



## LJMU Research Online

**Razzolini, NL, Vila, B, Castanera, D, Falkingham, PL, Barco, JL, Canudo, JI, Manning, PL and Galobart, A**

**Intra-trackway morphological variations due to substrate consistency: the El Frontal dinosaur tracksite (Lower Cretaceous, Spain).**

<http://researchonline.ljmu.ac.uk/id/eprint/599/>

### Article

**Citation** (please note it is advisable to refer to the publisher's version if you intend to cite from this work)

**Razzolini, NL, Vila, B, Castanera, D, Falkingham, PL, Barco, JL, Canudo, JI, Manning, PL and Galobart, A (2014) Intra-trackway morphological variations due to substrate consistency: the El Frontal dinosaur tracksite (Lower Cretaceous, Spain). PLoS One. 9 (4). ISSN 1932-6203**

LJMU has developed [LJMU Research Online](http://researchonline.ljmu.ac.uk/) for users to access the research output of the University more effectively. Copyright © and Moral Rights for the papers on this site are retained by the individual authors and/or other copyright owners. Users may download and/or print one copy of any article(s) in LJMU Research Online to facilitate their private study or for non-commercial research. You may not engage in further distribution of the material or use it for any profit-making activities or any commercial gain.

The version presented here may differ from the published version or from the version of the record. Please see the repository URL above for details on accessing the published version and note that access may require a subscription.

For more information please contact [researchonline@ljmu.ac.uk](mailto:researchonline@ljmu.ac.uk)

<http://researchonline.ljmu.ac.uk/>



# Intra-Trackway Morphological Variations Due to Substrate Consistency: The El Frontal Dinosaur Tracksite (Lower Cretaceous, Spain)

Novella L. Razzolini<sup>1\*</sup>, Bernat Vila<sup>1,2</sup>, Diego Castanera<sup>2</sup>, Peter L. Falkingham<sup>3,4</sup>, José Luis Barco<sup>5</sup>, José Ignacio Canudo<sup>2</sup>, Phillip L. Manning<sup>6,7</sup>, Àngel Galobart<sup>1</sup>

**1** Institut Català de Paleontologia Miquel Crusafont, Carrer de l'Escola Industrial, 23, Sabadell, Barcelona, Catalonia, **2** Grupo Aragosaurus-IUCA, Departamento de Paleontología, Facultad de Ciencias, Universidad de Zaragoza, Zaragoza, Spain, **3** Department of Comparative Biomedical Sciences, Structure and Motion Laboratory, Royal Veterinary College, London, United Kingdom, **4** Department of Ecology and Evolutionary Biology, Division of Biology and Medicine, Brown University, Providence, Rhode Island, United States of America, **5** Paleoymás S.L. Pol. Empresarium, Zaragoza, Spain, **6** Department of Earth and Environmental Science, University of Pennsylvania, Philadelphia, Pennsylvania, United States of America, **7** School of Earth, Atmospheric and Environmental Sciences, University of Manchester, Manchester, United Kingdom

## Abstract

An ichnological and sedimentological study of the El Frontal dinosaur tracksite (Early Cretaceous, Cameros basin, Soria, Spain) highlights the pronounced intra-trackway variation found in track morphologies of four theropod trackways. Photogrammetric 3D digital models revealed various and distinct intra-trackway morphotypes, which reflect changes in footprint parameters such as the pace length, the track length, depth, and height of displacement rims. Sedimentological analyses suggest that the original substrate was non-homogenous due to lateral changes in adjoining microfacies. Multidata analyses indicate that morphological differences in these deep and shallow tracks represent a part of a continuum of track morphologies and geometries produced by a gradient of substrate consistencies across the site. This implies that the large range of track morphologies at this site resulted from similar trackmakers crossing variable facies. The trackways at the El Frontal site present an exemplary case of how track morphology, and consequently potential ichnotaxa, can vary, even when produced by a single trackmaker.

**Citation:** Razzolini NL, Vila B, Castanera D, Falkingham PL, Barco JL, et al. (2014) Intra-Trackway Morphological Variations Due to Substrate Consistency: The El Frontal Dinosaur Tracksite (Lower Cretaceous, Spain). PLoS ONE 9(4): e93708. doi:10.1371/journal.pone.0093708

**Editor:** Andrew A. Farke, Raymond M. Alf Museum of Paleontology, United States of America

**Received:** November 28, 2013; **Accepted:** March 5, 2014; **Published:** April 3, 2014

**Copyright:** © 2014 Razzolini et al. This is an open-access article distributed under the terms of the Creative Commons Attribution License, which permits unrestricted use, distribution, and reproduction in any medium, provided the original author and source are credited.

**Funding:** This paper is a contribution to the projects CGL2011-30069-C02-01 and CGL2010-16447, subsidized by the Ministerio de Economía y Competitividad of Spain. LiDAR data acquisition was funded by the Institut Català de Paleontologia "Miquel Crusafont". N. L. Razzolini acknowledges support from BES-2012-051847 subsidized by the Ministerio de Economía y Competitividad. B. Vila acknowledges support from Subprograma Juan de la Cierva (MICINN-JDC-2011). D. Castanera is the beneficiary of a grant from the Ministry of Education (AP2008-01340). P. L. Falkingham was supported by a Marie Curie International Outgoing Fellowship within the 7th European Framework Programme.

**Competing Interests:** José Luis Barco declares the affiliation to the company Paleoymás. There are no patents, products in development or marketed products to declare. This does not alter the authors' adherence to PLOS ONE policies on sharing data and materials.

\* E-mail: novella.razzolini@icp.cat

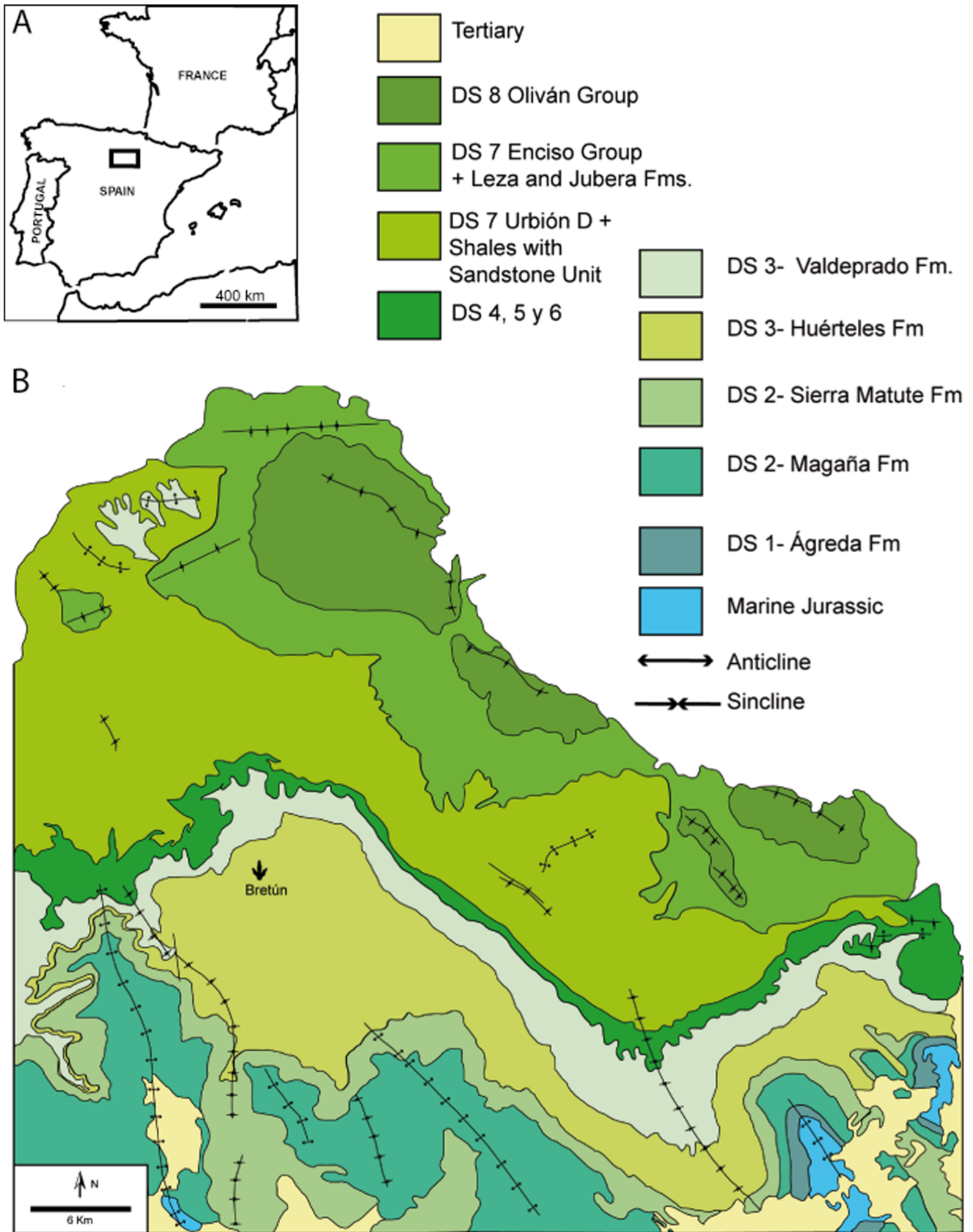
## Introduction

Track morphology is determined by both the trackmaker and the substrate characteristics [1–4]. Although it is widely accepted that the substrate is a major control in determining the final track morphology [1,2,5–7], studying this dynamic formation process is challenging given the fact that most foot-sediment and sediment-sediment interactions are highly complex, rapid and hidden from view [8–10]. Baird [11] stated that a trackway is not a simple record of anatomy; instead, it is a record of how a foot behaves under a distinct locomotory pattern as it makes contact with a particular substrate. The way in which sediments behave before, during, and after a track is formed and the subsequent processes that may further modify, enhance, or disguise a track has been much neglected [1,12]. Hence, to understand the formation and preservation of tracks, it is essential to understand the mechanics of soils and rheology [13–15].

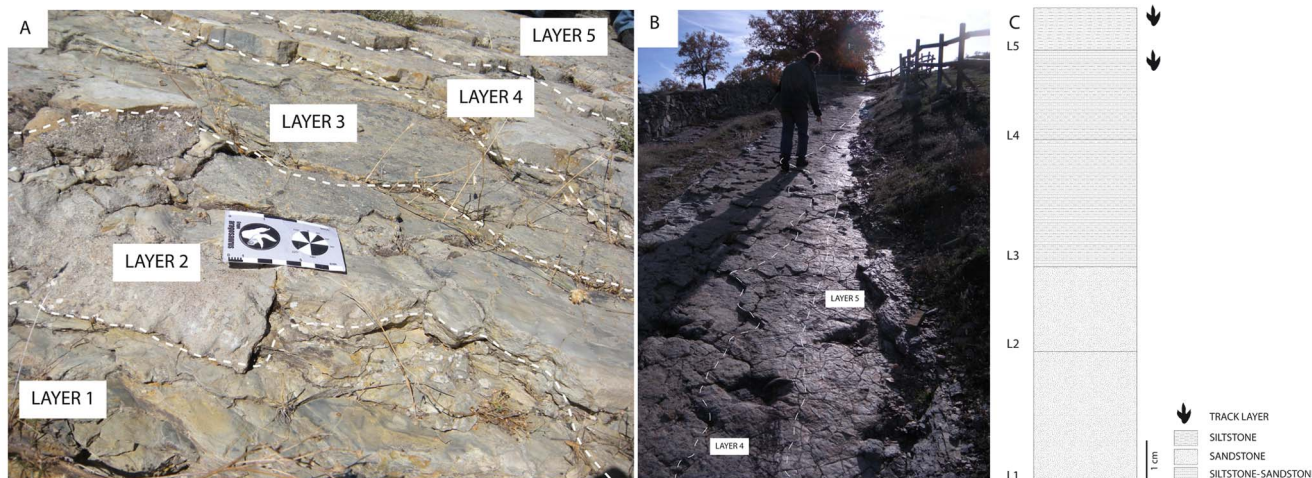
Traditionally, ichnology has primarily studied tracks and trackways as two-dimensional traces (e.g., [16,17]), rarely considering the substrates mechanics and prevailing condition at the time

a track-maker's foot made contact with a sediment. For ichnological analyses to be well founded, footprints must be documented by methods that avoid inaccurate representations of track morphology, which can distort or obscure potentially important data [18]. Recent advances agree that the foot's contact with a substrate can only be understood by taking a three-dimensional approach to explain track formation [1,5,19–21]. The variation in track morphology due to sediment consistency can be observed and quantified through the use of three-dimensional (3-D) technologies (i.e. using laser scanning or photogrammetry to show depth analyses and vertical cross sections) [22–25] with the intention to integrate quantitative analytical techniques with the traditional ichnotaxonomic definition. Light Detection And Range (LiDAR) techniques [22,23,26] together with photogrammetry methods [27] complement the classic ichnological data acquisition by providing accurate data on 3-D specimens.

The present study concentrates on the quantification of morphological variability of tridactyl dinosaur tracks documented at the El Frontal tracksite (Lower Cretaceous, NW Iberian Peninsula), which were briefly mentioned in previous works [28–



**Figure 1. Geographical and geological setting of the El Frontal tracksite (Bretun, Soria).** The location of Bretun locality within the Iberian Peninsula is inside the black square. The tracksite locates in DS-3 of the Huerteles Fm [32]. Geological map modified from Castanera et al. [33]. doi:10.1371/journal.pone.0093708.g001

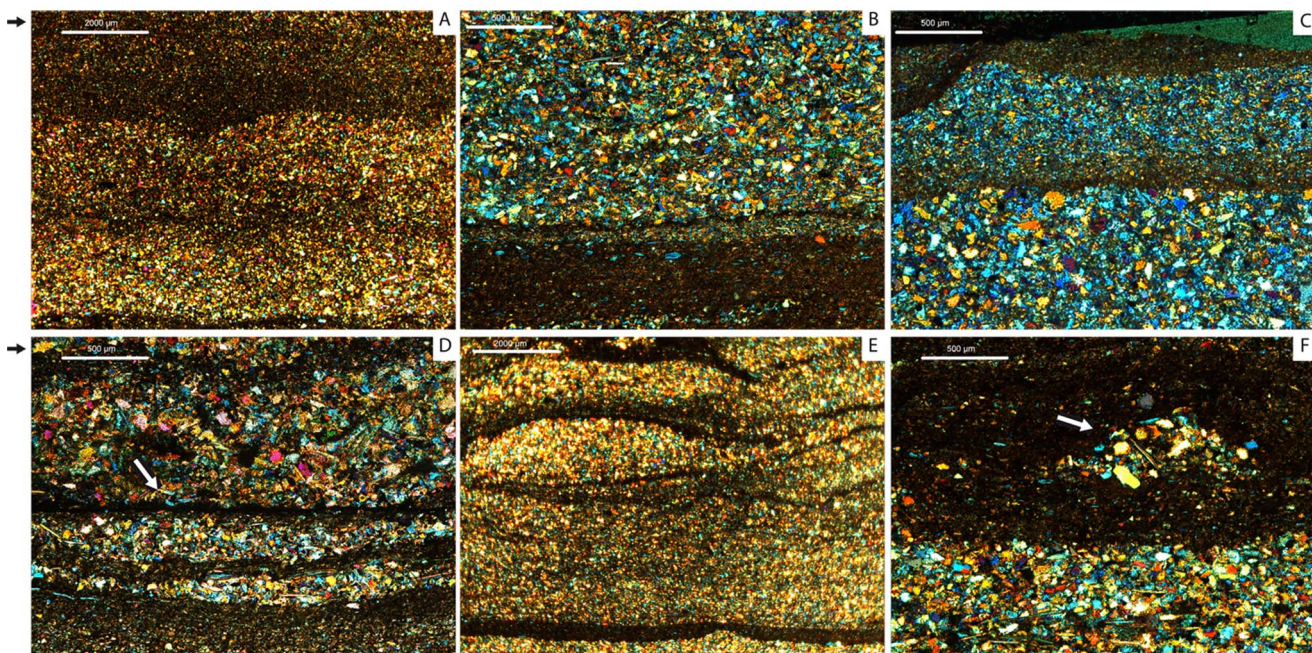


**Figure 2. Tracksite microlayers organization.** A) The El Frontal tracksite is composed of 5 different centimeter-thick layers that intercalate gray siltstones, limestone and sandy-siltstones. Scale bar equals 8 cm. B) El Frontal track layers 4 (penetrative tracks) and 5 (tracking surface), where all the studied tracks originated. When thin layer 5 is not preserved, tracks are found in level 4. C) Stratigraphical log of the five layers found in the El Frontal tracksite. Theropod tracks are found in layers 5 and 4. Scale bar equals 1 cm. doi:10.1371/journal.pone.0093708.g002

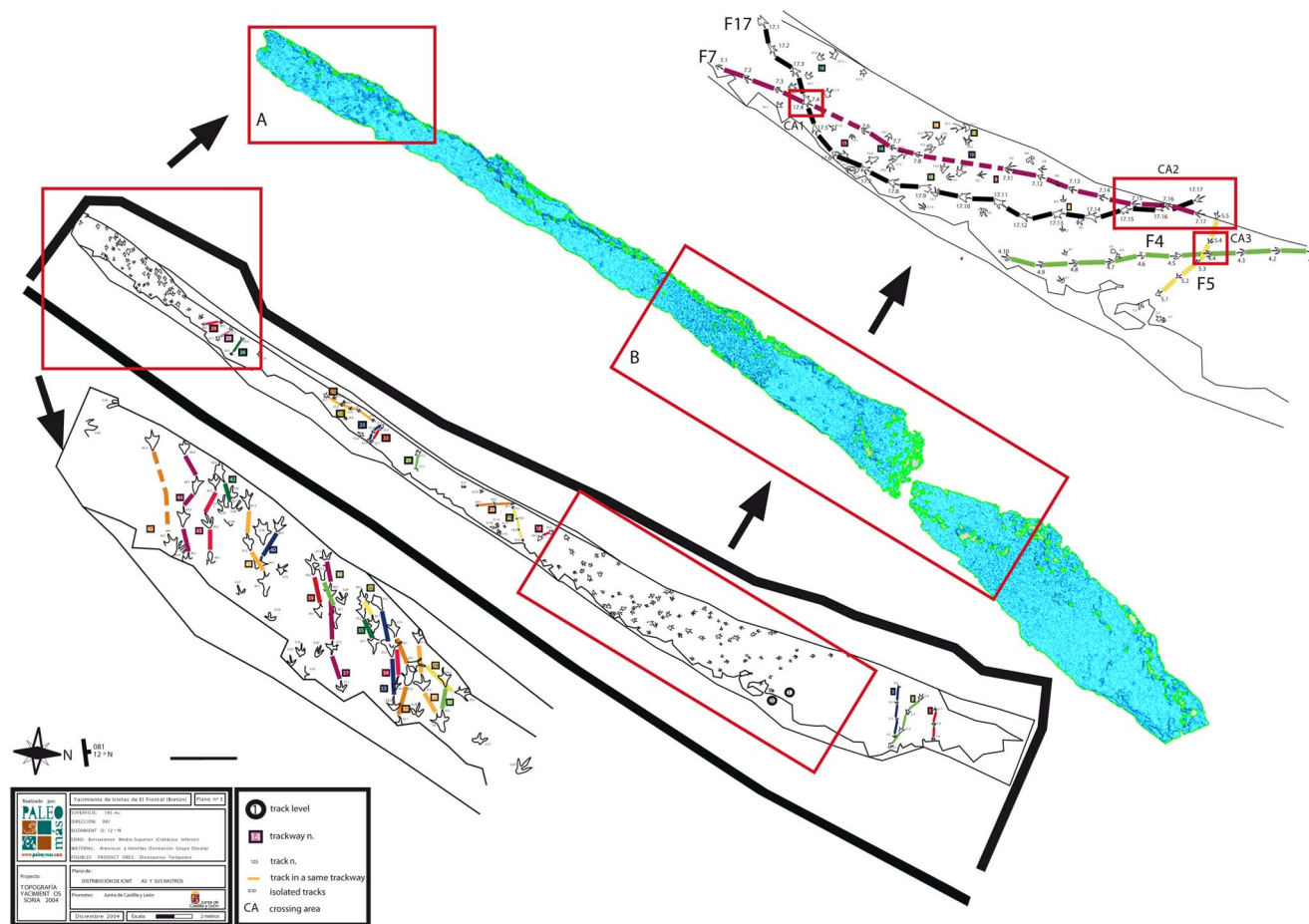
31] but never studied in detail. The aim of this work is to quantify the inter- and intra-trackway morphological variation *sensu* [10] recorded in different track shapes to underpin the variability in track morphology when track-maker is kept constant. This study will focus on four long trackways that are characterized by a range of track morphologies that are considered as indicators of rheological conditions.

**Geological Setting**

The El Frontal site is found in the Cameros Basin (Soria, Spain), which is located northwest of the Iberian range. The sedimentary infill of the Cameros basin was divided in eight depositional sequences, with deposits predominantly from continental environments [32] (Fig. 1). The sedimentation was dominantly continental as demonstrated by alluvial and lacustrine deposits [34], but includes some sporadic marine incursions [35–38]. The tracksite



**Figure 3. Thin sections IPS-82477a-d of layers 4 and 5 of the El Frontal tracksite.** A) Sandstone intercalated by siltstone-mudstone bands in which chlorite minerals are scarce (<5%); B) High quartz concentration (>60%) and scarce presence of clay minerals in mud bands; C) Sandstone-siltstone in which the grain size decreases from the bottom to the top; D) Sandstone intercalated by siltstone, mineral clays are abundant (>60%); E) Deformation structures (mud drapes); F) Deformation structure (symmetrical ripple). Black arrows indicate the top of the laminae, white arrows indicate deformation structures. Scale bars are 2000 μ for A, 500 μ for B, C, D, F and 2000 μ for E. doi:10.1371/journal.pone.0093708.g003



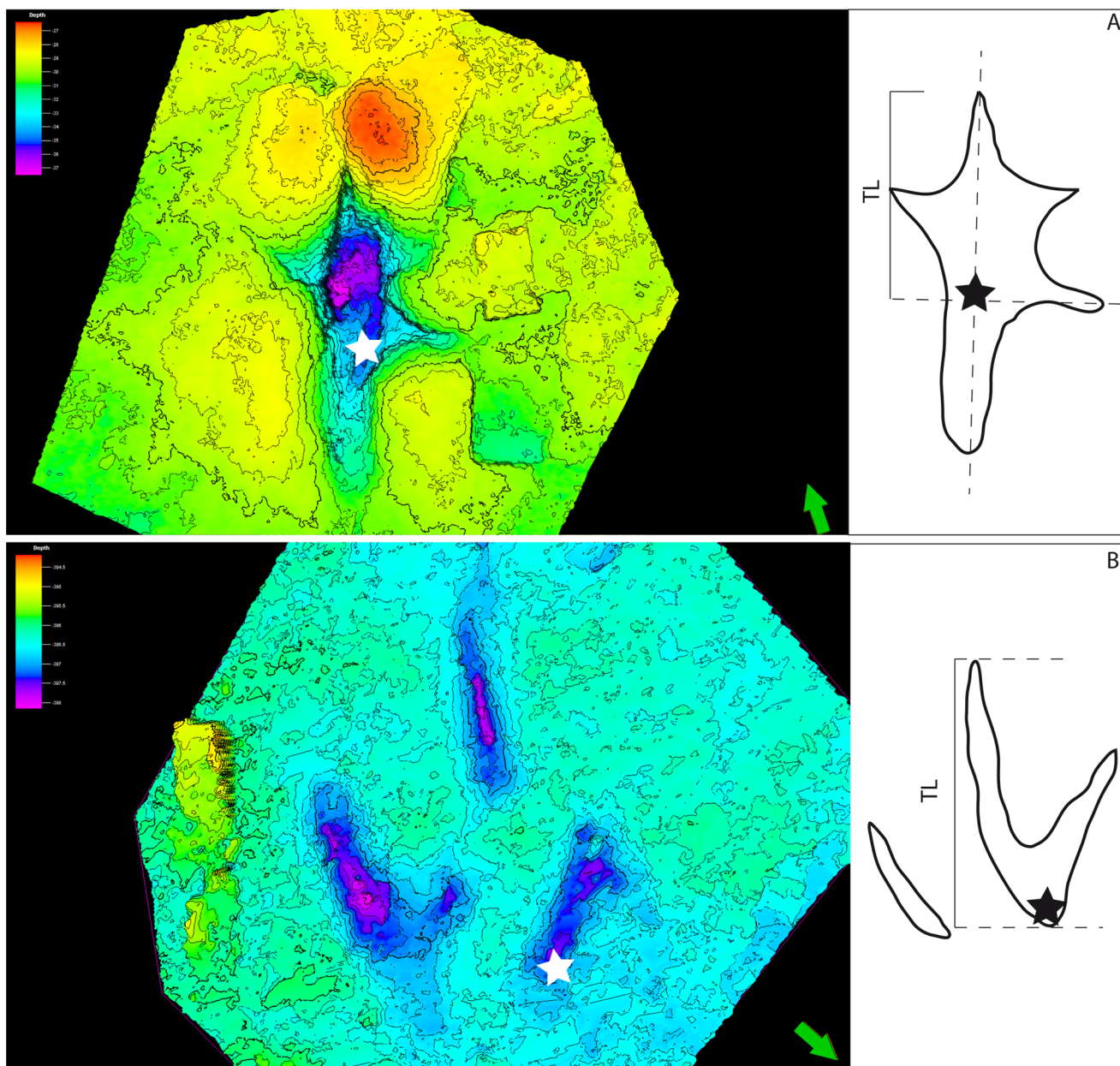
**Figure 4. Cartography and 3-D model of the El Frontal tracksite resulting from the LiDAR scanning (grey colour), modified from Barco et al. [44] and designed by Paleoymas SL.** In the red rectangles (A and B) are the details of the areas with the highest density of tracks. The studied area is detailed in the rectangle B. Studied trackways F17, F7, F5 and F4 are coloured respectively with black, pink, yellow and green. doi:10.1371/journal.pone.0093708.g004

falls in the DS-3 (depositional sequence, Fig. 1) and belongs to the Oncala Group. It is subdivided into the Huérteles (which includes the El Frontal tracksite) and Valdelprado formations and dates to the Berriasian [32,35,39,40]. The depositional sequence DS-3 follows a pattern of alluvial fans and lacustrine sediments that thin laterally (northwesterly) to fluvial and fluvio-lacustrine deposits [32].

Recent interpretation describes the Huérteles Formation as characterized by terrigenous sediments (fluvial system) in the western sector of the eastern Cameros Basin and by an increase in shallow, coastal, carbonate-sulphate water bodies to the east, implying that the connection of the Cameros basin with marine areas was much stronger [38] than previously considered [34]. A series of sedimentary structures that crop out near the El Frontal tracksite, (i.e. inclined heterolithic stratification, flaser, rhythmic alternations of sandstones and lutites, symmetrical ripples, mud-drapes) are indicative of a tidally-influenced fluvial-deltaic environment [37,38].

The El Frontal tracksite is 150 meters apart from the outcrops of the Fuente Lacorte tracksite reported by Aguirrezabala and Viera [28] and Sanz et al. [29]. The latter is stratigraphically lower with respect to the studied locality. The lithology of the El Frontal tracksite is composed of 5 different layers that include intercalated organic rich gray siltstones mudstones and sandy-siltstones (Fig. 2A–C). In detail, trackways and isolated tracks are produced

and impressed in layer 5 (tracking surface). Layer 5 is a 1 cm-thick siltstone with occasional mud cracks (Fig. 2C) which sometimes is not preserved, and tracks and trackways are found as undertracks in the underlying layer, layer 4. This is characterized by a 2–3 cm thick sandstone-siltstone (Fig. 2B–C). The first set of laminas (Fig. 3A–C) observed at different areas of the tracksite reveals that layer 4 and 5 are composed of quartz (>60%), and minor abundance of phyllosilicates, and chlorite minerals (Fig. 3A–C). It has a grain-supported fabric with quartz ranging from fine to medium size, yielding a moderately sorted composition. The chlorite minerals (<5%) and other planar minerals are very scarce in the mudstone band in the clay matrix. The second set of laminas including layer 4 and 5 was collected not far from the first (Fig. 3D–F) and is composed of sandstone intercalated with mudstone. In these laminas, sedimentary structures, such as mud drapes (Fig. 3D, 3E) and symmetric ripples (Fig. 3F) are observed. These are characterized by a higher percentage of chlorite minerals (>60%) that concentrate in the mudstone band in the clay matrix. Mud drape structures form when a sediment undergoes intermittent flows, leading to alternating sand and mud layers [37]. Symmetrical ripples are formed when a horizontal oscillation generates wave ripples formed by rolling grains in shallow water [41], and they are commonly found associated with mud drapes [37].



**Figure 5. Standards for measuring track length (TL) and depths (black and white stars).** A) TL excludes the elongated metatarsal impression, and the depth is taken approximately where the phalanx 1 of metatarsal IV should be. B) In this case there is no metatarsal impression and the TL is easier to measure. Depth is taken in the same point for every track. Color scale green indicates the track layer, purple is the deepest point recorded and red is the highest point recorded.  
doi:10.1371/journal.pone.0093708.g005

## Materials and Methods

A complete digital model of the track-bearing outcrop was generated using a RIEGL LMS-Z420i long range 3D laser scanner capable of 5–10 mm resolution [22,23,42] (for three-dimensional El Frontal tracksite caption see Appendix S1). The three-dimensional surface of the tracksite El Frontal is available as a polygon file in the Supplementary Information. This overview scan was complemented with close-range photogrammetric models [27] of individual tracks (Appendix S2), produced from 10 to 20 photographs per track and processed using VisualSFM (<http://ccwu.me/vsfm/>) [43].

Four trackways (F17, F7, F5, and F4) spanning the site were studied in detail, comprising a total of 49 tracks (17, 17, 5, and 10 tracks from the respective trackways) (Fig. 4). These trackways were chosen for their high morphological variability and their proximity to each other, with the aim of reflecting any effect of spatial variation in substrate consistency. For each track, several metrics were measured from the photogrammetric models using both ImageJ software and Schlumberger package Petrel: track length (TL, measured from tip of digit III, excluding metatarsal pad when present), track width (TW), interdigital angles ( $\text{II}^{\wedge}\text{III}$ ,  $\text{III}^{\wedge}\text{IV}$ ,  $\text{II}^{\wedge}\text{IV}$ ), displacement rim height (DR), maximum depth ( $D_{\text{max}}$ ), and depth of the metatarsal pad impression ( $D_{\text{mp}}$ ). The two

**Table 1.** Table with measurements taken for all the theropod trackways in the El Frontal.

TRACK	TL	TW	II^III	III^IV	II^IV	PL	SL
45.1	34.4	15.6	23.6	18.4		85.4	166.1
45.2	34	x	x	x		85.1	x
45.3	32.2	18.7	22.8	26.3		x	x
44.1	19.2	17.4	24.2	38.7		68.8	115.7
44.2	31.3	20	24.8	19		54.1	108.2
44.3	27.8	16.8	17.2	20.9		62	x
44.4	20.1	15.8	24.7	21.3		x	x
43.1	26	18.5	21.4	42.9		62.2	124.9
43.2	26	17.7	20.3	30		66	x
43.3	21	18.5	25.3	22.5		x	x
42.1	31	18.1	x	x		65.4	x
42.2	26.4	15	x	x		x	x
41.1	39	24.3	22.2	19.3		71.4	133
41.2	27.2	14.3	19.7	17		66.5	x
41.3	22	19.2	20.4	16.5		x	x
40.1	21	23.7	29.6	34.1		66.1	x
40.2	25	20.6	32.3	38.6		x	x
39.1	22.1	16.5	35.7	19.24		71.1	x
39.1	22.4	22.4	25.6	17.44		x	x
38.1	22.7	17.8	17	14.6		74.1	x
38.2	22	x	x	x		x	x
37.1	18.1	15.5	32.5	39.2		66.1	150.5
37.2	14.6	18	24.4	44		82.3	139.1
37.3	24.7	18.1	27.6	22.5		54.1	x
37.4	14.7	18.7	54.2	51.37		x	x
36.1	21	16.5	13.7	25.6		39.6	x
36.2	19.6	15.8	28.05	39.3		x	x
35.1	19.3	22.1	16.4	18.4		52	x
35.2	21.7	15	x	x		x	x
34.1	25	14	25.2	30.8		87.5	x
34.2	26.2	19.8	17.4	25.6		x	x
33.1	23.4	19.4	38.8	36.3		90.3	163.7
33.2	23.4	27.1	21.3	25.4		69.1	x
33.3	23.7	16.3	x	x		x	x
32.1	15.5	15.3	x	25.2		70.8	146.3
32.2	23	19.3	19.8	24.4		76.2	x
32.3	23.6	19	28.9	30.8		x	x
31.1	21.4	18.2	33.3	18.6		67.1	131
31.2	26.2	23.4	35.8	35.1		76.1	x
31.3	22.7	15.5	27.5	29.8		x	x
30.1	17.8	13.5	54.7	x		62.9	x
30.2	19.6	18.8	24.2	26.5		x	x
29.1	18.3	21.7	26.3	28.3		86.1	x
29.2	20.4	18.7	18.5	17.6		x	x
28.1	5.8	5.3	29.5	30.3		28	x
28.2	5.2	4.6	24.9	32.3		x	x
27.1	7.8	6.4	21.07	30.6		25.1	x
27.2	7.7	7.8	35.5	32.7		x	x

**Table 1. Cont.**

TRACK	TL	TW	II^III	III^IV	II^IV	PL	SL
26.1	6	5	26.5	26.7		28.6	x
26.2	5.1	6.2	23.52	30		x	x
25.1	21	27	68.2	77.6		97.2	202
25.2	28.4	21.8	45.3	36.5		107	x
25.3	28.4	20.2	43.8	36.5		100.6	x
25.4	25.2	19	40.2	42.5		x	x
24.1	6.5	5.3	24.9	26.8		26.8	x
24.2	5	5.5	40.8	28		x	x
23.1	5.5	6.3	31.2	39.1		16.1	29.5
23.2	6.9	5.8	27.2	39.2		15.3	x
23.3	3.8	5	42	42		x	x
22.1	5.8	5	20.5	20.1		23.7	x
22.2	5	5.6	25.3	25		x	x
21.1	7	6.3	36.7	28.5		26.3	x
21.2	7	6.1	12.9	26.1		x	x
20.1	8.2	5.4	20	27.1		27.4	53.5
20.2	6.4	5.9	30.4	36.8		26.3	x
20.3	5.4	4.3	40.2	47		x	x
19.1	4.6	4.4	60.2	43		18.4	34.6
19.2	4	4.1	34.4	30		17.2	35
19.3	4.2	3.5	52.3	30.5		18	x
19.4	5.3	4.8	32.4	40.8		x	x
18.1	4.6	4.4	17.7	28.6		x	x
18.2	7.3	7.9	55.7	49.1		20.1	x
17.1	30	14.7	27.3	25.6		91	180
17.2	29	13.2	22.1	19.9		95	182
17.3	27	17.7	26.7	32.7		96	177
17.4	31	14.0	19.7	22.1		85	175
17.5	29	14.3	21.5	30.3		90	
17.6	27	11.876	30	37.2	76.8	100	195
17.7	26	12.881	28.4	25.1	68.9	94	193
17.8	32	13.333	31.6	31.3	54.2	98	213
17.9	27	14.465	24.6	28.6	57.5	115	222
17.10	32	14.165	24.7	26.6	46.3	108	193
17.11	34	14.712	26.4	27.3	48.3	88	178
17.12	31	13.535	18.9	26	52.3	98	200
17.13	30	16.311	22.6	27	48.4	103	198
17.14	29	16.078	19.8	25.8	51.4	100	198
17.15	32	14.307	32	35.7	54	100	197
17.16	27	12.819	27.3	30.9	68.8	94	
17.17	30	16.297				100	
16.1	10.9	10.3	46.7	36.5		33.5	62.4
16.2	12	9.1	29.8	42		30.2	61.8
16.3	12.2	10.9	37.9	35.7		33.9	x
16.4	12.3	10.6	36	35		x	x
15.1	8	9.1	35.6	30.6		48.8	x
15.2	10.3	8.7	28.7	42		x	x
14.1	12.7	10.2	26.7	36		44.8	x
14.2	13.4	9.8	30.4	32.1		x	x

Table 1. Cont.

TRACK	TL	TW	II'III	III'IV	II'IV	PL	SL
13.1	15.5	12.1	23.9	38.8		40.4	81.2
13.2	15.7	12.1	22.8	29.1		42.4	x
13.3	15.5	14.6	20.1	33.2		x	x
12.1	18.3	11.8	29.1	32.2		44.8	91
12.2	18.8	11.8	33.2	20.2		48.8	95.2
12.3	16.2	11.4	26.2	21.6		49.4	x
12.4	20.5	10.8	27.5	35.1		x	x
11.1	12	11.3	42.7	34.4		62	119.4
11.2	13.4	11.8	32.8	31.3		58.7	119
11.3	13.8	10.5	32.9	34.6		60.8	x
11.4	13	11.3	30.7	33		x	x
10.1	18.2	15.2	25.4	38.4		50	x
10.2	16.4	14.3	31	35.2		x	x
9.1	12.3	10.2	30.4	45		48.6	92.4
9.2	14.2	10.2	34.4	33.5		44.6	93.1
9.3	14.2	11.5	30	40.4		49.6	x
9.4	11.7	11.3	31.9	44.2		x	x
8.1	8.9	9	28.5	46.6		36.1	80.6
8.2	8.7	9.3	36.6	47.1		45	68
8.3	11.3	9.3	48.7	48.7		24.9	x
8.4	10	8.8	38.6	46.5		x	x
7.1	20	11.0	29.2	27.9	57.6	88	185
7.2	21	13.9	26.3	39.6	59.3	98	193
7.3	23	13.9	27.3	42.4	65.9	100	
7.4	22	10.2	26.1	35.7	55.7	95	181
7.5	22	11.1	x			95	176
7.6	24	10.1	25	18.6	47.1	93	167
7.7	25	11.0	41.4	36.3	58.3	84	174
7.8	25	9.5	22.2	22	53.6	82	184
7.9	26	11.4				94	
7.10	22	10				93	196
7.11	25	11.2	45.9	17.2	59.6	103	201
7.12	24	13.0	27.9	29.6	51.3	102	195
7.13	21	10.7	32.9	26.2	66.7	99	193
7.14	22	12.0	30.6	31.7	59.6	97	197
7.15	25	10.9	38.3	52.6	64.7	100	195
7.16	19	10.7	32.5	38	73.4	99	
7.17	25	11.0				99	
5.1	17	12.0	29	41.4	77.7	74	142
5.2	19	12.1	31.6	48	78.6	68	134
5.3	16	8.7	39.4	43.2	80.3	69	135
5.4	18	10.4	28.5	39.7	70	66	134
5.5	22	11.7				65	
4.1	23	11.0	36.2	28.4	67.1	106	207
4.2	26	10.3	39.7	40.8	85.9	103	204
4.3	26	9.7	38.7	34.6	66.8	101	202
4.4	27	14.0	52.2	42.4	72.2	103	204
4.5	30	11.0	34	24.5	58.6	105	202
4.6	23	12.6	45	30.1	73.8	102	201

Table 1. Cont.

TRACK	TL	TW	II'III	III'IV	II'IV	PL	SL
4.7	26	11.0	31.2	23	51.3	100	204
4.8	22	12.6	37.6	35	70.7	100	200
4.9	24	12.2	36.4	38.7	61.8	103	
4.10	27	12.3	43	31.5	70.2	100	
3.1	5.6	5.5	42.1	45.8		25.5	50.5
3.2	5.4	5.5	44.3	43.5		25.2	50
3.3	6.2	6.1	30.2	53.3		24.9	x
3.4	2.9	5.3	38	25.3		x	x
2.1	42.08	33.05	57.12	34.74		103	
2.2	38.18	39.13	61.8	44.25		113	210
2.3	46.87	41.1	51.03	51.25			
2.4							
1.1	7.5	6.8	32.2	36.6		22.8	44.9
1.2	8.6	6.6	34.6	37		21.8	x
1.3	7.3	5.6	25.2	37.3		x	x

TL (track length); TW (track width); II'III (interdigital angle between II'III); III'IV (interdigital angle between III'IV); II'IV (interdigital angle between II'IV taken for trackways 17, 7, 5 and 4); PL (Pace length); SL (Stride length). All measurements in Table 1 are in CM.  
doi:10.1371/journal.pone.0093708.t001

depth metrics were recorded, because many of the tracks show signs of post-formational sealing of the track walls around the digit impressions. Maximum depth is therefore interesting to note, but is of no use for comparisons between tracks (though it remains a useful metric in tracks where no sealing has occurred). The metatarsal pad, conversely, rarely suffers from such wall collapse due to the width of the impression, and so depth recordings from this homologous point between tracks can be comparatively informative. Unfortunately, the metatarsal pad is not always impressed. However, by recording both depth metrics where possible, an indication of the track morphology can be conveyed (Fig. 5A–B). Additionally, pace length (PL) and stride length (SL) were measured both in the field and using the whole-outcrop digital model. Statistical analyses on the 49 tracks refer to linear correlation and dispersion plots that interpolate track length (TL), depth (D) and displacement rim height (DR) parameters.

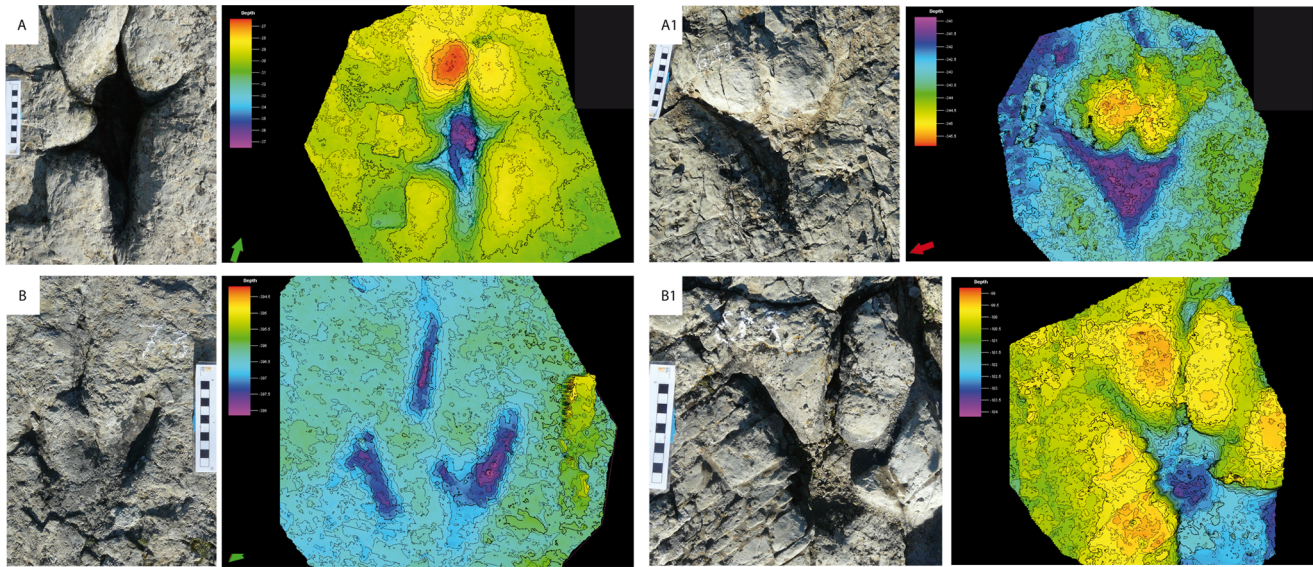
To quantify the substantial intra-trackway depth and length variations, four graphs for trackways F17, F7, F5 and F4 were built using TL, PL (left Y axis) and depth measurements (right y axis). A sedimentological analysis (4 thin sections in total, IPS-82477a-d housed at the Institut Català de Paleontologia “Miquel Crusafont”-ICP) for layers 4 (undertracks) and 5 (tracking surface) was undertaken to quantify lithology and mineral composition of the sediment. Pictures of the four polished thin sections (Fig. 3A–F) were taken using light microscopy via a Leica DM 2500 photomicroscope.

## Results

The El Frontal tracksite consists of a southwest-northeast orientated outcrop containing more than 200 tridactyl tracks and 45 trackways (see Table 1) [28–31], distributed along 185 m<sup>2</sup> surface area (Fig. 4). Track density is of more than one track/m<sup>2</sup>, although tracks are not homogeneously distributed (Fig. 4A–B).

We describe the position in the tracksite, spatial distance and possible interaction of trackways F17, F7, F4 and F5 with one





**Figure 6. Morphological characterization of the El Frontal tracks.** A) Morphotype A is track 17.2, A1) Variation of morphotype A, track 17.17, B) Morphotype B is track 7.3, B1) Variation of morphotype B, track 7.13. Color scale green and yellow indicates the track layer, purple is the deepest point recorded for depth and red is the highest point recorded for displacement rims.  
doi:10.1371/journal.pone.0093708.g006

another. At the northeastern edge of the outcrop, the first tracks of trackway F17 are separated from those of trackway F7 by one meter (Fig. 4B). F17 crosses with trackway F7 (crossing area 1, Fig. 4B) at one meter from its origin. The crossing area includes tracks 17.4 and 7.4 (Fig. 4B), respectively, the former track overlapping the latter, and thus indicates the sequence of trampling. Trackway F17 turns east and aligns parallel to trackway F7 (Fig. 4B). They follow a north-northwest direction for about 10 meters. Trackway F17 finally crosses again with trackway F7 in a region that includes tracks 17.15 - 17.17 and 7.15 - 7.17 (crossing area 2, Fig. 4B). No overlapping of tracks is found in this area, although tracks are located very close to each other. Trackway F4 is parallel to but with an opposite direction to trackways F17 and F7, from which it is separated by 2 meters (Fig. 4B). Trackway F5 has a subperpendicular direction to trackways F17, F7 and F4, and intercepts trackway F4 at track 4.4 (crossing area 3, Fig. 4B) without evidence of overlapping.

### 1. Morphological Variation

Field observations and photogrammetric models (Appendix S2) revealed various intra-trackway morphotypes (Fig. 6). The morphological variation is exemplified by four different track shapes from the starting (17.2–7.3, Fig. 6A–B) and ending portions (17.17–7.13, Fig. 6A1–B1) of trackways F17 and F7. Track 17.2 in Figure 6A is characterized by being deep and poorly detailed with thin digital impressions (particularly Digit III), bounded by substantial displacement rims. In this regard, it is not uncommon to observe the exit hole *sensu* [45] p.39 of digit III. When digit III is long, and distinguishable, digits II and IV tend to be narrow due to wall collapse (e.g., tracks 17.1, 17.3, and 17.10; see three-dimensional model capture of El Frontal tracksite in Appendix S1). Conversely, when digit III is sealed and bounded by sediment ridges, impressions of digits II and IV are thicker (e.g., Tracks 17.17 and 5.2, Fig. 6A1, see Appendix S2). Track 17.2 (Fig. 6A) shows a deep central area and a deeply impressed and elongated metatarsal mark similar to that reported by Kuban [46]. Sometimes, tracks preserving impressions of digits II, III, and IV exhibit a posteromedially oriented hallux mark in the rear margin.

Track 17.17 (Fig. 6A1) belongs to the same trackway as track 17.2, yet 17.17 lacks the hallux and metatarsal impressions which dominate the morphology of 17.2. On the other hand, track 7.3 (Fig. 6B) is a shallow track with a typical tridactyl appearance, digits II and IV usually well impressed, and digit III marked only in its distal part (eg., tracks 4.8 and 7.3, Fig. 6B, see Appendix S2 and Table 2). Track 7.3 shows very little extraneous substrate deformation. In the same trackway, track 7.13 (Fig. 6B1) is found, which differs substantially from 7.3, being considerably deeper and with displacement rims between digits II–III and III–IV. There is also a deep impression where the digits converge at the metatarsal pad – an impression almost entirely absent from track 7.3. The tracks differ according to characters such as the presence/absence of hallux or metatarsus impressions, interdigital rims, mud collapse structures and pad impressions.

### 2. Quantification of Morphological Variation

The shape variation described above is reflected in changes in track parameters such as the measurable pace length (PL), track length (TL) and depth (D), and maximum height of the associated displacement rims (DR) (Table 2). This morphological variation is presented quantitatively in Figures 7–9.

The D versus DR graphic (see Table 2 and Fig. 7) shows the relationship between the depth (D) of the tracks, and the maximum sediment height of the associated displacement rims (DR) (Fig. 7). It shows that these two parameters are positively correlated (Pearson's correlation matrix  $r = 0.871$  and Spearman's correlation matrix  $r = 0.820$ ). Deeper tracks show the highest displacement rims between the digits.

The figure 8 shows considerable intra-trackway variation. More importantly, in deep tracks, measurable track length (TL) appears influenced by track depth. Thus, trackways F17, F7 and F5 show a wide range of values, displaying a very pronounced variability in both depth (D) and track length (TL) parameters (see Table 2). By contrast, tracks forming trackway F4 are more closely grouped, and the values are somewhat more conservative and consistent along the trackway. In trackway F17 (17 measurements), the D parameter ranges from 48 mm to 13 mm (mean: 31.5 mm,

**Table 2.** Table with measurements taken for all tracks belonging to trackways F17, F7, F5 and F4 of the El Frontal tracksite.

TRACKS	D	TL	PL	DR
7.1	0,18	20	88	0.700
7.2	0,18	21	98	0.320
7.3	0,17	23	100	0.320
7.4	0,9	22	95	0.917
7.5	0,8	22	95	0.600
7.6	0,8	24	93	0.710
7.7	0,7	25	84	0.8
7.8	0,18	25	82	0.75
7.9	1,3	26	94	0.64
7.10	1,4	22	93	0.6
7.11	1,5	25	103	0.7
7.12	2,1	24	102	0.782
7.13	2,25	21	99	1.709
7.14	2,1	22	97	0.9
7.15	2	25	100	0.8
7.16	1,9	19	99	1.189
7.17	1,2	25	99	0.553
average	1,16	23	95,35	
desvst	0,74	2,06	5,97	
max	1,9	25,06	101,32	
min	0,41	20,94	89,38	
17.1	3,18	30	91	1.582
17.2	4,8	29	95	4.179
17.3	4,32	27	96	2.951
17.4	3,1	31	85	1.132
17.5	2,9	29	90	2.267
17.6	2,8	27	100	1.771
17.7	2,6	26	94	1.432
17.8	2,6	32	98	1.679
17.9	2,9	27	115	1.965
17.10	2,7	32	108	2.256
17.11	3,1	34	88	2.332
17.12	3,5	31	98	1.879
17.13	3,8	30	103	2.277
17.14	4,7	29	100	2.653
17.15	3,1	32	100	2.331
17.16	2	27	94	2.014
17.17	1,3	30	100	1.511
average	3,15	29,59	97,35	
desv	0,88	2,27	7,29	
max	4,03	31,85	104,64	
min	2,27	27,32	90,06	
5.1	2,2	17	74	1.1
5.2	2,2	19	68	1.1
5.3	1,7	16	69	0.365
5.4	1,4	18	66	0.2
5.5	0,8	22	65	0.398
average	1,66	18,4	68,4	

**Table 2.** Cont.

TRACKS	D	TL	PL	DR
desv	0,59	2,3	3,51	
max	2,25	20,7	71,91	
min	1,07	16,1	64,89	
4.1	0,5	23	106	0.35
4.2	0,6	26	103	0.593
4.3	0,7	26	101	0.574
4.4	0,7	27	103	0.798
4.5	0,9	30	105	0.8
4.6	1	23	102	1.01
4.7	1,2	26	100	1.02
4.8	0,66	22	100	0.987
4.9	0,69	24	103	0.617
4.10	0,8	27	100	0.781
average	0,78	25,4	102,3	
desv	0,21	2,41	2	
max	0,98	27,81	104,3	
min	0,57	22,99	100,3	

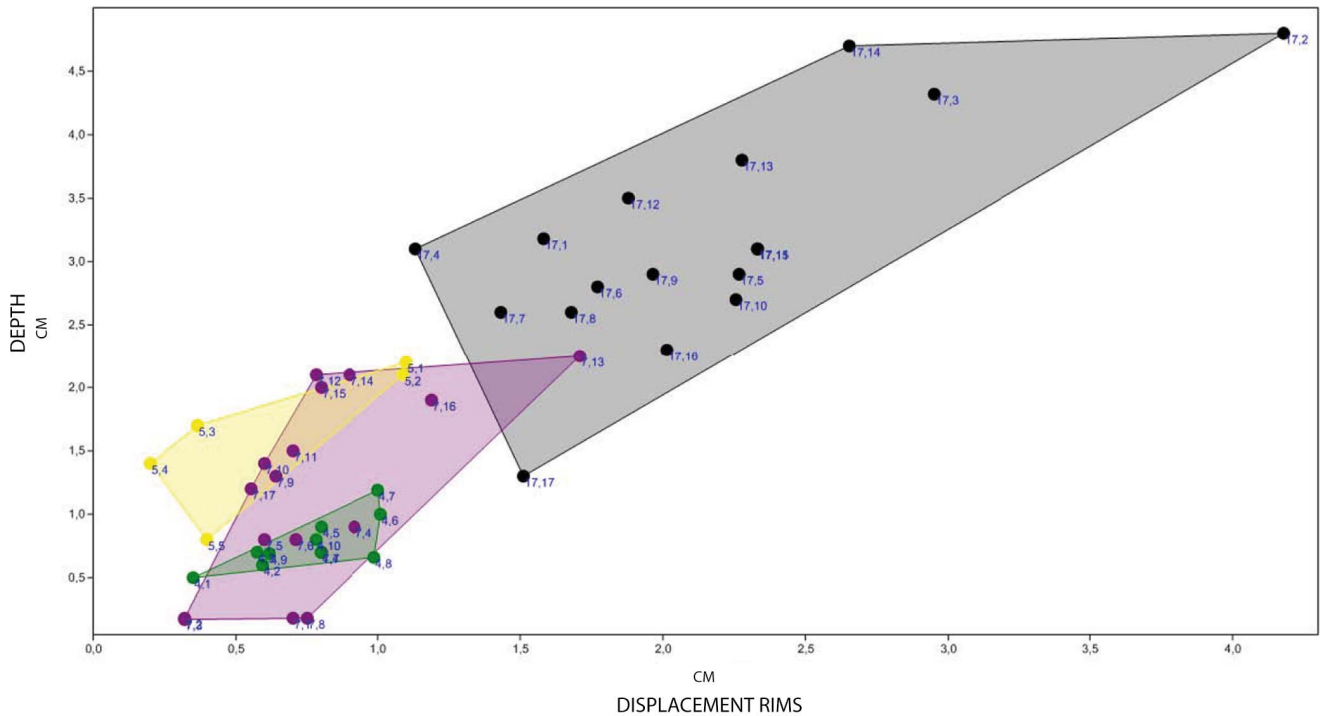
DR (depth rims); D (depth of the track); TL (track length); PL (pace length). For each measurement average (media), standard deviation (desv) and values with the maximum and minimum standard deviation (max and min) are calculated. All measurements in Table 2 are in CM.  
doi:10.1371/journal.pone.0093708.t002

SD±0.88, Table 2). It is noteworthy that the highest values are in the tracks that show evidence of hallux and metatarsal impression marks. Depth in trackway F7 (17 measurements) displays a range of values from 1.7 mm to 22.5 mm (mean = 11.6, SD±0.74, table 2). In trackway F5 (5 measurements), depth ranges from 22 mm to 8 mm (mean = 16.6 mm, SD±0.59, table 2). Trackway F4 (10 measurements) shows a range from 5 mm to 12 mm in depth (mean = 7.8 mm, SD±0.2, table 2).

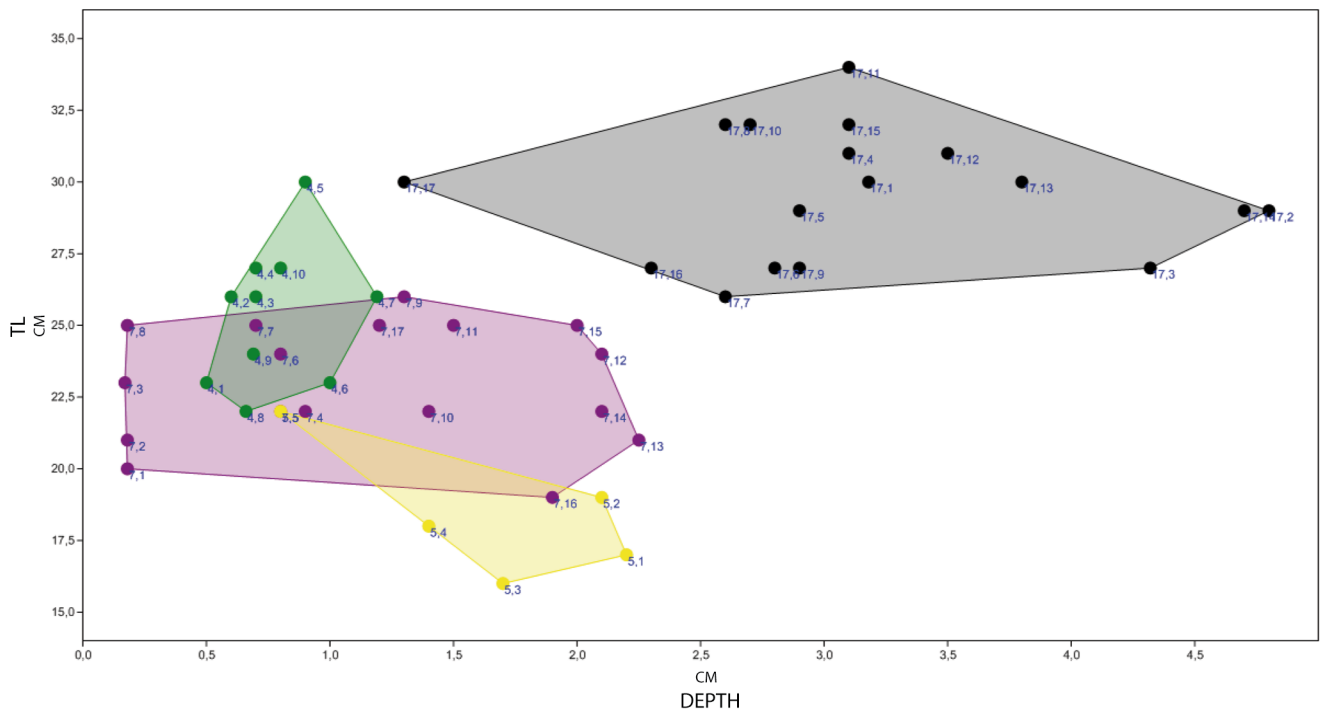
To quantify this substantial intra-trackway depth and length variations, four graphs for trackways F17, F7, F5 and F4 were built using TL, PL (left Y axis) and depth measurements (right y axis) (Fig. 9).

Figure 9 shows that: a) the pace length (PL) displays some variations, especially in trackways F17 and F7, b) the track length (TL) changes to a lesser extent than the depth parameter, and c) that the depth (D) is the most variable measurement. In particular, in F17, there is considerable variation in depth among the first few tracks, yet track length and pace length remain relatively consistent. Between tracks 17.4 (crossing area 1) and 17.9, TL and D both display a decrease of a 19% and 16%, respectively, while PL increases by 26%. From track 17.9 to 17.14, the PL decreases by 13% and D increases by 42%. From track 17.14 to track 17.17 (crossing area 2), D strongly decreases a 58%, while PL and TL do not show remarkable variations.

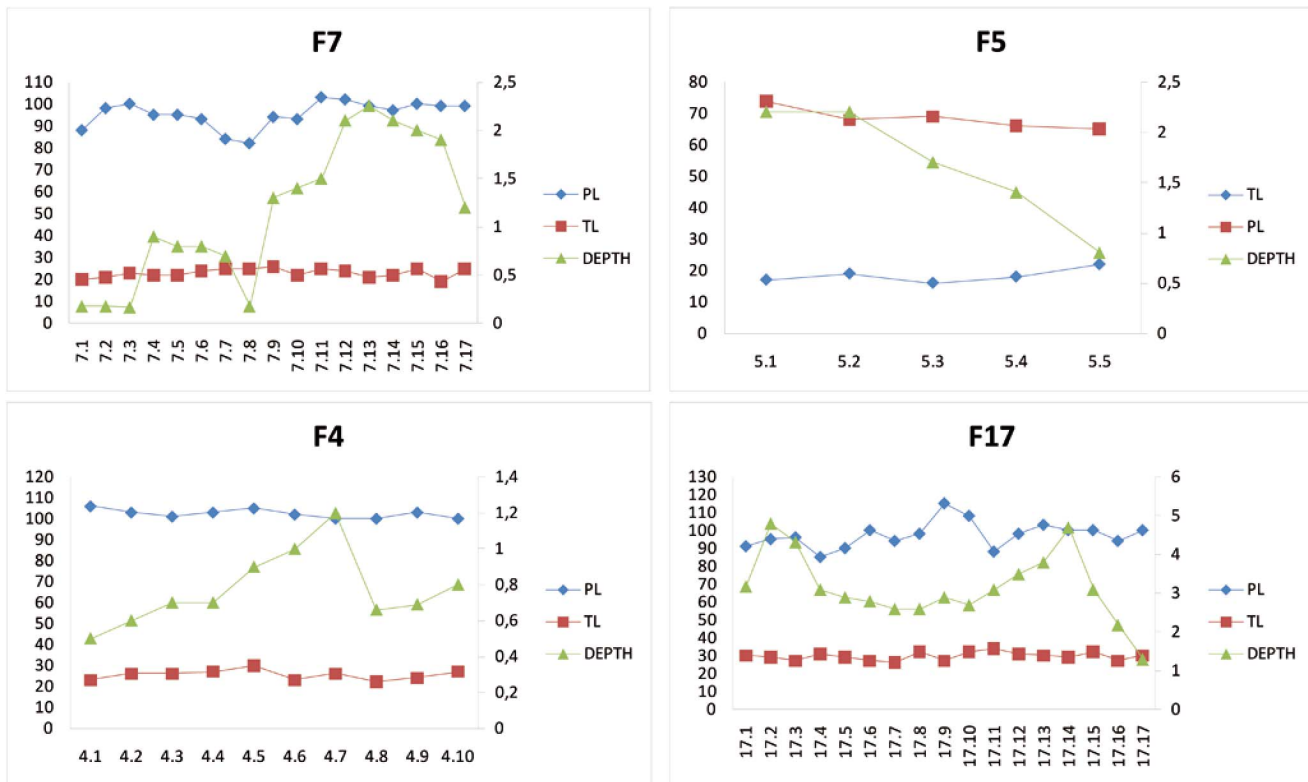
Between tracks 7.1 and 7.3, trackway F7 displays an increasing PL with constant D and TL. From track 7.4 to 7.8, PL slightly decreases, while D increases until track 7.7 to then decrease in track 7.8. From this point of the graphic, a remarkable increase in D (92%) is recorded from track 7.8 to track 7.13, but PL increases only by 17%. Finally, from track 7.14 to track 7.17 (crossing area 2), although PL and TL remain averagely constant, D strongly decreases by 43%. Trackway F5 displays a decreasing D (63%) from track 5.1 to 5.5, with an average constant PL and TL.



**Figure 7. The linear correlation graphic of depth (D) vs displacement rims (DR) shows a positive correlation among these two values.** Pearson’s correlation matrix results in  $r=0.871$  and Spearman’s correlation matrix in an  $r=0.820$ . F17 (black colour), F7 (purple colour), F5 (yellow colour) and F4 (green colour). Units are in centimeters.  
doi:10.1371/journal.pone.0093708.g007



**Figure 8. Dispersion graph of depth (D) vs track length (TL) shows a wide range distribution among the tracks of trackways F17, F7 and F5 (respectively black, purple and yellow colours) of the El Frontal tracksite.** The most concentrated cluster in that of trackway F4 (green colour), in which values are quite consistent and only weakly vary along the trackway. Units are in centimeters.  
doi:10.1371/journal.pone.0093708.g008



**Figure 9. Quantification of the intra-trackway depth and length variations, four graphs for trackways F17, F7, F5 and F4 are built using TL, PL (left Y axis) and depth measurements (right y axis). Units are in centimeters.**  
doi:10.1371/journal.pone.0093708.g009

Trackway F4 shows an increasing D from track 4.1 to track 4.7 (58%), a variable TL (23% of variation) and a weakly decreasing PL (6%). From track 4.7 to track 4.10, D decreases quite abruptly (33%), while PL and TL do not display significant variations.

## Discussion

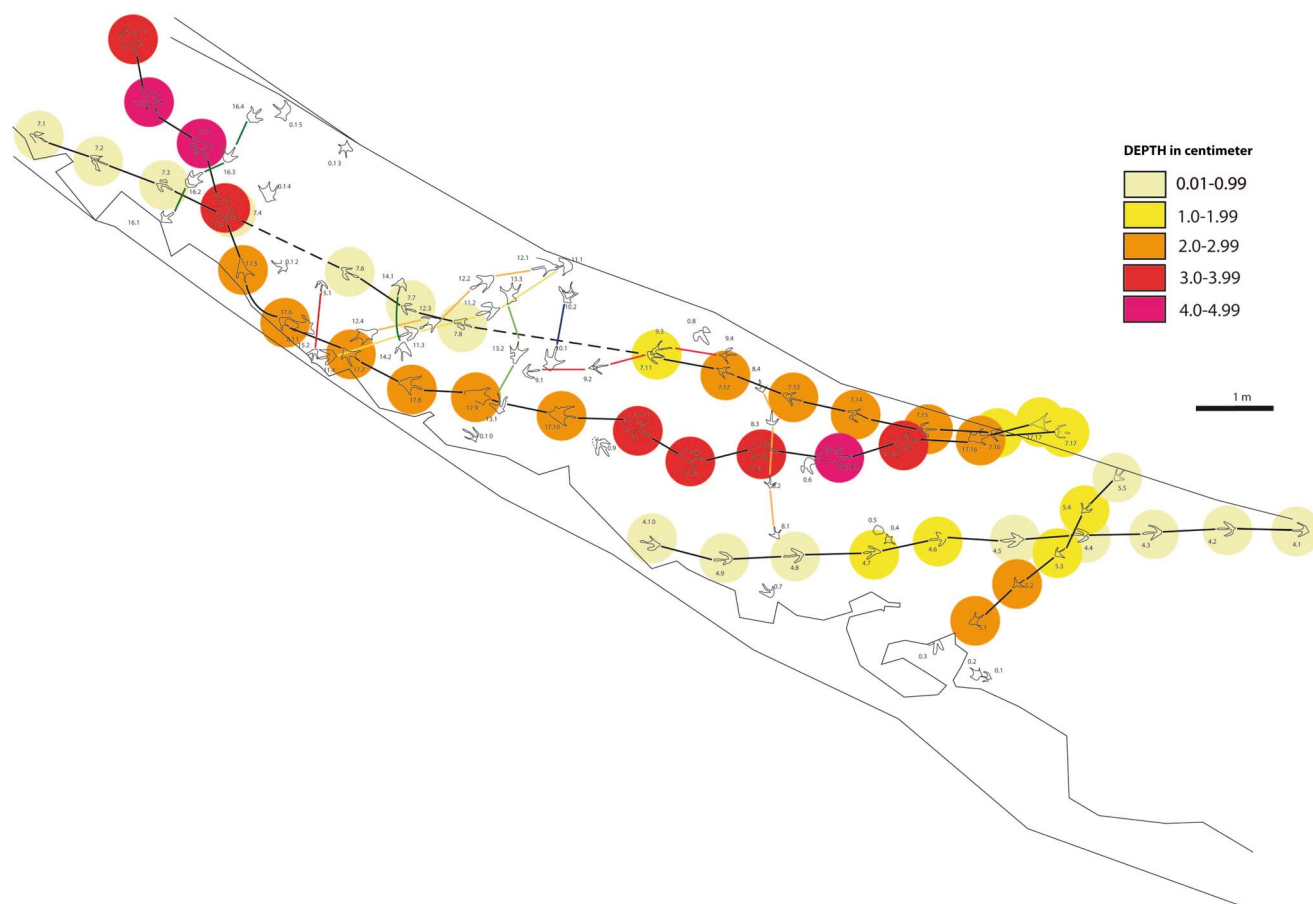
The El Frontal site is an exceptional example of high within-trackway morphological variation. The final morphology of a track is determined by the shape of the track maker's foot, the dynamics of that foot, and the substrate conditions [4,47,48]. Within-trackway morphological variation cannot come from variations in foot anatomy, and therefore must originate from horizontal sediment heterogeneity, differences in limb dynamics, or a combination of the two.

The morphological variation of all tracks (Fig. 6 and Appendix S2) is highly influenced by the depth to which the animal sank. By observing the position of each track in the El Frontal tracksite (Fig. 4) and comparing the graphics with each other, it is noticed that similar depth trends are recorded for F4, F7 and F17, which are located parallel in the tracksite. Among trackway segments 4.6–4.10, 17.11–17.14 and 7.11–7.14, a progressive depth increase is recorded, while among trackway segments 4.5–4.1, 17.15–17.17 and 7.15–7.17 depth decreases. Trackways F17 and F7 differ in PL, TL and D values quite strongly (see Table 2), although they behave similarly along three different intra-trackway segments: between 17.2–17.8 and 7.4–7.8, depth decreases in both trackways, between 17.8–17.14 and 7.8–7.13 depth increases and finally, between 17.14–17.17 and 7.14–7.17 depth strongly decreases in both trackways (72% and 43% respectively). This last zone corresponds to the crossing area 2 (Fig. 4), in which

trackways are closely located and, although displaying different absolute values of the parameters (Fig. 9), they present a similar trend in responding to the substrate (depth decrease).

Trackway F5, which crosses the site perpendicular to the other trackways, does not display any intra-trackway variation, or similar trends to those of F4, F7 and F17. Nevertheless, tracks 5.4 and 5.5 decrease depth values when approaching to the crossing area 2, where the general tendency is for tracks to be deeper (eg. F17 and F7).

It has been accepted for a long time that the depth to which a foot sinks is a determinant parameter in understanding the soil mechanics that control track formation [1,2,6,10,19–21,49]. The deep tracks at the El Frontal site represent part of a continuum that must have been produced on a laterally heterogeneous substrate (Figures 6 and 10). Hence, tracks change their morphology in accordance to their relative position along a substrate consistency gradient that persisted across the site (Figures 9 and 10). Scrivner and Bottjner [50] and Allen [51] suggested that there is a positive correlation between the foot penetration and the degree of deformation in a sediment. At the El Frontal tracksite, the D versus DR and D versus TL graphics (Figures 7 and 8) show a high difference of values for the 49 tracks considered in the sample as a whole and within single trackways. The dispersion graphics underpin the importance of substrate response with respect to track length and depth variations during the indentation of the foot. If the substrate conditions of the El Frontal tracksite were uniform throughout the trampled surface, foot loads made by comparably sized animals moving in a dynamically similar fashion (see PL in Fig. 9) would have produced similar tracks (same track length and depth) along single and associated trackways. On the contrary, we observe that track depth



**Figure 10. Morphological continuum of the tracks of the El Frontal tracksite. Color scale bar is based on depth intervals.**  
doi:10.1371/journal.pone.0093708.g010

and morphology are extremely variable both within a single trackway and the whole track sample (Fig. 9).

Tracks can be used to provide additional information on the conditions of the substrate at the time of track formation. Various works [1,8,10,20,52–54] underscored the fact that substrate properties such as consistency, sediment composition (e.g. proportion of clay minerals), grain size, texture, water content and rate of consolidation control and bias the resulting track morphology. The sedimentological analyses performed on the El Frontal site support with the idea that the original substrate was non-homogenous due to lateral changes in adjoining microfacies. Thin sections of layers 4 and 5 (Fig. 3D–F) reveal sedimentary structures (mud drapes and symmetrical ripples) that are usually found when the surrounding environment is characterized by interruptions in the continuity of water flows, such as the current produced in environments with tidal influence [37,38]. This implies that the energetic episodes are frequent, fluctuant and intermittent (Fig. 3D–F). A substrate with a higher water content offers more favourable conditions to produce deep tracks (Fig. 3A–C, 6A,B1). A drier substrate of firmer quartz dominant sandstone is more likely to have produced shallow tracks (Fig. 3A–B,6A1,B). The El Frontal tracks exhibit different depths and morphologies resulting from varying rheological conditions due to a lateral facies of changeable consistency, perpendicular to F17, F7 and F4, but parallel to F5, which is affected to a lesser extent.

Finally, in the current state of knowledge it seems difficult to assign any of the studied tracks to a particular group of tridactyl trackmakers, especially regarding the difficulties distinguishing

between theropods and ornithopods in the Iberian Range during Berriasian times [33,55]. The presence of hallux marks and large steps might indicate a probable theropod origin [56], though they are not exclusive characters of this group. Several theropod ichnotaxa have been described in the Huérteles Formation: *Megalosauropus isp.* [57], *Kalohipus bretunensis* [58], “*Fillichnites gracilis*” [59] and *Archaeornithipus mejjidei* [60]. Moreover, some grallatorid [61] and *Buckerburgichnus*-like tracks have been reported [62]. Inferences on possible ichnotaxa in the El Frontal tracksite are tangled by the morphological variability observed in the site. The substrate bias in the morphology prevents us from assigning any of the tracks to a particular ichnotaxon, and the strong substrate bias affects track morphology in such a way that it rarely correlates with real foot anatomy of the trackmaker. Interestingly, this study opens a new window into the interpretation of the aforementioned ichnotaxa in the Huérteles Formation and questions whether some of them might represent taphotaxa *sensu* [63].

## Conclusions

The El Frontal tracksite displays a variety of tridactyl track morphologies and provides a valuable example of how track geometry might be dominantly affected by substrate conditions during formation, implying that rheology is the major factor in track formation. The photogrammetry models and depth analyses spotlighted that the deep and shallow tracks are part of a continuum of track morphologies and depths. Sedimentological analyses revealed that the site was a non-homogenous substrate

that experienced lateral changes due to fluctuating and intermittent flow episodes in a fluvial-deltaic environment. The tracksite differentiation of substrate consistencies and the vast range of intra-trackway morphologies suggest that tracks were produced by similar trackmakers crossing the lateral gradient of heterogeneous substrate consistencies.

The presented analyses underline the influence of substrate on the final track morphology and length. The within-trackway variation highlights that ichnotaxonomic assignments of sediment-biased tracks should be avoided.

## Supporting Information

**Appendix S1** Caption of three-dimensional El Frontal tracksite. Scale bar 1 meter. (TIF)

**Appendix S2** Photogrammetry and depth analysis respectively undertaken with free software VisualSFM and Schlumberger package Petrel of the El Frontal tracksite. Tracks are disposed

## References

- Manning PL (2004) A new approach to the analysis and interpretation of tracks: examples from the Dinosauria. In: The application of ichnology to palaeoenvironmental and stratigraphic analysis. McIlroy, D. (Ed.), Geological Society, London, Special Publications 228: 93–123.
- Milán J, Bromley RG (2006) True tracks, undertracks and eroded tracks, experimental work with tetrapod tracks in laboratory and field. *Palaeogeography, Palaeoclimatology, Palaeoecology* 231: 253–264.
- Díaz-Martínez I, Pérez-Lorente F, Canudo JI, Pereda-Suberbiola X (2009) causas de la variabilidad en icnitas de dinosaurios y su aplicación en icnotaxonomía. In: Actas de las IV Jornadas internacionales sobre paleontología de dinosaurios y su entorno (P. Huerta And F. Torcida, Eds), Salas De Los Infantes, Burgos: 207–220.
- Falkingham PL (2014) Interpreting ecology and behaviour from the vertebrate fossil track record. *Journal of Zoology*: doi:10.1111/jzo.12110.
- Padian K (1999) Dinosaur tracks in the computer age. *Nature* 399 (6732): 103–104.
- Gatesy SM (2003) Direct and indirect track features: what sediment did a dinosaur touch? *Ichnos* 10: 91–98.
- Milán J, Bromley RG (2008) The impact of sediment consistency on trackand undertrack morphology: experiments with emu tracks in layered cement. *Ichnos* 15: 19–27.
- Marty D, Strasser A, Meyer CA (2009) Formation and taphonomy of human footprints in microbial mats of present-day tidal-flat environments: implications for the study of fossil footprints. *Ichnos* 16(1): 127–142.
- Ellis RG, Gatesy SM (2013) A biplanar X-ray method for three-dimensional analysis of track formation. *Palaeontologia Electronica* 16 (1)1t, 16p; [Palaeo-Electronica.Org/Content/2013/371-X-Ray-Track-Analysis](http://Palaeo-Electronica.Org/Content/2013/371-X-Ray-Track-Analysis).
- Morse SA, Bennet MR, Liutkus-Piercem C, Thackeray F, McClymont J, et al. (2013) Holocene footprints in Namibia: the influence of substrate on footprint variability. *American Journal of Physical Anthropology* 151: 265–279.
- Baird D (1957) Triassic reptile footprint faunules from Milford, New Jersey. *Bulletin of the Museum of comparative zoology, Harvard University* 117: 449–520.
- Manning PL (2008) *T.rex* speed trap. In: *Tyrannosaurus rex*, The Tyrant King. Farlow, O (Ed.), Indiana University Press, 204–231.
- Barnes HA (2000) A handbook of elementary rheology. The University of Wales Institute of non-newtonian fluid mechanics, department of Mathematics, University Of Wales Aberystwyth, Penglais, Aberystwyth, Dyfed, Wales, Sy23 3bz. 201p.
- Craig RF (2004) Craig's soil mechanics. Spon Press, Abingdon, 447p.
- Schanz T, Lins Y, Viehhaus H, Barciaga T, Låbe S, et al. (2013) Quantitative interpretation of tracks for determination of body mass. *PLoS ONE* 8(10): e77606. doi:10.1371/journal.pone.0077606.
- Currie PJ, Sarjeant WAS (1979) Lower Cretaceous dinosaur footprints from the Peace River Canyon, British Columbia, Canada. *Palaeogeography, Palaeoclimatology, Palaeoecology* 28: 103–115.
- Casamiquela RM, Demathieu GR, Haubold H, Leonardi G, Sarjeant WAS (1987) Glossary and manual of tetrapod footprint palaeoichnology. Leonardi, G. (ed.): 22–25.
- Gatesy SM, Shubin NH, Jenkins FA Jr (2005) Anaglyph stereo imaging of dinosaur track morphology and microtopography. *Palaeontologia Electronica* 8 (1)10a: 10p, 693kb; [Http://Palaeo-Electronica.Org/Palco/2005\\_1/Gatesy10/Issue1\\_05.Htm](http://Palaeo-Electronica.Org/Palco/2005_1/Gatesy10/Issue1_05.Htm).

vertically to underpin the intra-trackway morphological variation. Color scale green and yellow indicates the track layer, purple is the deepest point recorded for depth and red is the highest point recorded for displacement rims.

(TIF)

## Acknowledgments

We thank Arnau Bolet and Frank Rarity for field assistance. Oriol Oms, Soledad Esteban de Trivigno, Ignacio Díaz-Martínez and Carlos Pascual made useful comments on a previous version of the manuscript.

## Author Contributions

Conceived and designed the experiments: NLR BV DC PLF JLB JIC PLM AG. Performed the experiments: NLR BV DC PLF JLB JIC PLM AG. Analyzed the data: NLR BV DC PLF JLB JIC PLM AG. Contributed reagents/materials/analysis tools: NLR BV DC PLF JLB JIC PLM AG. Wrote the paper: NLR BV DC PLF JLB JIC PLM AG.

- Gatesy SM, Middleton KM, Jenkins FA, Shubin NH (1999) Three-dimensional preservation of foot movements in triassic theropod dinosaurs. *Nature* 399: 141–144.
- Jackson SJ, Whyte MA, Romano M (2010) Range of experimental dinosaur (hypsilophodon foxii) footprints due to variation in sand consistency: how wet was the track? *Ichnos* 17: 197–214.
- Falkingham PL, Margetts L, Manning PL (2010) Fossil vertebrate tracks as paleopneumotometers: confounding effects of foot morphology. *Palaios* 25: 356–360.
- Bates K, Manning PL, Vila B, Hodgetts D (2008a) Three-dimensional modelling and analysis of dinosaur trackways. *Palaeontology* 51: 999–1010.
- Bates K, Rarity F, Manning PL, Hodgetts D, Vila B, et al. (2008b). High-resolution lidar and photogrammetric survey of the Fumanya dinosaur tracksites (Catalonia): implications for the conservation and interpretation of geological heritage sites. *Journal Geological Society* 165: 115–127.
- Breithaupt BH, Southwell EH, Adams T, Matthews NA (2001) Innovative documentation methodologies in the study of the most extensive dinosaur tracksite in Wyoming. 6th Fossil Research Conference Proceedings Volume: 113–122.
- Falkingham PL, Margetts L, Smith I, Manning PL (2009) Reinterpretation of palmate and semi-palmate (webbed) fossil tracks; insights from finite element modelling. *Palaeogeography, Palaeoclimatology, Palaeoecology* 271: 69–76.
- Bellian JA, Kerans C, Jennette DC (2005) Digital outcrop models: applications of terrestrial scanning lidar technology in stratigraphic modeling. *Journal Of Sedimentary Research* 75: 166–176.
- Falkingham PL (2012) Acquisition of high resolution 3D models using free, open-source, photogrammetric software. *Palaeontologia Electronica* 15(1); 1t:15p; [Palaeo-Electronica.Org/Content/93-Issue-1-2012-Technical-Articles/92-3d-Photogrammetry](http://Palaeo-Electronica.Org/Content/93-Issue-1-2012-Technical-Articles/92-3d-Photogrammetry).
- Aguirrezabala LM, Viera LI (1980) Icnitas de dinosaurios en Bretún (Soria). *Munibe* 32: 257–279.
- Sanz JL, Moratalla JJ, Rubio JL, Fuentes C, Mejjide M (1997) Huellas de dinosaurios de Castilla y Leon. Ed Junta de Castilla y Leon: 87 p.
- Barco JL, Ruiz-Omeñaca JI (2005) la ruta de las icnitas de dinosaurio de Soria. una excursión a los comienzos del Cretácico. *Cuadernos De Paleontología Aragonesa* 4: 60–90.
- Razzolini NL, Vila B, Barco JL, Galobart À, Falkingham PL, et al. (2012) First approach to the El Frontal tracksite (Berriasián, Soria, Spain): perspectives on morphological variability in theropod tracks. 207–210 P. In: Royo-Torres, R., Gascó, F. And Alcalá, L., Coord. (2012). 10th Annual Meeting Of The European Association Of Vertebrate Palaeontologists. *Fundamental!* 20: 1–290.
- Moratalla JJ, Hernán J (2010) Probable palaeogeographic influences of the Lower Cretaceous Iberian rifting phase in the eastern Cameros basin (Spain) on dinosaur trackway orientations. *Palaeogeography, Palaeoclimatology, Palaeoecology* 295: 116–130.
- Castanera D, Pascual C, Razzolini NL, Vila B, Barco JL, et al.(2013a) Discriminating between medium-sized tridactyl trackmakers: tracking ornithopod tracks in the base of the Cretaceous (Berriasián, Spain). *PLoS ONE* 8(11): e81830. doi:10.1371/journal.pone.0081830.
- Gómez-Fernández JC, Meléndez N (1994) Climatic control on Lower Cretaceous sedimentation in a playa-lake system of a tectonically active basin (Huérteles Alloformation, Eastern Cameros Basin, North-Central Spain). *Journal of Paleolimnology* 11: 91–107.
- Salas R, Guimerà J, Mas R, Martín-Closas C, Meléndez A, et al. (2001) Evolution of the Mesozoic central Iberian Rift System and its Cainozoic

- inversion (Iberian Chain). In: Ziegler PA, Cavazza W, Robertson AHF, Crasquin-Soleau S, editors. Peri-Tethys Memoir 6: Peri-Tethyan Rift/Wrench Basins and Passive Margins: Museum National d'Histoire Naturelle, Paris. 145–186.
36. Mas R, García A, Salas R, Meléndez A, Alonso A, et al. (2004) Segunda fase de rifting: Jurasico Superior- Cretácico Inferior, In Vera, J.A. (Ed.) Geología De España. Sge-Igme, España, 503–510.
  37. Quijada EI, Suárez-González P, Benito MI, Mas JR, Alonso A (2010) Un ejemplo de llanura fluvio-deltaica influenciada por las mareas: el yacimiento de icnitas de Serrantes (Grupo Oncala, Berriasiense, Cuenca de Cameros, N. de España). Geogaceta 49: 15–18.
  38. Quijada EI, Suarez-Gonzalez P, Benito MI, Mas R (2013) New insights on stratigraphy and sedimentology of the Oncala Group (eastern Cameros basin): implications for the paleogeographic reconstruction of NE Iberia at Berriasian times. Journal of Iberian Geology 39 (2): 313–334.
  39. Martín-Closas C, Alonso-Millán A (1998) Estratigrafía y bioestratigrafía (Charophyta) del Cretácico Inferior en el sector occidental de la Cuenca de Cameros (Cordillera Ibérica). Revista de la Sociedad Geológica de España 11 (3–4): 253–269.
  40. Schudack U, Schudack M (2009) Ostracod biostratigraphy in the Lower Cretaceous of the Iberian Chain (eastern Spain). Journal of Iberian Geology 35: 141–168.
  41. Allen JRL (1982) Sedimentary Structures: Their Character And Physical Basis. Amsterdam: Elsevier.
  42. Bates KT, Falkingham PL, Rarity F, Hodgetts D, Purslow A, et al. (2010) Application Of high-resolution laser scanning And photogrammetric techniques to data acquisition, analysis and interpretation in paleontology. International Archives Of Photogrammetry, Remote Sensing And Spatial Information Series, Commission. V Symposium, Newcastle Upon Tyne, Uk, Xxxviii 5: 68–73.
  43. Wu C (2013) Towards linear-time incremental structure from motion. 3DV 2013.
  44. Barco JL, Castilla D, Ibáñez D, Morós A, Rasal S, et al. (2005) Optimización de recursos en la elaboración de planimetrías de yacimientos de icnitas mediante la utilización de aparatos topográficos y aplicaciones de entorno cad. XXI Jornadas de la Sociedad Española de Paleontología: 169–171.
  45. Milan J (2003) Experimental Ichnology: Experiments with track and undertrack formation using emu tracks in sediments of different consistencies, with comparison to fossil dinosaur tracks. Msc Thesis, Geological Institute, University Of Copenhagen. 91p.
  46. Kuban GJ (1989) Elongate dinosaur tracks. In: Dinosaur tracks and traces. Ed. D. D. Gillette Y M. G. Lockley. Cambridge University Press, Ed. D. D. Gillette Y M. G. Lockley. Cambridge University Press: 57–72.
  47. Padian K, Olsen PE (1984) The fossil trackway *Pteraichnus*: not pterosaurian, but crocodylian. Journal of Paleontology 58: 178–184.
  48. Minter NJ, Braddy SJ, Davis RB (2007) Between a rock and a hard place: arthropod trackways and ichnotaxonomy. Lethaia 40: 365–375.
  49. Falkingham PL, Bates KT, Margetts L, Manning PL (2011) The ‘Goldilocks’ effect: preservation bias in vertebrate track assemblages. Journal Of The Royal Society Interface 8: 1142–1154.
  50. Scrivner PJ, Botjter DJ (1986) Neogene avian and mammal tracks from Death Valley National Monument, California: their context, conservation and preservation. Paleogeography, Paleoclimatology, Paleoecology 57: 285–331.
  51. Allen JRL (1989) Fossil vertebrate tracks and indenter mechanics. Journal Of The Geological Society 146: 600–602.
  52. Manning PL (1999) Dinosaur track formation, preservation and interpretation: fossil and laboratory simulated track studies. University of Sheffield, PhD Thesis, 440 pages.
  53. Allen JRL (1997) Subfossil mammalian tracks (flandrian) in the Severn Estuary, SW Britain: mechanics of formation, preservation and distribution. Philosophical Transactions Of The Royal Society Of London, Series B, Biological Sciences 352: 481–518.
  54. Scott JJ, Robin W, Renaut R, Bernhart O (2010) Taphonomic controls on animal tracks at saline, alkaline Lake Bogoria, Kenya Rift Valley: impact of salt efflorescence and clay mineralogy. Journal Of Sedimentary Research 80: 639–665.
  55. Castanera D, Vila B, Razzolini NL, Falkingham PL, Canudo JI, et al. (2013b) Manus track preservation bias as a key factor for assessing trackmaker identity and quadrupedalism in basal ornithopods. PLoS ONE 8(1): e54177. doi:10.1371/journal.pone.0054177.
  56. Lockley MG (2009) New perspectives on morphological variation in tridactyl footprints: clues to widespread convergence in developmental dynamics. Geological Quarterly 53 (4): 415–432.
  57. Barco JL, Canudo JI, Ruiz-Omeñaca JI, Pérez-Lorente F, Rubio De Lucas JL (2004) Ichnological evidence of the presence of gigantic theropods in the Berriasian (Lower Cretaceous) of Spain. In: Abstract Book – First international congress on ichnology, Trelew, Argentina L.A. Buatois & G. Mángano: 18–19.
  58. Fuentes Vidarte C, Meijide Calvo M (1998) Icnitas de dinosaurios terópodos en el Weald de Soria (España). Nuevo icnogénero *Kalohipus*. Estudios Geológicos 54: 147–152.
  59. Moratalla García JJ (1993) Restos indirectos de dinosaurios del registro español: paleoicnología de la Cuenca de Cameros (Jurásico Superior-Cretácico Inferior) y paleoología del Cretácico Superior. Universidad Complutense De Madrid PhD Thesis, 727p.
  60. Fuentes Vidarte C (1996) Icnitas de dinosaurios en Soria (España). Zúbia 14: 57–64.
  61. Pascual-Arribas C, Hernández-Medrano N (2011) Posibles huellas de crías de terópodo en el yacimiento de Valdehijuclos (Soria, España). Studia Geologica Salmanticensis 47: 77–110.
  62. Hernández Medrano N, Pascual Arribas C (2008) Los yacimientos de icnitas de dinosaurio y de otros reptiles en la provincia de Soria. Arevacon 28: 18–31.
  63. Lucas SG (2001) Taphotaxon. Lethaia in Lethaia Seminar 34: 30.