption after log transformation, Boxcox transformation  
233 was used. Normality was assessed by visual examination of 4 different residual plots, namely  
234 histogram of residuals, fitted-value plot, normal plot and half-normal plot. Backward stepwise  
235 multiple regression was performed with CO<sub>2</sub> and CH<sub>4</sub> as response variables and nutrient  
236 concentration and other physico-chemical properties as fitted terms. Statistical significance  
237 was assessed at  $p < 0.05$  for all analyses.

## 238 **3. Results**

### 239 **3.1. Greenhouse gas emissions**

240 CO<sub>2</sub> emissions in the burnt region were twice as high as the CO<sub>2</sub> emissions from the non-burnt  
241 region in the dry season one month after the fire incident (Fig 3a; Table 2). In wet season, five  
242 months after the fire incident, CO<sub>2</sub> emissions in the burnt region levelled off to the same level  
243 as non-burnt region. The difference between the seasons were significant, while the interactions  
244 between regions and season were not significant (Table 2).

245 CH<sub>4</sub> emissions also varied significantly between the burnt and non-burnt regions (Fig 3b; Table  
246 2). CH<sub>4</sub> emissions were multi-fold higher in the burnt region than in the non-burnt region for  
247 both seasons, while the seasonal variations and interaction terms were not significant (Table  
248 2).

### 249 **3.2. Surface peat/soil properties**

250 Peat surface temperature did not significantly differ between burnt and non-burnt regions, with  
251 significantly higher temperature in dry season than in wet season (Fig 4a; Table 2). The wet  
252 season temperature was slightly yet significantly different between the regions, resulting in  
253 significant interactions terms. Gravimetric moisture did not significantly vary between the  
254 burnt and non-burnt regions, while both regions had significantly higher gravimetric moisture  
255 content in wet season than in dry season with no significant interaction between region and  
256 season (Fig 4b; Table 2). During the dry season, electrical conductivity was more than 3 times  
257 higher in the burnt region than in the non-burnt region resulting in significant difference  
258 between the regions. Electrical conductivity was greatly reduced in wet season compared to  
259 that of the dry season, resulting in significant seasonal variation (Fig 4c; Table 2). During the  
260 wet season there was no significant difference in electrical conductivity between the two  
261 regions while the difference was significant in dry season, resulting in significant interactions  
262 (Fig 4c; Table 2). Both redox potential and pH were significantly higher in the burnt region

263 than in the non-burnt region, with no significant difference between seasons nor significant  
264 interactions (Fig 4d,e; Table 2).

### 265 **3.3. Surface peat/soil C and nutrient content**

266 C content in the non-burnt region was more than double the content in the burnt region, with  
267 no significant difference between seasons or interactions (Fig 5a; Table 2). N content was also  
268 significantly higher in the non-burnt region than in the burnt region, and both regions had  
269 significantly higher N content in wet season than in the dry season, with no significant  
270 interactions between region and season (Fig 5b; Table 2). The C:N ratio was higher in the non-  
271 burnt region during the dry season and the ratio was higher in the non-burnt region for wet  
272 season, resulting in significant seasonal variations and significant interaction terms (Fig 5c;  
273 Table 2).

274 All the rest of the macronutrients except N and S were greater in the burnt region than in the  
275 non-burnt region and only S and K varied significantly between seasons with lower  
276 concentrations for the burnt region in wet season (Fig 5d-h; Table 2). All macronutrients except  
277 K exhibited significant interactions between seasons, as most nutrient concentrations did not  
278 significantly differ between seasons in the non-burnt region (Fig 5d-h; Table 2).

### 279 **3.4. Changes with peat depth**

280 Electrical conductivity and pH were significantly higher in the burnt than in non-burnt region,  
281 while the variations with depth and the interactions were not significant (Fig 6a,c; Table 3).  
282 Gravimetric moisture in the non-burnt region was significantly higher than that of burnt region  
283 in all depths (fig 6b,c; Table 3). The gravimetric moisture increased up to 20-30 cm and 30-40  
284 cm for burnt and non-burnt region respectively (fig 6b; Table 3) and then showed a slight  
285 decrease in deepest layers. Redox potential was significantly lower in the burnt region than in  
286 the non-burnt region in the top 3 surface layers upto the depth of 30 cm (Fig 6d; Table 3).

287 Redox potential increased with depth in the burnt region while the redox potential in the surface  
288 was slightly lowered in 10-20 cm and stayed at the same level in the deeper layers, resulting in  
289 significant seasonal variations and interaction terms (Fig 6d; Table 3).

290 C content was significantly greater in the non-burnt region than in the burnt region with no  
291 significant variations with depth nor interactions (Fig 7a; Table 3). N content and C:N ratio did  
292 not significantly vary between regions or depth, with no significant interactions between region  
293 and depth (fig 7b,c; Table 3).

294 All macronutrient concentrations except S and N were significantly greater in the burnt than in  
295 the non-burnt region (Fig 7d-h; Table 3). Mg, P and Ca concentrations significantly decreased  
296 with depth, which is more pronounced in the burnt region (Table 3). Mg and Ca exhibited  
297 significant interactions between region and depth, as the concentration in the burnt region  
298 decreased with depth whilst the concentrations in the non-burnt region stayed at the same level  
299 throughout the depths. K and S did not show any significant variations with depth, nor any  
300 significant interaction between region and depth (Fig 7f,g; Table 3).

### 301 **3.5. GHG emissions and environmental controls**

302 Backward stepwise multiple regression showed that CO<sub>2</sub> emissions were positively correlated  
303 with pH and electrical conductivity and negatively correlated with P and K concentrations  
304 (Table 4).

305 Similar regression for CH<sub>4</sub> emissions showed that the CH<sub>4</sub> emissions were negatively  
306 correlated with pH and P concentrations, and positively correlated with Ca concentrations  
307 (Table 4).

## 308 **4. Discussion**

309 Drainage for agriculture in peatlands were known to increase the CO<sub>2</sub> emissions with exposure  
310 of peat to aerobic decomposition (Couwenberg et al., 2010; Dhandapani et al., 2019a; 2019c),  
311 this study shows that the use of fire in management practices results in further increase in CO<sub>2</sub>  
312 emissions post fire (Fig 3a). There are several factors that may have played a part in this  
313 increase in emissions in fire affected region, such as increased pH and increased concentrations  
314 of macronutrients. It should be noted that the CO<sub>2</sub> emissions decreased and levelled off with  
315 that of non-burnt region after few months in the wet season, while the CO<sub>2</sub> emissions from the  
316 non-burnt region did not significantly vary between the seasons, suggesting that this artefact  
317 was not just seasonal variations of CO<sub>2</sub> emissions, but rather a reverting of emissions along a  
318 temporal gradient since the fire. However, while CO<sub>2</sub> emissions of pristine, secondary peat  
319 forests and first generation oil palm showing no significant seasonal variations, seasonal  
320 variation have been observed in second generation agricultural systems in the area (Dhandapani  
321 et al., unpublished; Dhandapani et al., 2019a; 2019c; 2020).

322 The observed increase in CO<sub>2</sub> emissions post fire may be due to intense disturbance during fire  
323 and subsequent changes in peat physicochemical properties such as increased pH, electrical  
324 conductivity and nutrient concentrations (Fig 4c-e; Fig 5d-h), along with breakdown of  
325 complex organic matter due to fire and new exposure of deeper peat layers to aerobic conditions  
326 (Sazawa et al., 2018). However, this increase is short lived as the surface conditions are  
327 stabilised over time and also possibly because of greater loss of newly acquired nutrients over  
328 time (Fig 4c; Fig 5d-h). This trend of initial increase and consequent decline in CO<sub>2</sub> emissions  
329 post fire was previously observed by Astiani et al. (2018) in West Kalimantan peatlands. It  
330 should be noted that in Astiani et al. (2018)'s study, the peak in emission was observed in the  
331 9<sup>th</sup> month post fire, and emissions stabilising at the 11<sup>th</sup> month post fire. The stabilisation of  
332 emissions in 5<sup>th</sup> month post fire (Fig 3a) in our study may be due to a lower intensity localised  
333 fire in our agricultural study site, compared to high intensity fire burning much greater forest

334 biomass in Astiani et al. (2018)'s study that would likely result in greater disturbance, greater  
335 nutrient addition from resultant ash, and thus would require greater time to recover and  
336 stabilise.

337 CH<sub>4</sub> emissions were also significantly higher in the burnt region than in the non-burnt region  
338 and stayed at similar higher levels even into the wet-season, five months after the fire event  
339 (Fig 3b). It was previously observed that fire would make the peat dry and effect lower CH<sub>4</sub>  
340 emissions from the burnt areas because of the lack of moisture (Davidson et al., 2019), but this  
341 study has shown that at similar moisture levels, burnt peat has a potential for higher CH<sub>4</sub>  
342 emissions than the non-burnt peat. This higher CH<sub>4</sub> emissions in the fire-affected region may  
343 be due to the availability of more labile C due to the denaturation of peat by fire and resultant  
344 heat (Sazawa et al., 2018). Once the moisture level for anaerobic conditions are met, CH<sub>4</sub>  
345 emissions are dependent on the availability of labile C in peatlands (Couwenberg, 2009). Fire  
346 in peatlands are also known to reduce methanotrophic activity (Danilova et al., 2015), further  
347 helping the increase in CH<sub>4</sub> concentration in the burnt areas. However, it should be noted that  
348 CH<sub>4</sub> emissions are very low overall, with mean values under 0.1 mg m<sup>-2</sup>hr<sup>-1</sup>, which is in line  
349 with the lower near zero fluxes observed in different oil palm agricultural plantations in the  
350 region. Southeast Asian peatlands are naturally low methane emitting landscapes (Dhandapani  
351 et al., 2019c; Hatano et al., 2016) compared to peatlands in other regions such as neotropics  
352 (Wright et al., 2013), boreal (Kettunen et al., 1996) or temperate regions (Abdalla et al., 2016).  
353 This validates our first hypothesis that greenhouse gas emissions are significantly increased in  
354 the fire affected burnt region in relation to the non-burnt region.

355 The surface peat temperature and moisture did not significantly vary between the burnt and  
356 non-burnt region, possibly because both regions are open with no ground cover to provide  
357 shade and cool down surface peat (Dhandapani et al., 2019a). The significant seasonal changes  
358 in both these properties were as expected (Dhandapani et al., 2019a), with higher moisture and

359 lower temperature in both regions during wet season, because of increased rainfall in wet  
360 season compared to the dry season (Global Environmental Centre, 2014). Similarly, higher  
361 moisture with increasing depth were also as expected, as the top layers in these agricultural  
362 peatlands are actively drained, and the surface layers are exposed to the sun, further facilitating  
363 increased temperature and evaporation in the surface layers. Most other physico-chemical  
364 properties and nutrient concentrations showed that the surface layers such as 0-10cm and 10-  
365 20cm showed significant difference between burnt and non-burnt regions before gradually  
366 narrowing in on the difference, to the same level in deeper layers starting from 20-30cm layer.

367 Redox potential is an important soil property which shows the electron exchange capacity for  
368 reduction and oxidation reactions that have great impact on nutrient availability and dynamics  
369 in soil by changing their electric charges (Niedermeier and Robinson, 2007; Søndergaard,  
370 2009). The reduction in redox potential in the burnt region shows the possible lack of oxidants  
371 in the burnt area (Fiedler, 2000). This is likely to be caused by fire, as this difference is observed  
372 only in the surface layers, with deeper layers having same level of redox potential in the deepest  
373 layers for both burnt and non-burnt regions. Fire were known to accumulate new particulate C  
374 forms in mineral soil humus layers that are highly resistant to redox reactions (González-Pérez  
375 et al., 2004), a similar trend is observed here in peat soil where fire affected surface layers have  
376 low redox potential.

377 Electrical conductivity and pH increased one month after fire, however electrical conductivity  
378 levelled off with that of non-burnt region during wet season, five months after the fire activity.  
379 The increase in pH after fire has been widely reported in different soil environments (Chungu  
380 et al., 2019; Heydari et al., 2017; Kennard and Gholz, 2001; Scharenbroch et al., 2012; Zaccone  
381 et al., 2014), that also holds true for highly acidic tropical peatlands where pH was almost twice  
382 as high in the burnt region compared to non-burnt region (Fig 4e). This increase may be directly  
383 related to the addition of ash from burning (Bang-Andreasen et al., 2017; Zaccone et al., 2014),

384 as shown by the significant increase only in the top 20 cm surface layers. Electrical conductivity  
385 exhibited very similar trend of change with depth as pH, possibly because of the same influence  
386 of ash addition in the top layers (Bang-Andreasen et al., 2017; Zacccone et al., 2014). Electrical  
387 conductivity indicates the salt content and is considered a rough estimate of soil nutrients that  
388 influence important soil processes such as GHG emissions (Visconti and De Paz, 2016). The  
389 significant reduction of electrical conductivity and levelling off during the wet season possibly  
390 shows that many of the nutrients that are made available after fire (Beest et al., 2019), were  
391 leached off in the proceeding few months (Beliveau et al., 2015), aided by heavy rainfall during  
392 the wet season (Global Environmental Centre, 2014). This is also evident in reduced  
393 concentration of all macronutrients in the wet season (Fig 5).

394 All of the macronutrients except N increased in the burnt region one month after fire and  
395 slightly decreased over time, yet remained higher than the concentrations in the non-burnt  
396 region throughout the seasons, except for S that had lower concentration in the burnt region  
397 than in the non-burnt region in wet season. The nutrients that are strongly basic and form  
398 important salts such as Mg and Ca (Visconti and De Paz, 2016), explicablely exhibited the same  
399 trend of change with depth as pH and electrical conductivity. P is also widely reported to  
400 increase in concentration in burnt peat and soil, as fire causes the release of P from organic  
401 materials (Beest et al., 2019; Wang et al., 2015). The increase in Ca, Mg and K concentrations  
402 in the burnt region can also be attributed to high content of these nutrient in wood ash that  
403 accumulated in the region as a result of burning previous generation oil palm stems (Kennard  
404 and Gholz, 2001). These nutrients are also found to be highly susceptible to leaching (Kennard  
405 and Gholz, 2001), explaining the lower concentrations in wet season sampling, few months  
406 after the fire event. C content was reduced to half in the surface layers, because of the C lost  
407 on the burning of organic peat on the surface (Wiggins et al., 2018). The peat cores were  
408 selectively collected in areas where the peat is not completely burned off, thus the difference

409 in C content between the surface layers in the cores and high replicate surface samples  
410 associated with GHG emission measurements (Fig 5a and 7a). However, it should be noted that  
411 lower number of replicate peat cores may also be a factor limiting the usability of these results.  
412 This validates our second hypothesis that peat physico-chemical properties and nutrient  
413 concentrations were significantly affected by fire, and most of these properties showed  
414 significant variations with depth.

415 These changes in physico-chemical properties were significantly correlated with CO<sub>2</sub> and CH<sub>4</sub>  
416 emissions, most visibly CO<sub>2</sub> emissions exhibited the same pattern of change as electrical  
417 conductivity and pH of surface peat. Multiple regression analyses showed that peat physico-  
418 chemical properties and total nutrient concentrations were strong predictors of change in CO<sub>2</sub>  
419 and CH<sub>4</sub> emissions (Table 4). CO<sub>2</sub> emissions from tropical peat of different land uses are  
420 known to exhibit positive correlations with pH (Dhandapani et al., 2019a) and our results show  
421 that the fire event resulting in sudden and steep increase in pH does not override this  
422 relationship. Inversely, CH<sub>4</sub> was negatively related with pH, this negative correlation has also  
423 been observed in northern peatlands (Weslien et al., 2009). Acidic condition is a prerequisite  
424 to functionalise precursors such as H<sub>2</sub>/CO<sub>2</sub>, acetic acid, formic acid, methanol and  
425 methylamine, for CH<sub>4</sub> production (Qing-Yu et al., 2019), thus explaining this functional  
426 relationship between pH and CH<sub>4</sub> emissions in the field.

427 The positive relationship between electrical conductivity or salinity and CO<sub>2</sub> emissions has  
428 been observed in other ecosystems (Capooci et al., 2019), however this is the first time such  
429 relationship is explored and reported in tropical peatlands. Though very high or excess salinity  
430 can be toxic and can negatively impact microbial activity and reduce CO<sub>2</sub> production (Setia et  
431 al., 2013; Yan et al., 2015), tropical peatlands are naturally nutrient poor (Sjögersten et al.,  
432 2011) and this increase in electrical conductivity due to fire boosted CO<sub>2</sub> production (Table 4).  
433 P concentrations were negatively related with both CO<sub>2</sub> and CH<sub>4</sub> emissions. Additionally, CO<sub>2</sub>

434 emissions were negatively related with K, and CH<sub>4</sub> emissions were positively related with Ca.  
435 Nutrient dynamics (with the exception of N) and their interactions with other peat processes  
436 such as C cycling were not well documented in agricultural tropical peatlands. Addition of  
437 manures containing these macronutrients had different effects on GHG emissions in different  
438 soil types (Ren et al., 2017), and their cause and effect pattern in tropical peatlands need more  
439 *in situ* research coupled with controlled experiments to understand the interactions, especially  
440 considering that they exhibit significant functional correlations with GHG emissions. However,  
441 P is known to support methane oxidation reactions in soil (Veraart et al., 2015) and similarly  
442 limited addition of Ca compounds were known to support methane production in varied set ups  
443 (Khor et al., 2015; Lar et al., 2010), suggesting similar functional relationship in tropical  
444 peatlands. In line with this, Murakami et al. (2005) also found CH<sub>4</sub> potential increased with  
445 liming (CaCO<sub>3</sub> addition) in agricultural tropical peatlands. This validates our third hypothesis  
446 that the changes in soil physico-chemical properties and nutrient concentrations correlate with  
447 changes in GHG emissions.

448 As oil palm is a high biomass crop, there is a potential alternative for the biomass removed  
449 from previous generation to be used for biochar production, which then could be used to enrich,  
450 and return C back to agricultural peatlands, potentially retaining some of the benefits of slash-  
451 and-burn practice such as increased pH and nutrient concentrations (Bista et al., 2019; Kong et  
452 al., 2014; Liew et al., 2018). Such biochar can be produced using methods ranging from  
453 constructing a conical soil pit in the ground, and simple conical or retort kilns to modern  
454 microwave pyrolysis equipment (Liew et al., 2018; Arafat Hossain et al., 2017). However  
455 socio-economic feasibility, and practicality of such alternatives for biomass management are  
456 yet to be explored and researched. There is a need for increased research in understanding the  
457 holistic environmental and socio-economic impacts of slash-and-burn practice in tropical  
458 peatlands, and other alternative practices for land clearing and biomass management between

459 oil palm generations. This study is a step towards such understanding of the impacts of different  
460 practices. Further research addressing these issues will help us make informed suggestions on  
461 land use policies and management practices, for rapidly increasing area of agricultural tropical  
462 peatlands.

## 463 **5. Conclusion**

464 The use of fire for slash-and-burn management practice severely alters peat physico-chemical  
465 properties and increases post-fire greenhouse gas emissions and peat nutrient concentrations.  
466 While some of the properties such as CO<sub>2</sub> emissions, and electrical conductivity reverted back  
467 to normal level after few months, other properties such as CH<sub>4</sub> emissions, pH and nutrient  
468 concentrations remained high in the burnt region, though nutrient concentrations were  
469 significantly reduced over time, possibly due to leaching. This study has shown very high loss  
470 of surface peat C content due to fire, which is irreversible. This has important implications if  
471 such a loss occurred in all agricultural peatlands in the region through slash-and-burn activity.  
472 The practice of intercropping was previously known to prolong defining peat properties and  
473 ameliorate environmental impacts of oil palm agriculture in peat, but the current results show  
474 that use of fire as a management practice overrides any such benefits from intercropping,  
475 because of high C loss. The peat surface physico-chemical properties and nutrient  
476 concentrations also exhibited functional correlations with GHG emissions, providing new  
477 insights into complex interactions between different biogeochemical processes in tropical  
478 peatlands. The results also show that peat layers from surface to 20 cm depth were most  
479 affected by this particular slash-and-burn activity in oil palm agriculture. However, the  
480 intensity of fire can vary widely between different oil palm management and needs further  
481 research to fully understand the long term and regional impacts of such slash-and-burn activity  
482 in tropical peatlands.

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490

## 491 **6. References**

- 492 Abdalla M, Hastings A, Truu J, Espenberg M, Mander Ü, Smith P. Emissions of methane from  
493 northern peatlands: a review of management impacts and implications for future  
494 management options. *Ecology and Evolution* 2016; 6: 7080-7102.
- 495 Abdullah AZ, Salamatinia B, Mootabadi H, Bhatia S. Current status and policies on biodiesel  
496 industry in Malaysia as the world's leading producer of palm oil. *Energy Policy* 2009;  
497 37: 5440-5448.
- 498 Arafat Hossain M, Ganesan P, Jewaratnam J, Chinna K. Optimization of process parameters  
499 for microwave pyrolysis of oil palm fiber (OPF) for hydrogen and biochar production.  
500 *Energy Conversion and Management* 2017; 133: 349-362.
- 501 Astiani D, Curran LM, Burhanuddin, Taherzadeh M, Mujiman, Hatta M, et al. Fire-driven  
502 biomass and peat carbon losses and post-fire soil CO<sub>2</sub> emission in a West Kalimantan  
503 peatland forest. *Journal of Tropical Forest Science* 2018; 30: 570-575.
- 504 Bang-Andreasen T, Nielsen JT, Voriskova J, Heise J, Rønn R, Kjøller R, et al. Wood Ash  
505 Induced pH Changes Strongly Affect Soil Bacterial Numbers and Community  
506 Composition. *Frontiers in Microbiology* 2017; 8.

507 Basiron Y. Palm oil production through sustainable plantations. *Eur. J. Lipid Sci. Technol.*  
508 2007; 109: 289-295.

509 Beest, Petrone, Nwaishi, Waddington, Macrae. Increased peatland nutrient availability  
510 following the Fort McMurray Horse River wildfire. *Diversity* 2019; 11: 142.

511 Beliveau A, Davidson R, Lucotte M, Do Canto Lopes LO, Paquet S, Vasseur C. Early effects  
512 of slash-and-burn cultivation on soil physicochemical properties of small-scale farms  
513 in the Tapajós region, Brazilian Amazon. *The Journal of Agricultural Science* 2015;  
514 153: 205-221.

515 Bista P, Ghimire R, Machado S, Pritchett L. Biochar Effects on Soil Properties and Wheat  
516 Biomass vary with Fertility Management. *Agronomy* 2019; 9: 623.

517 Capooci M, Barba J, Seyfferth AL, Vargas R. Experimental influence of storm-surge salinity  
518 on soil greenhouse gas emissions from a tidal salt marsh. *Science of The Total*  
519 *Environment* 2019; 686: 1164-1172.

520 Chazdon RL. Ecology - Tropical forests - Log 'em or leave 'em? *Science* 1998; 281: 1295-  
521 1296.

522 Cheong KH, Ngiam NJ, Morgan GG, Pek PP, Tan BYQ, Lai JW, et al. Acute Health Impacts  
523 of the Southeast Asian Transboundary Haze Problem-A Review. *International Journal*  
524 *of Environmental Research and Public Health* 2019; 16: 18.

525 Chungu D, Ng'andwe P, Mubanga H, Chileshe F. Fire alters the availability of soil nutrients  
526 and accelerates growth of *Eucalyptus grandis* in Zambia. *Journal of Forestry Research*  
527 2019.

528 Corley RHV. How much palm oil do we need? *Environmental Science & Policy* 2009; 12: 134-  
529 139.

530 Couwenberg J. Methane emissions from peat soils (organic soils, histosols) Facts, MRV-  
531 ability, emission factors., Bonn, Germany, 2009.

532 Couwenberg J, Dommain R, Joosten H. Greenhouse gas fluxes from tropical peatlands in  
533 south-east Asia. *Global Change Biology* 2010; 16: 1715-1732.

534 Danilova OV, Belova SE, Kulichevskaya IS, Dedysh SN. Decline of activity and shifts in the  
535 methanotrophic community structure of an ombrotrophic peat bog after wildfire.  
536 *Microbiology* 2015; 84: 624-629.

537 Davidson SJ, Van Beest C, Petrone R, Strack M. Wildfire overrides hydrological controls on  
538 boreal peatland methane emissions. *Biogeosciences* 2019; 16: 2651-2660.

539 Dennis RA, Mayer J, Applegate G, Chokkalingam U, Colfer CJP, Kurniawan I, et al. Fire,  
540 people and pixels: Linking social science and remote sensing to understand underlying  
541 causes and impacts of fires in Indonesia. *Human Ecology* 2005; 33: 465-504.

542 Dhandapani S. Biodiversity Loss Associated with Oil Palm Plantations in Malaysia: Serving  
543 the Need versus Saving the Nature. 4th International Conference On Biodiversity. 4:3.  
544 *Journal of Ecosystem and Ecography*, Las Vegas, USA, 2015.

545 Dhandapani S, Girkin NT, Evers S, Ritz K, Sjögersten S. Is intercropping an environmentally-  
546 wise alternative to established oil palm monoculture in tropical peatlands? *Frontiers in*  
547 *Forests and Global Change* 2020; 3:70.

548 Dhandapani S, Ritz K, Evers S, Sjögersten S. Environmental impacts as affected by different  
549 oil palm cropping systems in tropical peatlands. *Agriculture, Ecosystems &*  
550 *Environment* 2019a; 276: 8-20.

551 Dhandapani S, Ritz K, Evers S, Sjögersten S. GHG emission under different cropping systems  
552 in some Histosols of Malaysia. *Geoderma Regional* 2019b; 18: e00229.

553 Dhandapani S, Ritz K, Evers S, Yule CM, Sjögersten S. Are secondary forests second-rate?  
554 Comparing peatland greenhouse gas emissions, chemical and microbial community  
555 properties between primary and secondary forests in Peninsular Malaysia. *Science of*  
556 *The Total Environment* 2019c; 655: 220-231.

557 Evers S, Yule CM, Padfield R, O'Reilly P, Varkkey H. Keep wetlands wet: the myth of  
558 sustainable development of tropical peatlands - implications for policies and  
559 management. *Global Change Biology* 2017; 23: 534-549.

560 Fiedler S. In Situ Long-Term-Measurement of Redox Potential in Redoximorphic Soils. In:  
561 Schüring J, Schulz HD, Fischer WR, Böttcher J, Duijnsveld WHM, editors. *Redox:  
562 Fundamentals, Processes and Applications*. Springer Berlin Heidelberg, Berlin,  
563 Heidelberg, 2000, pp. 81-94.

564 Gay-Des-Combes JM, Sanz Carrillo C, Robroek BJM, Jassey VEJ, Mills RTE, Arif MS, et al.  
565 Tropical soils degraded by slash-and-burn cultivation can be recultivated when  
566 amended with ashes and compost. *Ecology and Evolution* 2017; 7: 5378-5388.

567 Giardina CP, Sanford RL, Døckersmith IC, Jaramillo VJ. The effects of slash burning on  
568 ecosystem nutrients during the land preparation phase of shifting cultivation. *Plant and  
569 Soil* 2000; 220: 247-260.

570 Global Environmental Centre G. *Integrated Management Plan for North Selangor Peat Swamp  
571 Forest 2014-2023 for Selangor State Forestry Department*, 2014, pp. 183.

572 Godfray H CJ, Beddington JR, Crute IR, Haddad L, Lawrence D, Muir JF, et al. *Food Security:  
573 The Challenge of Feeding 9 Billion People*. *Science* 2010; 327: 812-818.

574 González-Pérez JA, González-Vila FJ, Almendros G, Knicker H. The effect of fire on soil  
575 organic matter—a review. *Environment International* 2004; 30: 855-870.

576 Hansen MC, Potapov PV, Moore R, Hancher M, Turubanova SA, Tyukavina A, et al. High-  
577 Resolution Global Maps of 21st-Century Forest Cover Change. *Science* 2013; 342:  
578 850-853.

579 Hatano R, Toma Y, Hamada Y, Arai H, Susilawati HL, Inubushi K. Methane and nitrous oxide  
580 emissions from tropical peat soil. In: Osaki M, Tsuji N, editors. *Tropical Peatland  
581 Ecosystems*. Springer, Tokyo, 2016.

582 Heydari M, Rostamy A, Najafi F, Dey DC. Effect of fire severity on physical and biochemical  
583 soil properties in Zagros oak (*Quercus brantii* Lindl.) forests in Iran. *Journal of Forestry*  
584 *Research* 2017; 28: 95-104.

585

586 Hu Y, Christensen EG, Amin HMF, Smith TEL, Rein G. Experimental study of moisture  
587 content effects on the transient gas and particle emissions from peat fires. *Combustion*  
588 *and Flame* 2019; 209: 408-417.

589 Hu YQ, Fernandez-Anez N, Smith TEL, Rein G. Review of emissions from smouldering peat  
590 fires and their contribution to regional haze episodes. *International Journal of Wildland*  
591 *Fire* 2018; 27: 293-312.

592 Islam MS, Pei YH, Mangharam S. Trans-Boundary Haze Pollution in Southeast Asia:  
593 Sustainability through Plural Environmental Governance. *Sustainability* 2016; 8.

594 Kennard DK, Gholz HL. Effects of high- and low-intensity fires on soil properties and plant  
595 growth in a Bolivian dry forest. *Plant and Soil* 2001; 234: 119-129.

596 Kettunen A, Kaitala V, Alm J, Silvola J, Nykänen H, Martikainen PJ. Cross-correlation  
597 analysis of the dynamics of methane emissions from a boreal peatland. *Global*  
598 *Biogeochemical Cycles* 1996; 10: 457-471.

599 Khor WC, Rabaey K, Vervaeren H. Low temperature calcium hydroxide treatment enhances  
600 anaerobic methane production from (extruded) biomass. *Bioresource Technology*  
601 2015; 176: 181-188.

602 Kleinman PJA, Pimentel D, Bryant RB. The ecological sustainability of slash-and-burn  
603 agriculture. *Agriculture, Ecosystems & Environment* 1995; 52: 235-249.

604 Koh LP, Wilcove DS. Is oil palm agriculture really destroying tropical biodiversity? *Conserv.*  
605 *Lett.* 2008; 1: 60-64.

606 Kong S-H, Loh S-K, Bachmann RT, Rahim SA, Salimon J. Biochar from oil palm biomass: A  
607 review of its potential and challenges. *Renewable and Sustainable Energy Reviews*  
608 2014; 39: 729-739.

609 Kutiel P, Naveh Z. The effect of fire on nutrients in a pine forest soil. *Plant and Soil* 1987; 104:  
610 269-274.

611 Lam WY, Kulak M, Sim S, King H, Huijbregts MAJ, Chaplin-Kramer R. Greenhouse gas  
612 footprints of palm oil production in Indonesia over space and time. *Science of the Total*  
613 *Environment* 2019; 688: 827-837.

614 Lar JS, Li R, Li X. The Influence of Calcium and Iron Supplementation on the Methane Yield  
615 of Biogas Treating Dairy Manure. *Energy Sources, Part A: Recovery, Utilization, and*  
616 *Environmental Effects* 2010; 32: 1651-1658.

617 Liew RK, Nam WL, Chong MY, Phang XY, Su MH, Yek PNY, et al. Oil palm waste: An  
618 abundant and promising feedstock for microwave pyrolysis conversion into good  
619 quality biochar with potential multi-applications. *Process Safety and Environmental*  
620 *Protection* 2018; 115: 57-69.

621 Luskin MS, Potts MD. Microclimate and habitat heterogeneity through the oil palm lifecycle.  
622 *Basic and Applied Ecology* 2011; 12: 540-551.

623 Miettinen J, Hooijer A, Tollenaar D, Malins C, Vernimmen R, Shi C, et al. Historical Analysis  
624 and Projection of Oil Palm Plantation Expansion on Peatland in Southeast Asia.  
625 International Council on Clean Transportation, Washington DC, 2012.

626 Murakami M, Furukawa Y, Inubushi K. Methane production after liming to tropical acid peat  
627 soil. *Soil Science and Plant Nutrition* 2005; 51: 697-699.

628 Murdiyarso D, Hergoualc'h K, Verchot LV. Opportunities for reducing greenhouse gas  
629 emissions in tropical peatlands. *Proceedings of the National Academy of Sciences of*  
630 *the United States of America* 2010; 107: 19655-19660.

631 Murdiyarso D, Lilleskov E, Kolka R. Tropical peatlands under siege: the need for evidence-  
632 based policies and strategies. *Mitigation and Adaptation Strategies for Global Change*  
633 2019; 24: 493-505.

634 Myllyntaus T, Minna H, Jan K. Sustainability in Danger?: Slash-and-Burn Cultivation in  
635 Nineteenth-Century Finland and Twentieth-Century Southeast Asia. *Environmental*  
636 *History* 2002; 7: 267-302.

637 Niedermeier A, Robinson JS. Hydrological controls on soil redox dynamics in a peat-based,  
638 restored wetland. *Geoderma* 2007; 137: 318-326.

639 Ohlemiller TJ. Modeling of smoldering combustion propagation. *Progress in Energy and*  
640 *Combustion Science* 1985; 11: 277-310.

641 Page SE, Siegert F, Rieley JO, Boehm HDV, Jaya A, Limin S. The amount of carbon released  
642 from peat and forest fires in Indonesia during 1997. *Nature* 2002; 420: 61-65.

643 Posa MRC, Wijedasa LS, Corlett RT. Biodiversity and Conservation of Tropical Peat Swamp  
644 Forests. *BioScience* 2011; 61: 49-57.

645 Qing-Yu J, Wen-Ying Y, Li Z, Ri-Hong W, Yan-bing X, Ni-Na C. Methane emissions from  
646 typical paddy fields in Liaohe Plain and Sanjiang Plain, Northeast China.  
647 *Environmental Research Communications* 2019; 1: 011006.

648 Rashid AHAM, Hamzah KA, Joseph KT. Land use change in Malaysia. Reports from the  
649 technical panels of the 2nd green house gas. Roundtable on Sustainable palm oil, 2013.

650 Reijnders L, Huijbregts MAJ. Palm oil and the emission of carbon-based greenhouse gases.  
651 *Journal of Cleaner Production* 2008; 16: 477-482.

652 Rein G. Smouldering Fires and Natural Fuels. *Fire Phenomena and the Earth System: An*  
653 *Interdisciplinary Guide to Fire Science*, 2013, pp. 15-33.

654 Ren F, Zhang X, Liu J, Sun N, Wu L, Li Z, et al. A synthetic analysis of greenhouse gas  
655 emissions from manure amended agricultural soils in China. *Scientific Reports* 2017;  
656 7: 8123.

657 Sazawa K, Wakimoto T, Fukushima M, Yustiawati Y, Syawal MS, Hata N, et al. Impact of  
658 Peat Fire on the Soil and Export of Dissolved Organic Carbon in Tropical Peat Soil,  
659 Central Kalimantan, Indonesia. *ACS Earth and Space Chemistry* 2018; 2: 692-701.

660 Scharenbroch BC, Nix B, Jacobs KA, Bowles ML. Two decades of low-severity prescribed  
661 fire increases soil nutrient availability in a Midwestern, USA oak (*Quercus*) forest.  
662 *Geoderma* 2012; 183-184: 80-91.

663 Schmidt JH. Life cycle assessment of five vegetable oils. *Journal of Cleaner Production* 2015;  
664 87: 130-138.

665 Setia R, Gottschalk P, Smith P, Marschner P, Baldock J, Setia D, et al. Soil salinity decreases  
666 global soil organic carbon stocks. *Science of The Total Environment* 2013; 465: 267-  
667 272.

668 Shevade VS, Loboda TV. Oil palm plantations in Peninsular Malaysia: Determinants and  
669 constraints on expansion. *PLOS ONE* 2019; 14: e0210628.

670 Sjögersten S, Cheesman AW, Lopez O, Turner BL. Biogeochemical processes along a nutrient  
671 gradient in a tropical ombrotrophic peatland. *Biogeochemistry* 2011; 104: 147-163.

672 Smith TEL, Evers S, Yule CM, Gan JY. In Situ tropical peatland fire emission factors and their  
673 variability, as determined by field measurements in Peninsula Malaysia. *Global  
674 Biogeochemical Cycles* 2018; 32: 18-31.

675 Søndergaard M. Redox Potential. In: Likens GE, editor. *Encyclopedia of Inland Waters*.  
676 Academic Press, Oxford, 2009, pp. 852-859.

677 Tan KT, Lee KT, Mohamed AR, Bhatia S. Palm oil: Addressing issues and towards sustainable  
678 development. *Renewable and Sustainable Energy Reviews* 2009; 13: 420-427.

679 Turetsky MR, Benscoter B, Page S, Rein G, van der Werf GR, Watts A. Global vulnerability  
680 of peatlands to fire and carbon loss. *Nature Geoscience* 2015; 8: 11-14.

681 Uda SK, Hein L, Sumarga E. Towards sustainable management of Indonesian tropical  
682 peatlands. *Wetlands Ecology and Management* 2017; 25: 683-701.

683 Ulery AL, Graham RC, Amrhein C. Wood-ash composition and soil pH following intense  
684 burning. *Soil Science* 1993; 156: 358-364.

685 Van Reuler H, Janssen BH. Nutrient fluxes in the shifting cultivation system of south-west  
686 Côte d'Ivoire. *Plant and Soil* 1993; 154: 169-177.

687 Varkkey H. Patronage politics, plantation fires and transboundary haze. *Environmental*  
688 *Hazards* 2013; 12: 200-217.

689 Varkkey H. The haze problem in Southeast Asia: palm oil and patronage. Abingdon:  
690 Routledge, 2016.

691 Veraart AJ, Steenbergh AK, Ho A, Kim SY, Bodelier PLE. Beyond nitrogen: The importance  
692 of phosphorus for CH<sub>4</sub> oxidation in soils and sediments. *Geoderma* 2015; 259-260:  
693 337-346.

694 Visconti F, De Paz JM. Electrical Conductivity Measurements in Agriculture: The Assessment  
695 of Soil Salinity. InTech, 2016.

696 Wang G, Yu X, Bao K, Xing W, Gao C, Lin Q, et al. Effect of fire on phosphorus forms in  
697 Sphagnum moss and peat soils of ombrotrophic bogs. *Chemosphere* 2015; 119: 1329-  
698 1334.

699 Weslien P, Kasimir Klemetsson Å, Börjesson G, Klemetsson L. Strong pH influence on N<sub>2</sub>O  
700 and CH<sub>4</sub> fluxes from forested organic soils. *European Journal of Soil Science* 2009;  
701 60: 311-320.

702 Wiggins EB, Czimczik CI, Santos GM, Chen Y, Xu XM, Holden SR, et al. Smoke radiocarbon  
703 measurements from Indonesian fires provide evidence for burning of millennia-aged

704 peat. Proceedings of the National Academy of Sciences of the United States of America  
705 2018; 115: 12419-12424.

706 World Bank. The cost of fire : an economic analysis of Indonesia's 2015 fire crisis (English)  
707 . Indonesia sustainable landscapes knowledge; note no. 1. World Bank Group,  
708 Washington D.C., 2016.

709 Wright EL, Black CR, Turner BL, Sjögersten S. Environmental controls of temporal and spatial  
710 variability in CO<sub>2</sub> and CH<sub>4</sub> fluxes in a neotropical peatland. Global Change Biology  
711 2013; 19: 3775-3789.

712 Yan N, Marschner P, Cao W, Zuo C, Qin W. Influence of salinity and water content on soil  
713 microorganisms. International Soil and Water Conservation Research 2015; 3: 316-  
714 323.

715 Zaccone C, Rein G, D'Orazio V, Hadden RM, Belcher CM, Miano TM. Smouldering fire  
716 signatures in peat and their implications for palaeoenvironmental reconstructions.  
717 Geochimica et Cosmochimica Acta 2014; 137: 134-146.

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## 720 **Figure Captions**

721 **Figure 1:** Location of the study site.

722 **Figure 2:** Site pictures - a) picture of the study site during smouldering fire July 2018 b) picture  
723 of the site during the dry season measurements, one month after the fire event, in August 2018  
724 c) picture of the site during the wet season measurements, five months after the fire event, in  
725 December 2018.

726 **Figure 3:** Effect of burnt and non-burnt region, and season upon a) CO<sub>2</sub> emissions b) CH<sub>4</sub>  
727 emissions, during dry and wet season. Bars denote mean values (for dry season n=20 each  
728 region; for wet season n=10 each region). Whiskers denote standard errors.

729 **Figure 4:** Effect of burnt and non-burnt region, and season upon a) Peat surface temperature  
730 b) gravimetric moisture c) electrical conductivity d) redox potential e) pH during dry and wet  
731 season. Bars denote mean values (for dry season n=20 each region; for wet season n=10 each  
732 region). Whiskers denote standard errors.

733 **Figure 5:** Effect of burnt and non-burnt region, and season upon peat a) C content b) N content  
734 c) C:N ratio d) Mg concentrations e) P concentrations f) S concentrations f) K concentrations  
735 g) Ca concentrations during dry and wet season. Bars denote mean values (for dry season n=20  
736 each region; for wet season n=10 each region). Whiskers denote standard errors.

737 **Figure 6:** Effect of burnt and non-burnt region, and depth upon a) pH b) gravimetric moisture  
738 c) electrical conductivity d) redox potential during dry season. Points denote mean values (n=3  
739 for each region). Whiskers denote standard errors.

740 **Figure 7:** Effect of burnt and non-burnt region, and depth upon peat a) C content b) N content  
741 c) C:N ratio d) Mg concentrations e) P concentrations f) S concentrations f) K concentrations  
742 g) Ca concentrations during dry season. Bars denote mean values (n=3 for each region).  
743 Whiskers denote standard errors.

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