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Research

Assessing the potential of compaction techniques in tropical peatlands for effective carbon reduction and climate change mitigation

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Abstract

There is a pressing need to tackle carbon emissions from oil palm plantations on tropical peatland, which has garnered significant discussion and concern in recent years. In response, compaction techniques were introduced in Malaysia with the aim of mitigating CO₂ emissions by improving moisture levels and reducing soil aeration. This research investigates the impact of mechanical compaction on two distinct ecosystems: a peat swamp forest (PSF) and an oil palm plantation (OPP), characterized by their unique physicochemical properties. Using a specially designed compaction apparatus, significant changes in carbon emissions were observed in PSF but not in OPP, with means 1263 and 404 mg CO_{2-eq} m⁻² h⁻¹, respectively. This disparity can be due to substrate availability between the two ecosystems. Subsequently, in the PSF, a promising pattern of a percentage ratio of approximately 1:3.5 was observed, indicating a substantial reduction in CO₂ emissions (from 1295 to 468 mg m⁻² h⁻¹; 64%) alongside a corresponding increase in CH₄ emissions (from -50 to 60 μg m⁻² h⁻¹; 221%). This finding suggests that compaction alters the aerobic peat horizon, bringing the peat surface closer to the groundwater level. The study underscores the importance of considering confounding factors such as decomposition degree and groundwater fluctuation when assessing the effects of compaction on tropical peat. By shedding light on these complexities, the findings contribute to a better understanding of the efficacy of compaction techniques in reducing emissions of these special case atmospheric pollutants.

Article highlights

- The first data on *in-situ* compaction on tropical peatland carbon emissions.
- Peat physicochemical properties were not affected by compaction when in contact with groundwater.
- Appropriate clarity of mechanical compaction on tropical peatland.

Keywords Carbon emissions · Tropical peatland · Compaction · Groundwater · Oil palm plantations · Peat swamp forest

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1 Introduction

Tropical peatlands are highly delicate and sensitive ecosystems, and the conversion of these areas into agricultural farms requires intensive inputs and management practices that can have detrimental effects on the ecosystem, leading to land degradation and environmental issues [1–3]. Intensive and appropriate management practices are crucial for ensuring sustainable crop productivity in tropical peatlands. These management practices encompass various strategies, including water table management [4], soil compaction [5], and fire prevention [6]. In their natural state, tropical peatlands maintain a high water table, which is essential for the stability and carbon storage capacity of the peat layer. However, when converted into agricultural farms, drainage is often employed to lower the water table, making the land suitable for cultivation. Drainage significantly alters the physical and chemical properties of the soil and microbial activity by introducing oxygen into the soil, triggering microbial decomposition processes, and causing the breakdown of organic matter [7]. This alteration influences the release of special case atmospheric pollutants or greenhouse gases such as CO₂, CH₄, organic acids, and organic particulates [8]. Climate change exacerbates the rate of carbon loss, particularly through the establishment of oil palm plantations using drainage and slash-and-burn techniques, as demonstrated in previous studies [9]. To address the rising environmental concerns, a novel technique introduced for open land practice on tropical peatland is peat mechanical compaction, which claims to effectively minimize the risk of fire by enhancing soil moisture [10]. Several studies have explored the positive effects of peat compaction on mechanical anchorage for palm stands and nutrient retention [5, 11]. However, the impact of compressed soil on the composition and decomposition of aerobic and anaerobic peat materials, as well as its ability to reduce the risk of peat fires and lower CO₂ emissions, remains uncertain due to limited available data. Furthermore, the implications for CH₄ emissions, which have a significantly higher global warming potential than CO₂, due to induced anaerobic conditions through compaction, have yet to be fully addressed [12, 13].

In the early stages of establishing oil palm plantations on tropical peatlands, companies typically avoid using fire and instead opt for manual mechanical methods to clear the land. This approach is used whether they are clearing pristine forests, secondary forests, areas covered in *Imperata* sp. grass, or transitioning from other crops. The manual mechanical process involves a series of steps, including slashing, cutting, chopping, and

piling/staking [14], occasionally supplemented with pesticide spraying [15]. Once the land clearing phase is completed and aligns with the intended goals, the land is prepared for planting oil palm seedlings using a compaction method. Through 2 to 3 tractor passes, compaction can increase the dry bulk density of peat by as much as 0.20 g cm⁻³ [16]. However, before mechanical compaction takes place, it is customary to lower the water table level to a range of 0.5 m to 1.0 m. This is done to achieve self-subsidence and improve the suitability of peat for heavy machinery usage [5]. In such situations, two interconnected factors are expected to contribute to the increase in bulk density: compaction resulting from tractor passes and drainage, which leads to the shrinkage of organic and woody materials as water is lost [17]. Furthermore, the complexity surrounding the extent of compaction arises from a range of factors, including abiotic, biotic, and anthropogenic elements (man-made compaction) [12]. Nevertheless, differentiating compaction from shrinkage as distinct phenomena can be quite challenging. Most scholarly works [5, 11, 17, 18] have defined tropical peat compaction based on the specific characteristics of the study site. For instance, in ecosystems like logged-over or drained forests, peat compaction is often attributed to variations in the water table during seasonal changes, resulting in shrinkage (leading to bulk density increase) or oxidation (caused by microbial activity), which decreases bulk density [17]. Conversely, in developed peatland ecosystems such as agricultural systems, many authors [11, 17, 19] describe peat compaction as a combined process influenced by heavy machinery and shrinkage due to drainage, both contributing to increased bulk density. This landscape-based approach to defining compaction has led to confusion among researchers regarding the precise mechanism of compaction on tropical peatlands.

Samuel and Evers [12] underscored in their comprehensive review paper the pressing need for a profound understanding of the intricacies associated with tropical peat compaction. This phenomenon involves a complex interplay of various processes, including oxidative reactions, consolidation, shrinkage, and mechanical compaction, the latter being induced by human activities. Mechanical compaction is also commonly referred to as man-made compaction. Peat compaction can be dissected into two interlinked aspects: compression and consolidation. Compression refers to the reduction in the volume of the peat material due to the displacement of the oxic phase (which lacks water). On the other hand, consolidation entails an enhancement in the mechanical strength of the material. This strengthening is a result of particle-to-particle interactions that occur when peat soil collapses under its weight within the anoxic horizon [17]. Measuring

the compressibility of peat soil directly poses considerable challenges, primarily due to the multitude of interconnected factors, including peat depth, maturity, and bulk density [20]. Existing research on mechanical compaction and peat soil has predominantly viewed peat as a soil conditioner or an additive for mineral soils [21]. This perspective is largely influenced by the remarkably high organic matter content found in peat, typically hovering around 99% [22]. Consequently, peat soil of this nature exhibits a compressibility level within the range of 300% to 400%. It also possesses the remarkable ability to revert to its original state [23]—unless specific hydrophobic characteristics are deliberately introduced [24].

While some studies have investigated the effects of soil compaction on carbon emissions, soil water, and physicochemical properties [25–27], experimental data related to tropical peat soil-induced or artificial compaction on tropical peat physicochemical properties and carbon emissions is rare. The only available reference by Busman et al. [28] was conducted under laboratory conditions using destructive composite sub-surface samples (50 to 70 cm) and was solely based on the dry bulk density values. Hence, this study aims to assess the effect of artificial compaction of tropical peat on the peat physicochemical properties and carbon emissions. The objectives of this study are, therefore: (1) to induce compaction to vary magnitudes based on compression according to peat depth using a fabricated compaction apparatus; (2) to determine the peat physicochemical properties and quantify the peat carbon emissions according to the compaction treatment and site ecosystem (PSF: secondary peat swamp forest where pre-existing compaction via drainage will be minimal) and OPP: mature oil palm plantation (where pre-existing conditions will have been impacted by the conversion process and longer-term drainage effects); (3) to establish a relationship between carbon emissions and physicochemical properties to make recommendations on the efficacy of this approach within the two varying sites types.

2 Materials and methods

2.1 Study location

The study took place in two tropical peatlands located in the peninsular region of Malaysia (Fig. 1a), specifically in Tanjong Karang in Northern Selangor (3° 41' 49.2" N 101° 11' 06.0" E) (Fig. 1b) and Sepang in Southern Selangor (2° 44' 20.4" N 101° 39' 10.8" E) (Fig. 1c). The Tanjong Karang site was characterized by a drained peat swamp forest (PSF) that had been selectively logged in the past, while the Sepang site was represented by a second-generation oil palm plantation (OPP) that was established in 1978

and had been in operation for decades. During the study period, meteorological data from Department of Meteorology Malaysia (METMalaysia) recorded similar warm and humid weather patterns in both areas. Tanjong Karang experienced an average temperature of 28.1 °C, peaking at 33.5 °C, and dropping to 23.6 °C. Concurrently, Sepang maintained an average temperature of 28.3 °C, with a high of 33.6 °C and a low of 23.9 °C. Tanjong Karang experienced mean monthly rainfall of 127.8 mm, and Sepang saw 97.4 mm. The relative humidity was slightly higher in Tanjong Karang at 81% compared to Sepang's 79%, suggesting slightly warmer and less humid conditions in Sepang.

The history and management of each site influenced its vegetation. The Tanjong Karang site belongs to the Raja Musa Forest Reserve (RMFR), which is part of the North Selangor Peat Swamp Forest and covers 23,000 ha. RMFR joined the Queen's Commonwealth Canopy in 2017 and became a pioneer site for the Peat Swamp Forest Rehabilitation and Conservation Project in Southeast Asia and Malaysia. The drained PSF in Tanjong Karang had various tropical tree species common in peat swamp forests such as *Shorea* spp., *Koompassia* spp., and *Dipterocarpus* spp.. In contrast, the OPP in Sepang was characterized by a monoculture of oil palm trees (*Elaeis guineensis*), planted in orderly rows. The hydrological conditions of each site were influenced by their respective management practices. The PSF, previously drained, likely saw modifications in the natural water regime, leading to a potentially lower water table compared to its original state. The OPP's hydrological system was shaped by drainage networks and practices inherent to palm oil production, aiming to regulate water levels and soil moisture. The peat depth of the PSF was around 165.4 ± 3.9 cm (SD), while the OPP had a depth of approximately 95.2 ± 6.5 cm (SD), both considered moderate to shallow. USDA soil taxonomy classified the peat soils at both sites as Typic Fibric Tropohemist. In the drained PSF, the mechanical compaction experiment was positioned 1.0 km from the drain edge and entry point to mitigate the effects of passive compaction from corner sites due to the drainage system's development. Similarly, in the OPP, the experiment was placed between the palm planting rows, about 20 m from the main drain, to ensure data accuracy.

2.2 Fabrication of hybrid compaction apparatus

The compaction apparatus was fabricated using stainless steel based on a vertical plunger-like compression approach. This simple apparatus consisted of five key components: Fig. 2 (a) cover, (b) piston, (c) piston's ear, (d) soil collar, and (e) perforated mesh, with a total weight of approximately 3.0 kg. A soil collar cylinder

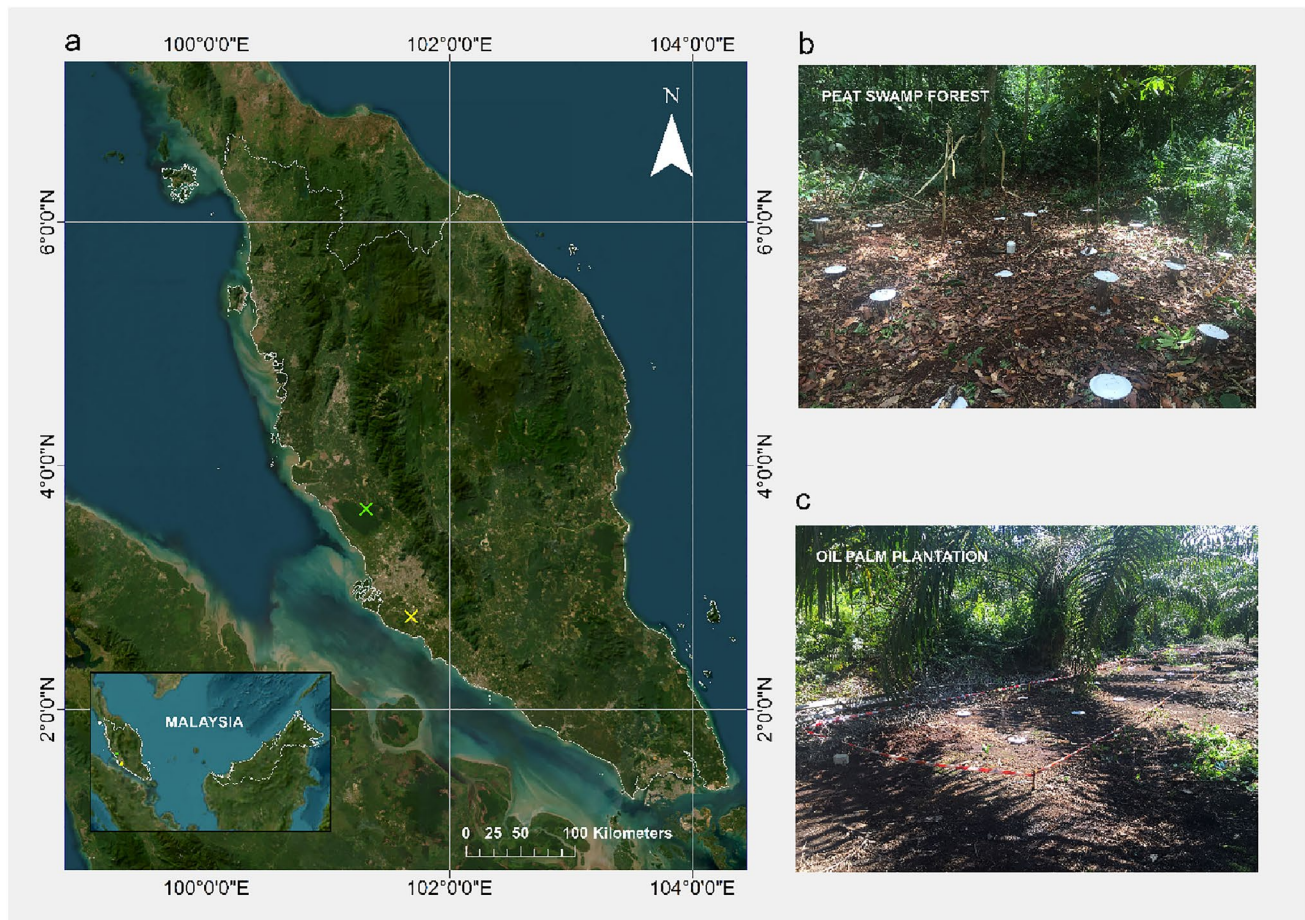


Fig. 1 **a** Map of Peninsular Malaysia that shows the locations of intact peat samples collected from two different ecosystems in Selangor state that are marked by colored X: **b** Peat Swamp Forest

with the green X and **c** Oil Palm Plantation with the yellow X. The map was generated using GIS software (ArcGIS Desktop 10.8, Version: 10.7.0.10450, <http://www.esri.com>).

measuring 40.0 cm long with a diameter of 10.16 cm, was used to standardize compaction magnitudes and maintain compaction degree throughout the experimental period—due to the peat’s high plasticity and recovery features [23]. A piston measuring 30.0 cm high, and 10.16 cm in diameter was used to compress the peat surface to the assigned soil depth. The piston also served as a static closed chamber for capturing gas. A piston cover that fitted two gas tubing ports was placed on top of the piston to facilitate gas sampling. The size of the piston was intentionally standardized for all treatments of gas concentration calculation (6.5 L). The piston’s ears, attached at various heights following the compaction treatment, were designed as a stopper during the compaction application. Finally, to address compressibility obstructions such as available water or air from peat removed by the compaction effort, the piston surface was perforated with \varnothing 1.0 cm meshes. This also allowed for the free transfer of peat carbon emissions.

2.3 Artificial compaction treatment

The soil collars were then driven into the peat profile and left for two months to allow resettlement and the root necrosis and exclusion of root biomass which would otherwise contribute to additional autotrophic emissions. Since flooding inside the soil collar was observed to occur due to rainwater, the soil collars were covered loosely using a polystyrene plate to prevent direct rainfall while also preventing vapor deficits that could affect the gas carbon emission. In addition, the slope of the soil was also considered. For OPP, sites used avoided areas within 3.5 m of the palm trunk to avoid the root cycle [29] and reduce microenvironment variability. The piston was left inside the soil collar after implementing the proposed depth to emulate static compaction by heavy machinery that commonly occurs in agricultural fields. To consider the mesh-perforated surface, which may interrupt the gas diffusion from the surface soil, an integrity test was done to evaluate the compaction apparatus

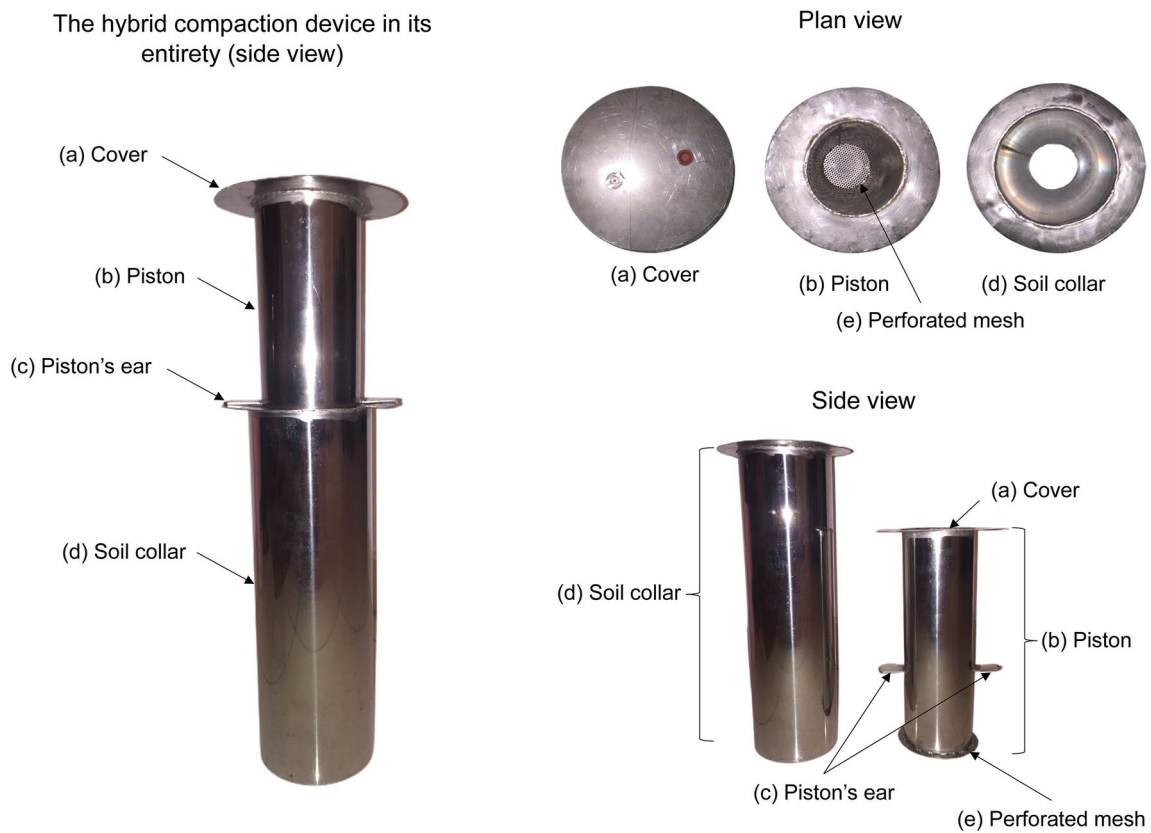
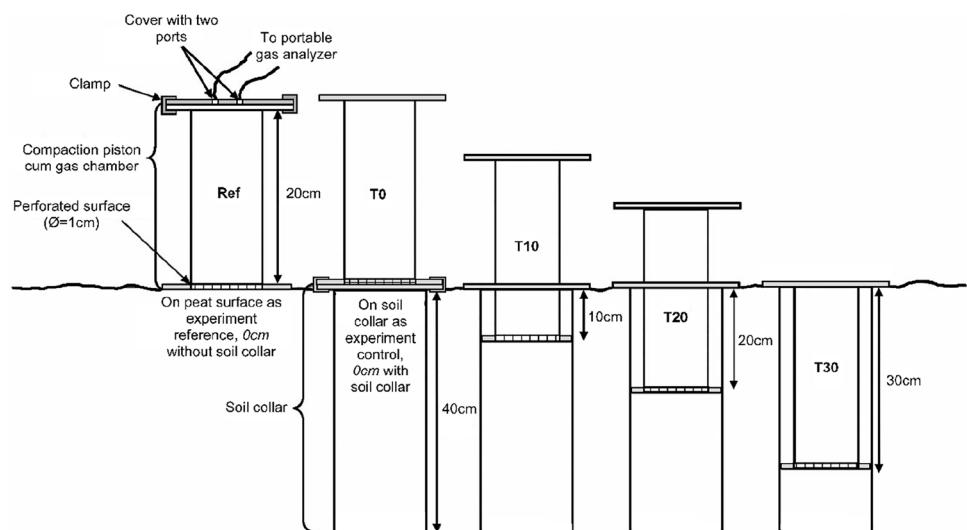


Fig. 2 Orthographic of the hybrid peat compaction apparatus used in this study

Fig. 3 Schematic of peat compaction apparatus installation at the peat swamp forest and the oil palm plantation



capability in terms of carbon emissions and degree of compactness (i.e., bulk density). The test was performed at the same time as the compaction experiment. Together, the compaction treatments were assigned as Ref (reference for control or no compaction without soil collar); T0 (no compaction with soil collar); T10 (soil

compression treatment to 10.0 cm); T20 (20.0 cm); and T30 (30.0 cm), as per Fig. 3. Four replications of the five compaction treatments (including the reference samples for the integrity test) in both ecosystems were established with a total number of samples of $n = 40$.

2.4 Gas and groundwater level measurement

Gas sampling was undertaken for 24 weeks for PSF and 21 weeks for OPP, over three-week intervals for each sampling point after mechanical compaction treatments. Gas carbon emissions were measured using an automated gas analyzer (Ultraportable gas analyzer by Los Gatos Research, USA). Mixed gas accumulation was analyzed in real-time for six minutes and each gas concentration point represented 20 s, which were recorded and stored automatically in the analyzer's data logger. Three concentration points for the first minute were discounted to account for the lid installation disruption.

Gas concentration in ppm was calculated by using ideal gas law, as given in Eq. (1).

$$PV = nRT \quad (1)$$

where P = Atmospheric pressure (101,325 pa); V = Volume of headspace (m^3); n = Number of moles (mol); R = Universal Gas Constant law ($8.314 \text{ J} \cdot \text{K}^{-1} \cdot \text{mol}^{-1}$) and T = Temperature in kelvin (K) with conversion number of 1 mol of gas to gram, $\text{CO}_2 = 44.01 \text{ g}$ and $\text{CH}_4 = 16.02 \text{ g}$, respectively.

The increase or decrease in the head-space concentration over time (hour, h^{-1}) over the head-space volume (L) from peat surface area covered (m^2) were fitted into linear regression (LR) [30]. The slope from LR (in this study $\sim R^2 > 0.90$) was considered as the rate of CO_2 emission or CH_4 emission (+ value) / uptake (-value) and presented in, $\text{mg} \cdot \text{m}^{-2} \cdot \text{h}^{-1}$ or $\mu\text{g} \cdot \text{m}^{-2} \cdot \text{h}^{-1}$. Groundwater levels were measured manually from piezometers (1.5 m) installed around the soil collars using rulers and distance laser beam that aimed to floating polystyrene. During gas sampling, all measurements were retrieved from three perforated piezometers and the average measurements were calculated.

To assess the relative impact of different greenhouse gas emissions on climate change, the values of CO_2 and CH_4 emissions were converted to a GWP index using Eq. (2). The GWP index is a way for comparing the contribution of different greenhouse gases to global warming over a certain period, usually 100 years. It considers the energy absorption capacity of each gas and its atmospheric residence time. The GWP index uses CO_2 as a reference, which has a GWP of 1 and assigns a GWP of 30 to CH_4 from fossil sources [31]. The results were presented in $\text{CO}_2\text{-eq}$.

$$\text{GWP} = \text{CO}_2 + \text{CH}_4(30) \quad (2)$$

2.5 Peat sampling and physicochemical properties analysis

The peat and bulk density samples were collected separately from the soil collars at the end of the compaction

experiment. Considering the small inner diameter of the peat surface inside the soil collar (i.e., 0.127 m diameter; 0.051 m^2 area), a bulk density core sub-sample with a volume of 50.0 cm^3 at 7.0 cm peat depth was harvested from the soil collar using an open-ended syringe (50.0 ml) modified earlier by cutting off the syringe's luer part (i.e., 0.027 m diameter with 0.0006 m^2 area). Then, 30.0 g of fresh peat was scooped out using a spatula for selected further peat physicochemical property analysis. In this study, peat surface sampling was emphasized because, within a range of 0 to 10.0 cm, peat soil exerts a major control on CO_2 and CH_4 emissions [32] and physicochemical properties, and this range is an effective depth for soil compaction [25].

In the laboratory, the analysis of dry bulk density was conducted using the core method prescribed by Al-Shammary et al. [33]. Soil cores containing fresh peat samples were weighed to obtain wet weight (W_w). Next, the samples were oven-dried at $105 \text{ }^\circ\text{C}$ until constant dry weight (D_w) was attained for 48 h. The oven-dried peat samples were placed in desiccators for cooling and to prevent moisture absorption. Alongside, the moisture content parameter was obtained as well. Both parameters were calculated using Eqs. (3) and (4) and expressed in $\text{g} \cdot \text{cm}^{-3}$ and percentage (%), respectively.

$$\text{BD} = \frac{D_w}{V} \quad (3)$$

$$\text{MC} = \left[\frac{W_w - D_w}{W_w} \right] \times 100\% \quad (4)$$

Loss on ignition (LOI) was determined from a small fraction of composite oven-dried peat (5.0 g) (D_w) via complete combustion by adhering to the method prescribed by Marwanto et al. [34]. After that, the weighed dried peat was transferred to muffle furnace (Thermo Scientific Thermolyne Industrial Benchtop Muffle Furnaces, Fisher Scientific, USA) at $550 \text{ }^\circ\text{C}$ for 12 h, including a cooling process to $60 \text{ }^\circ\text{C}$. Next, crucible cups containing ash were placed in the desiccator to gain weighed ash (DA) upon hitting room temperature at $30 \text{ }^\circ\text{C}$. The LOI was calculated using Eqs. (5) and (6) and expressed in percentage (%).

$$\text{DA} = \left[\left(D_w - \frac{W_a}{D_w} \right) \right] \times 100\% \quad (5)$$

$$\text{LOI} = 100\% - \text{DA} \quad (6)$$

Following the method suggested by Santos et al. [35], the value of particle density (DS) can be obtained using pycnometer or specific gravity flask (Phunque Flask). The peat particle density value was calculated using Eq. (7).

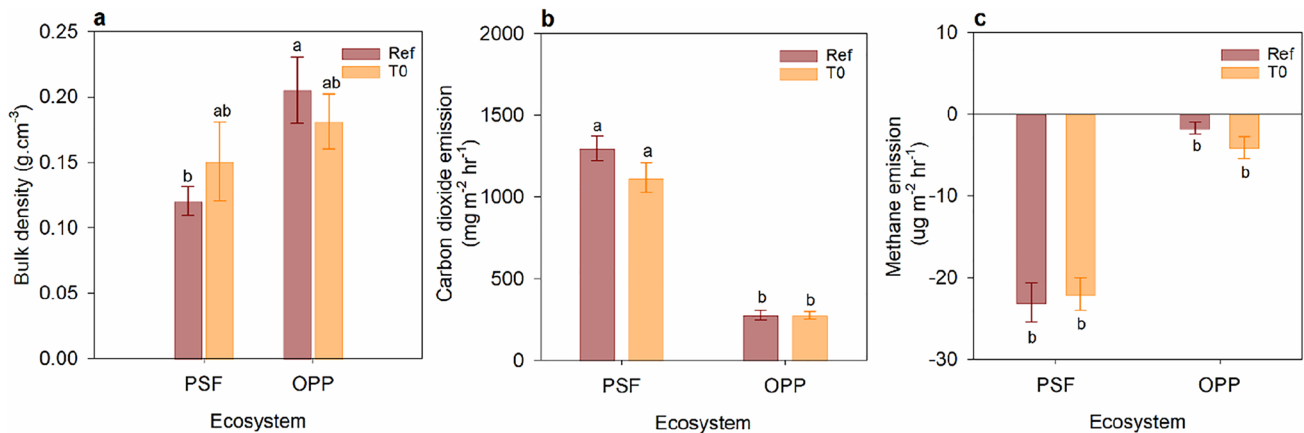


Fig. 4 The comparison of **a** bulk density **b** CO_2 and **c** CH_4 emissions between the reference and T0 samples at the PSF and the OPP. The error bars indicate the standard error of means. The different let-

ters indicate significant differences between treatments and land uses. The means were separated using Tukey's HSD post-hoc test ($\alpha=0.05$)

Where (D_w) refers to water density (g cm^{-3}) at the temperature; (W_s) denotes weight of peat soil sample; (W_{sw}) represents weight of pycnometer, peat soil, and water, while (W_w) stands for weight of pycnometer and water. The calculation of total porosity (TPS) had been based on peat particle density value using Eq. (8).

$$DS = \frac{D_w \times W_s}{(W_s) - (W_{sw} - W_w)} \% \quad (7)$$

$$TPS = 100 - \left(\left[\frac{BD}{D_s} \right] \times 100 \right) \quad (8)$$

Water-filled pore space ($WFPS$) was determined by using gravitational moisture content based on dry weight (GRA) and volume of moisture content (VOL) parameters, as given in Eqs. (9) and (10) [36]. The calculated values were used to estimate $WFPS$, as shown in Eq. (11).

$$GRA = \left[\frac{W_w - D_w}{D_w} \right] \times 100 \quad (9)$$

$$VOL = GRA \times BD \quad (10)$$

$$WFPS = (TPS \times VOL) \times 100 \quad (11)$$

The composite dry samples from soil core were crushed by mortar and pestle and passed through 2.0 mm sieve to determine their chemical properties [37]. The peat pH value was determined by using a benchtop pH meter model (Sartorius Benchtop Meters: pHBasic+ Series, Fisher Scientific, USA) by adding 5.0 g of dried and sieved peat sample into 12.5 ml deionized water (soil–water ratio of 1:2.5) in a 25.0 ml glass beaker. To determine the total carbon, total nitrogen, hydrogen, sulfur and CN ratio, the

samples were analyzed using the standard method of elemental analysis following ISO 13878:1998. This analysis was carried out using a CHNS analyzer (Vario MACRO cube, Elementary Analysis System, Germany).

2.6 Statistical analyses

A General Linear Model (GLM) was run to verify the integrity of the fabricated compaction apparatus and to assess the effect of the ecosystems and compaction treatments on the physicochemical properties and carbon emissions of peat. When statistically significant differences between parameters were observed, Tukey's HSD multiple comparison procedure at $p \leq 0.05$ was performed. To identify the relationship between physicochemical properties and water table fluctuation with carbon emissions, Pearson's correlation coefficient method, two-tailed at the 0.05 probability level, was performed. All statistical analyses were calculated and formulated with the help of Minitab software version 18.1 (Penn, USA, 2017) while the graphs were plotted using Sigmaplot 12.5 graphing software (Systat Software Incorporation, UK, 2014).

3 Results and discussion

3.1 Impact of fabricated compaction apparatus

An integrity test was conducted alongside the compaction treatment using uncompacted peat samples without a soil collar (Reference or 'Ref') and compared to a treatment control where the soil collar was installed but without compaction ('T0'). This comparison was carried out in both ecosystems. The analysis showed that there

were no significant differences in bulk density (used as an indicator for compaction), with values ranging from 0.12 to 0.15 g cm⁻³ for PSF and 0.18 to 0.21 g cm⁻³ for OPP (Fig. 4a). Additionally, there were no significant differences in carbon emissions in the form of CO₂, which ranged from 1234.3 to 1294.8 mg m⁻² h⁻¹ for PSF and 289.4 to 521.6 mg m⁻² h⁻¹ for OPP (Fig. 4b), and CH₄ emissions, ranging from -48.0 to -50.6 μg m⁻² h⁻¹ for PSF and -3.9 to -5.3 μg m⁻² h⁻¹ for OPP (Fig. 4c), between the Ref and T0 samples in both ecosystems. This suggests that the compaction apparatus used did not have a significant impact on bulk density or carbon emissions in the T0 (independently of any treatment), thereby allowing for subsequent analysis of compaction at varying magnitudes on physicochemical properties and carbon emissions.

3.2 Peat resilience in the presence of groundwater fluctuation

Notably, despite the depth-based compression applied by the compaction apparatus, there was no discernible relationship between the treatment and bulk density within each ecosystem, as indicated in Table 1. Similarly, it is evident that other physicochemical properties such as total porosity, gravimetric water content, total carbon, hydrogen, sulphur, and CN ratio remained unaffected by the compaction treatments. However, a significant disparity in bulk density between the two ecosystems is apparent, with the OPP showing a mean of 0.232 g cm⁻³ compared to that of the PSF, at 0.171 g cm⁻³. Correspondingly, peat porosity in OPP was lower than in PSF, with mean values of 85.5% and 89.7%, respectively. Furthermore, gravimetric water content was higher in PSF compared to OPP, with mean values of 225.3% and 175.3%, respectively. In OPP, total carbon, hydrogen, and CN ratios were higher than in PSF, with mean values of 42.5%, 13.8%, and 31.16,

Table 1 Physicochemical properties of the compaction treatments (mean ± standard error, n=4) between the two ecosystems. The difference in the uppercase (between ecosystems) and lowercase

(interaction effects) letters in bold in the same column indicates significant differences at p ≤ 0.05

Ecosystem	Treatment	Physical property					
		Bulk density (g cm ⁻³)	Loss on ignition (%)	Total porosity (%)	Water-filled pore space (%)	Gravimetric water content (%)	Volumetric water content (%)
PSF	T0	0.15 ^a ± 0.03	72.8 ^a ± 9.9	91.5 ^a ± 1.2	34.2 ^a ± 5.9	214.9 ^a ± 22.7	31.1 ^a ± 4.9
	T10	0.19 ^a ± 0.02	79.0 ^a ± 10.2	86.4 ^a ± 0.5	45.1 ^a ± 1.9	214.0 ^a ± 27.7	39.9 ^a ± 1.9
	T20	0.15 ^a ± 0.02	85.7 ^a ± 3.8	90.7 ^a ± 0.8	41.5 ^a ± 2.4	257.8 ^a ± 23.4	37.6 ^a ± 2.1
	T30	0.19 ^a ± 0.03	85.1 ^a ± 1.9	88.2 ^a ± 2.2	42.9 ^a ± 7.1	214.5 ^a ± 37.7	37.6 ^a ± 5.7
	Mean	0.17^B ± 0.01	80.7 ^A ± 3.6	89.7^A ± 0.7	40.9 ^A ± 2.4	225.3^A ± 13.6	36.5 ^A ± 2.0
OPP	T0	0.18 ^a ± 0.02	85.9 ^a ± 1.3	88.6 ^a ± 1.3	38.7 ^a ± 3.1	195.4 ^a ± 23.7	34.2 ^a ± 2.7
	T10	0.22 ^a ± 0.04	83.8 ^a ± 0.6	86.4 ^a ± 2.4	40.0 ^a ± 5.7	163.7 ^a ± 15.7	34.2 ^a ± 4.2
	T20	0.28 ^a ± 0.06	81.9 ^a ± 1.1	83.2 ^a ± 3.8	47.7 ^a ± 4.7	161.7 ^a ± 29.2	39.2 ^a ± 2.1
	T30	0.25 ^a ± 0.01	85.5 ^a ± 2.3	84.0 ^a ± 1.2	54.5 ^a ± 7.0	180.5 ^a ± 25.3	45.7 ^a ± 5.8
	Mean	0.23^A ± 0.02	84.3 ^A ± 0.8	85.5^B ± 1.2	45.2 ^A ± 2.9	175.3^B ± 11.3	38.3 ^A ± 2.1
Ecosystem	Treatment	Chemical property					
		pH value (1:2.5) (Peat:water)	Total carbon (%)	Total Nitrogen (%)	Hydrogen (%)	Sulphur (%)	C:N ratio (%)
PSF	T0	3.06 ^a ± 0.02	33.9 ^a ± 4.5	1.6 ^a ± 0.3	4.7 ^a ± 0.9	7.0 ^a ± 3.7	22.5 ^a ± 1.8
	T10	3.04 ^a ± 0.04	36.1 ^a ± 4.7	1.4 ^a ± 0.2	5.6 ^a ± 1.0	3.7 ^a ± 2.2	26.6 ^a ± 2.7
	T20	3.01 ^a ± 0.03	41.4 ^a ± 2.2	1.6 ^a ± 0.2	6.5 ^a ± 0.8	2.5 ^a ± 1.0	26.9 ^a ± 3.2
	T30	3.07 ^a ± 0.03	38.0 ^a ± 2.1	1.6 ^a ± 0.1	6.0 ^a ± 0.7	3.8 ^a ± 1.0	25.0 ^a ± 1.7
	Mean	3.05 ^A ± 0.01	37.4^B ± 1.8	1.5 ^A ± 0.1	5.7^B ± 0.4	4.3^A ± 1.2	25.2^B ± 1.2
OPP	T0	2.89 ^a ± 0.1	43.2 ^a ± 1.1	1.4 ^a ± 0.0	14.2 ^a ± 1.1	0.4 ^a ± 0.9	31.7 ^a ± 0.9
	T10	2.92 ^a ± 0.1	41.9 ^a ± 1.6	1.3 ^a ± 0.1	13.9 ^a ± 0.43	0.4 ^a ± 0.1	32.2 ^a ± 1.5
	T20	3.01 ^a ± 0.1	42.9 ^a ± 1.8	1.4 ^a ± 0.1	13.1 ^a ± 0.7	0.5 ^a ± 0.1	29.6 ^a ± 0.8
	T30	3.18 ^a ± 0.2	42.4 ^a ± 2.3	1.37 ^a ± 0.1	13.9 ^a ± 2.0	0.5 ^a ± 0.2	31.2 ^a ± 1.2
	Mean	3.00 ^A ± 0.1	42.5^A ± 0.8	1.37 ^A ± 0.0	13.8^A ± 0.6	0.44^B ± 0.1	31.16^A ± 0.6

respectively. Conversely, the sulphur content in PSF was higher than in OPP, with mean values of 4.26%, respectively. Notably, properties such as loss on ignition, water-filled pore space, volumetric water content, pH, and total nitrogen were not affected by either the ecosystem type or the compaction treatment.

The lack of response of peat properties to the compaction treatments was unexpected and suggested that at the forces applied here at least, the peat in both ecosystems was resilient towards mechanical compaction. There are several reasons why the peat from both ecosystems exhibited resilience. First, this effect is a result of groundwater fluctuation interaction, especially at maximum compaction, T30 (–30.0 cm compression), where the modified soil surface level was closest to the water table level. The water table level at maximum compaction during the study of PSF and OPP fluctuated from -6.26 ± 7.35 cm (SD) and -26.8 ± 8.59 cm (SD) when compaction went above –30.0 cm (Table 2). Groundwater level interference was also observed by Baker et al. [38], who suggested that groundwater level fluctuation in wetting and drying cycles could cause looseness or impaired pore structure or swelling. In addition, Pires et al. [39] also analyzed the changes in a pore system using gamma-ray computed tomography (CT) and found that wetting and drying cycles could result in a repair mechanism in the compacted soil structure. This is important when considering potential rebound post compaction. but does not account for the continuous compaction which occurred throughout the experimental period.

Second, the ineffectiveness of mechanical compaction on peat may be further enhanced by newly decomposed peat, or available labile C from above the peat surface, which has translocated to a deeper depth by vertical compaction and readily has sponge-like features [40]. Recent dead plant biomass is known as a source of labile C [41].

Table 2 Water table fluctuation according to the assigned treatments at the respective ecosystems

Treatment	Ecosystem	
	Peat swamp forest (PSF)	Oil palm plantation (OPP)
T0	-36.21 ± 7.34	-56.84 ± 8.59
T10	-26.21 ± 7.34	-46.84 ± 8.59
T20	-16.21 ± 7.34	-36.83 ± 8.59
T30	-6.21 ± 7.35	-26.83 ± 8.59

Values are presented in mean \pm (SD) of water table level at PSF ($n=32$) and OPP ($n=28$), respectively

T0=No compaction with soil collar; T10=compaction to –10 cm peat depth); T20=compaction to –20 cm peat depth; T30=compaction to –30 cm peat depth

During compaction treatment, labile C from the peat surface could be relocated to the deeper horizon to augment consolidation. Achieving instantaneously in this case is highly challenging due to the presence of labile C, which consists of intricate components derived from recently deceased plants with a sturdy structure that discourages herbivore-like traits [42, 43]. Hence, peat will return to its former state, as the peat pores could still be active and have elastic and plastic properties [22, 23]. Unless the peat pores are inactive due to excessive drying, which will manifest as hydrophobicity [24].

Third, peat failure to compress could be a result of a lack of mineral or clay content [44, 45], which serves as a cohesive or macropores filler to stabilize the anisotropic structure of peat. This is because mineral content such as fly ash, gypsum, and lime with peat admixtures can promote peat stabilization and could possibly support a more effective compaction [44–48]. If this were the situation, a more pronounced reaction to compaction treatments could be anticipated at the OPP sites, where lime is applied. The most plausible explanation may be that the compaction effort was insufficient to elicit a response by the physicochemical variables measured. According to Mutert et al. [5], the changes in bulk density may occur when the compaction of peat is drained to 0.50 m to 0.80 m from the peat surface, allowing self-subsidence to 1.0 m prior to heavy machine passes. Conversely, the induced compaction in this study was kept at a maximum of 30.0 cm depth, and the water table level was not controlled, which might not have been sufficient to initiate changes in physicochemical properties of peat, especially bulk density. Nevertheless, the prolonged drainage of the OPP site accounts for the consistently elevated bulk density values observed there.

Overall, while not expected, this result points to the impressive resilience of intact peatlands to compaction, and their ability to maintain intact pore structure, water-holding capacity, and porosity. These features are key to the hydrological ecosystem services provided by peatlands in terms of flood protection and water storage capacity [49]. This result also reinforces the direct role of drainage as the key driver of compaction and consolidation within peatlands [24], with the mechanical compaction only compounding the effect. The significant difference between ecosystems corroborate with the findings of several authors [50–54], who assessed the physicochemical changes of converted land from a peat swamp forest to an oil palm plantation. These typical changes are based on the origin ecosystem attributes and the different management involving the lowering of water table level by drainage, agricultural input (e.g., fertilizer, pesticide, and agri-implements), and litter input from surrounding plant biomass. Consequently, the distinct traits exhibited at the

OPP sites will also be associated with the degree of decomposition (e.g., bulk density, total porosity, water-filled pore space, gravimetric, and volumetric water content, and CN ratio) [36, 50].

Therefore, a key question to emerge from this work however is if drainage-based compaction is recoverable with rewetting. While the work from Pires et al. [39] suggests that the recovery of natural wet and dry cycles is key to the rehabilitation of pore structure, *in-situ* assessments of bulk density and pore structure recovery over time will need to be assessed in line with rehabilitation activities if key hydrological ecosystem services are to be recovered at previously drained and converted OPP sites.

3.3 Carbon emission changes due to mechanical compaction

Despite physicochemical properties being unaffected by site-specific mechanical compaction, carbon emissions in the form of CO₂ and CH₄ from the peat surface showed inverse patterns due to compaction, especially in the PSF ecosystem. During compaction, CO₂ emissions decreased according to the treatment, with the greatest reduction measured at T30, by 827.2 mg m⁻² h⁻¹ (64%) and followed by 465.1 mg m⁻² h⁻¹ (36%) at T20 (Fig. 5a). When combined, the average emissions decrease was 2% or 27.6 mg m⁻² h⁻¹ per cm of compression (when compressed beyond 20 cm). Various studies have reported CO₂ emissions from forest and oil palm plantations. The CO₂ emissions in this study ranged from 467.5 to 1294.8 mg m⁻² h⁻¹ and 290.18 to 521.6 mg CO₂ m⁻² h⁻¹ at SF and OP, respectively, which are similar to Dariah et al.

[55], Husnain et al. [56], and Melling et al. [53] (Table 3). However, compared to many other studies, the value for OPP emissions is generally low. For example, Matysek et al. [29] obtained a higher value in their study of the same oil palm plantation during dry season condition (means up to 1244.7 ± 149.2 (SE) mg m⁻² h⁻¹; Table 3) and found the heterotrophic emissions to be dependent on the age of the plantation, with those converted the longest (2nd generation) producing emissions comparable to this study (663.8 ± 102.2 (SE) mg m⁻² h⁻¹).

Contrary to CO₂, CH₄ emission showed a significant increase in T20 (from -49.7 to -7.5 μg m⁻² h⁻¹; 85%) and T30 (from -49.7 to 60.0 μg m⁻² h⁻¹; 224%), respectively, (by an average 150%) within the PSF ecosystem (Fig. 5b). Our

Table 3 Comparison of the CO₂ emissions of this study with other studies at secondary peat swamp forest (PSF) and oil palm plantation (OPP)

Ecosystem	Reported range (mg CO ₂ m ⁻² h ⁻¹)	Reference
PSF	1265 ± 110	This study (Integrity test)
OPP	405.5 ± 55.5	This study (Integrity test)
PSF	467.5 to 1294.8	This study (All treatments)
OPP	290.18 to 521.6	This study (All treatments)
PSF	366.4 to 1953.1	Melling et al. [53]
OPP	168.6 to 1227.6	Melling et al. [53]
PSF	411.0 to 981.7	Husnain et al. [56]
OPP	388.1 to 753.4	Husnain et al. [56]
OPP	389.0 to 435.9	Dariah et al. [55]
OPP	716 to 909	Matysek et al. [29]

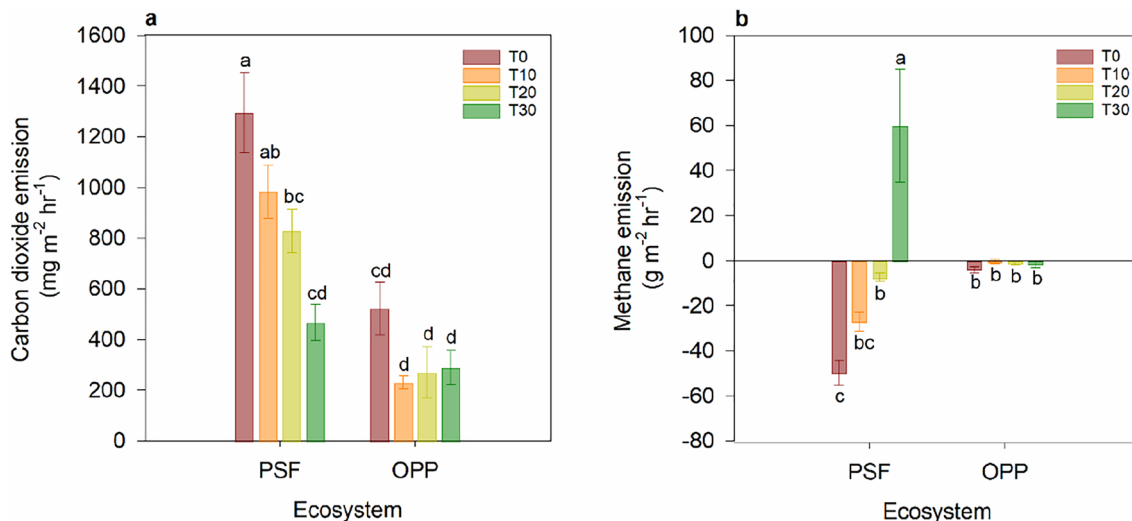


Fig. 5 The means of cumulative **a** CO₂ and **b** CH₄ emissions between compaction treatments at PSF (n = 32) and OP (n = 28). The error bars indicate the standard error of means. The different

letters indicate significant differences between treatments and land uses. The means were separated using Tukey's HSD post-hoc test (α= 0.05)

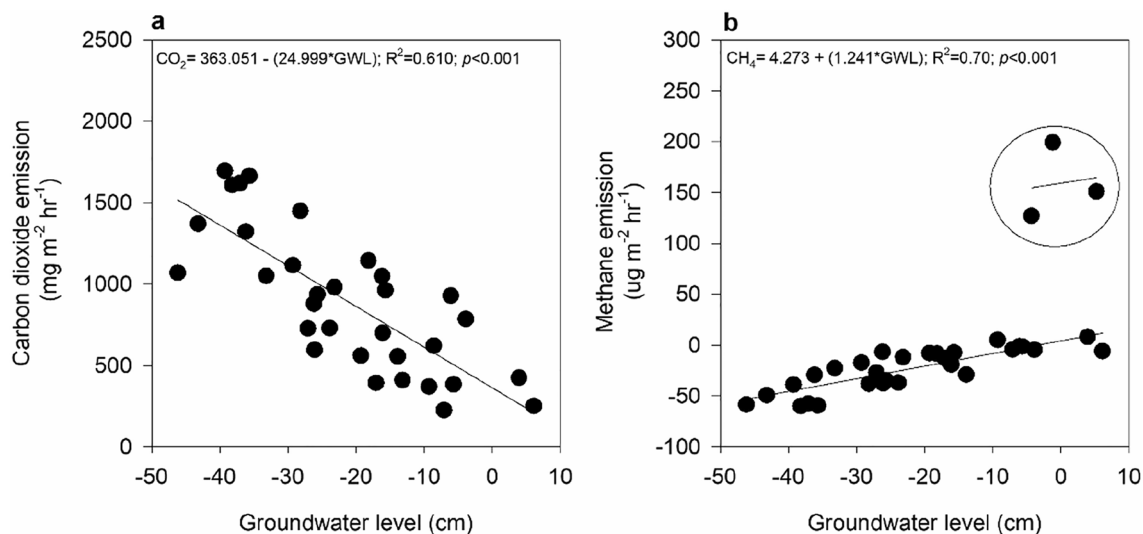
Table 4 Comparison of the CH₄ emissions of this study with other studies at secondary peat swamp forest (PSF) and oil palm plantation (OPP)

Ecosystem	Reported range ($\mu\text{g m}^{-2} \text{h}^{-1}$)	Reference
PSF	-49.3 ± 4.35	This study (Integrity test)
OPP	-4.64 ± 1.83	This study (Integrity test)
PSF	-48.4 to 59.95	This study (All treatments)
OPP	-3.94 to 3.63	This study (All treatments)
PSF	-6.03 to 11.19	Melling et al. [54]
OPP	-43.76 to 5.55	Melling et al. [54]
PSF	-23.2 to 27.86	Sakabe et al. [57]
Degraded SF	-9.63 to 10.4	Sakabe et al. [57]

results found uncompacted (T0) PSF to be a strong sink for CH₄ and OPP T0 to be a weak CH₄ sink, similar to Sakabe et al. [57] but opposed to the CH₄ emissions reported by Melling et al. [54] for both ecosystems (Table 4). This may have been due to their OPP site being newly converted (4 years) and thus potentially having a higher labile C surface peat content, a source of the substrate which has a CH₄ production priming effect [58, 59], as compared to our matured OPP. The effect of the compaction treatments worked to reduce the sink effect in both ecosystems. However, only PSF T30 became a mean source of CH₄ in this study. Given that there was no change in the physico-chemical variables with treatments, the only significant change between treatments was not the compaction (as indicated by bulk density value), but rather the increase in relative water table level (-36.21 at T0 and -6.21 at T30).

For this reason, we can use the data here to explore the direct relationship between emissions and groundwater levels from in-situ data. For CO₂ emission, a significant relationship was found in PSF ($p < 0.001$; $R^2 = 0.610$) (Fig. 6a). The CO₂ predicted emission from PSF is equal to $-24.9 \cdot \text{GWL} + 363.05$ when CO₂ emission is measured in $\text{mg m}^{-2} \text{h}^{-1}$. From this established model, the CO₂ emission can be reduced by $6.7 \text{ mg m}^{-2} \text{h}^{-1}$ for each centimeter of groundwater level (GWL) increased. On the other hand, a significant regression between CH₄ emission and GWL was also found at PSF ($p < 0.001$; $R^2 = 0.700$) (Fig. 6b) after eliminated pulse emissions (before was $\text{CH}_4 = 1.981 \cdot \text{GWL} + 35.98$ with $p = 0.008$; $R^2 = 0.212$). The predicted CH₄ emission in PSF is equal to $1.24 \cdot \text{GWL} + 4.27$ in centimeter when CH₄ emission is measured in $\mu\text{g CH}_4 \text{ m}^{-2} \text{h}^{-1}$. Thereby, CH₄ emission increased by $1.03 \mu\text{g CH}_4 \text{ m}^{-2} \text{h}^{-1}$ for each centimeter of GWL increased.

Therefore, to further explain the carbon reduction from peat in secondary peat swamp forests, a model illustrating the carbon emission exchange and GWL (Fig. 6) was employed. Thus, high, and low GWL scenarios (extreme inter-seasonal effect) were used representing high (-6.21 cm) and low (-66.21 cm) GWLs from the initial soil surface height (see Table 2 for water table mean at T0, control). The 30.0 cm value was derived from the effective change in carbon emission due to mechanical compaction, which affects peat depth. From the estimation presented in Table 5, 59.1% change in carbon emissions was observed when the water table was raised or lowered (e.g., inter-seasonal) by 30.0 cm .

**Fig. 6** Relationship between water table level and **a** CO₂ and **b** CH₄ emissions at PSF ($n = 29\text{--}32$). The point within solid circle indicates momentary emissions that eliminated from regression analysis.

Whereas solid line indicates regression line between water table level and carbon emission

Table 5 The estimation of carbon emission changes due to ground-water level in PSF of Tanjung Karang, Northern Selangor

Scenario	Carbon emission (t CO _{2eq} ha ⁻¹ yr ⁻¹)
Initial level (Mean)	111.0
High water table level (+ 30 cm)	45.4
Low water table level (−30 cm)	176.6

Interestingly, this estimation is similar to that of Dhandapani et al. [60], which measured CO₂ and CH₄ in a primary peat swamp forest in Setiu, Terengganu, and in a secondary peat swamp forest in Tanjung Karang, North Selangor. They observed a difference of more than 50% in CO₂ emissions. Thus, there is evidence to suggest that the secondary peat swamp forest in Tanjung Karang shares similar traits with the primary peat swamp forest in Terengganu, Peninsular Malaysia, which has been a protected area since the 1990s (27 years), particularly if the GWL has been raised to + 30.0 cm from the initial level. Besides, this is also indicating that rehabilitation of secondary tropical peat forest is progressing splendidly and in agreement with [61] that mentioned more than half of the tree species from primary forest can be found after 25 years of logging operations are ceased.

Hence, this study proposed that the GWL in the PSF of Tanjung Karang be maintained at 10.0 cm or higher to stabilise or reduce carbon emissions by 51.6%. However, it should be noted that the peat surface of this PSF will always produce carbon emissions due to the decomposition process of organic materials that are supplied continuously from available trees [41], besides root respiration [60]. Therefore, it is expected that approximately 53.7 t CO_{2-eq} ha⁻¹ yr⁻¹ will be emitted naturally by this forest, indicating a 10.0 cm GWL depth from the peat surface.

Yet importantly, this relationship only stands for relatively intact (i.e., PSF) peatlands. As illustrated, lower water table depths/greater drainage per se as found at all the OPP treatments (Fig. 7f) did not result in a greater sink potential for CH₄. Indeed, all OPP treatments were low CH₄ sinks despite having lower water level depths ranging from −56.45 (T0) to −26.83 (T30). In the OPP ecosystem, the litterfall input is undoubtedly limited [36] because of its monoculture system [62]. Despite ample oxic and anoxic conditions due to water table fluctuation, insufficient litter input could limit the source of substrate availability that promotes microbial activity [62]. Moreover, the effect of excessive periodical drainage in the matured OPP may also affect the water holding capacity resulting from irreversible drying condition [63] and an enhanced hydrophobic trait [24] due to disconnecting peat pores from water table

level in long period [52]. This phenomenon could explain why CO₂ and CH₄ at OPP (Fig. 7b, d) remain unchanged although the water table fluctuates and through depth alteration by compaction treatment.

Consideration of temporal water table fluctuations, of particular interest, is in the event of water table exceeds the peat surface on WAC 6 and WAC 9 in PSF with a water level difference of 20.1 cm (Fig. 7a, c, e) specifically on both carbon emissions. This may reflect the temporary water retention induced by anisotropic deformation on deeper peat [23]. In turn, the condition is likely to trigger temporary carbon emissions. Similarly, the CO₂ and CH₄ production was also found by Brouns et al. [64] during oxygenation treatment of deep peat layers that had not previously been exposed to air. This pulse and resulting emissions are often discounted in the long-term calculations (IPCC) [27, 52]. Therefore, the mechanism of changes in water tables above the ground level accompanied by mechanical compaction is important when undertaking a full C cost accounting for the process of oil palm conversion (as opposed to just a comparison of oil palm and forest). This pulse in emissions in association with the process of conversion agrees with that described by Hooijer et al. [17], but is crucially not accounted for in the IPCC 2014 accounting [65]. Nevertheless, the frequency is limited, and further study is needed to verify the reproducibility of the results reported here.

3.4 Implication of mechanical compaction to current OPP management practice

No correlation between carbon emission and physico-chemical properties was found when the data within the ecosystem was analyzed. Interestingly, the pooled data of physicochemical properties from both ecosystems showed several associations with carbon emissions (Table 6). These relationships can be interpreted in regard to the degree of decomposition resulting from oxidation and shrinkage. However, increases in bulk density and its decomposition proxies owing to ecosystem shifting, from PSF to OPP should not be interpreted as a result of mechanical compaction, but rather due to oxidation and shrinkage from periodic drainage alone [17, 66]. These coupled components have given rise to an important question regarding the actual definition of mechanical peat compaction. This issue could result in the misconception of mechanical compaction based on the use of heavy machinery during the opening of the plantation from a forest area. Here, through this study also, it is believed that mechanical compaction effectiveness could be deteriorated by the standard water table level from −40.0 cm to −60.0 cm [67],

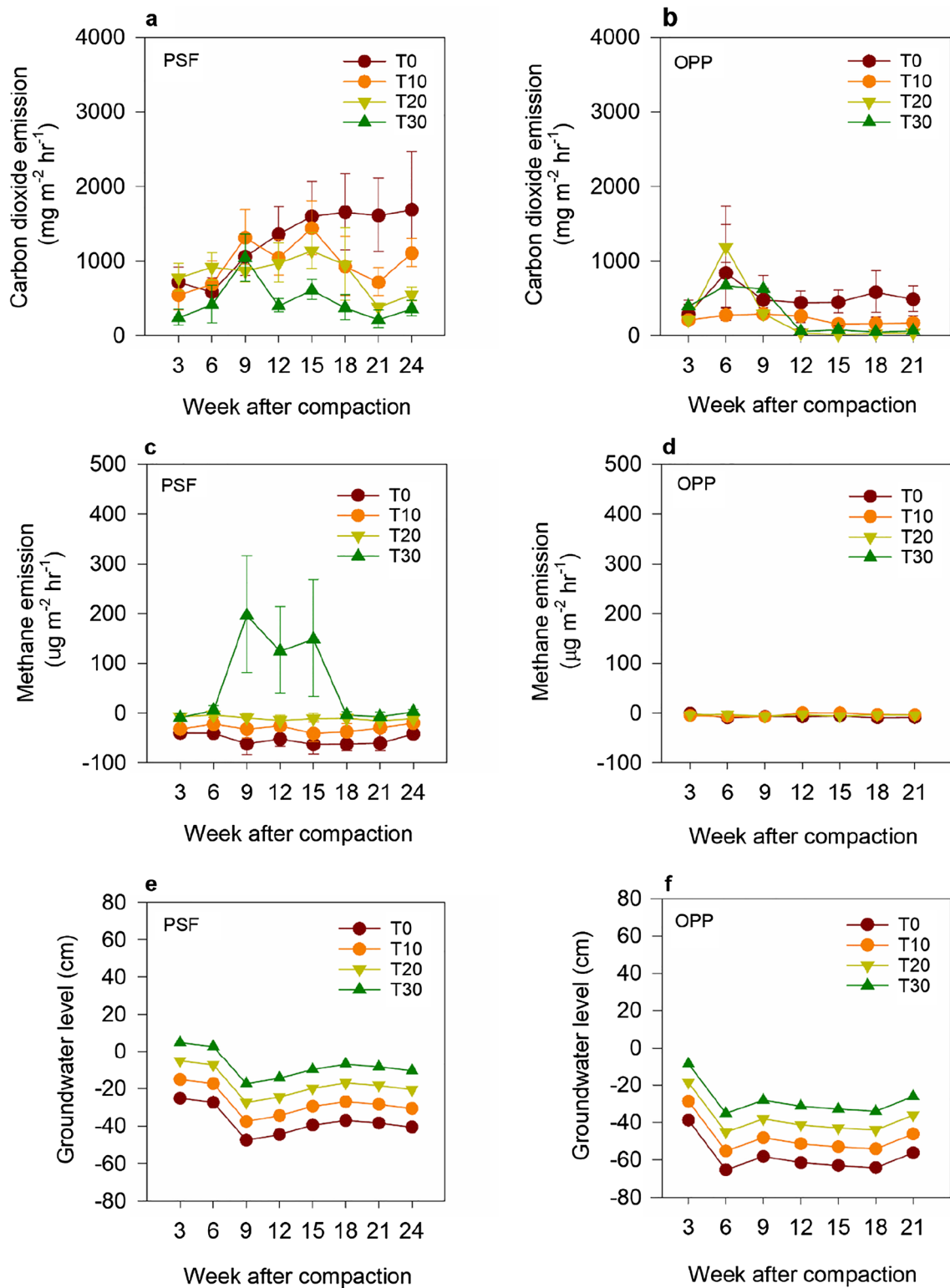


Fig. 7 The left and right figures display weekly measurements of CO₂ and CH₄ emissions, as well as groundwater levels, taken following compaction treatments in a peat swamp forest (PSF) and an oil palm plantation (OPP), respectively

Table 6 Pearson's correlation moment product between CO₂ and CH₄ emissions and physicochemical properties. The values represent Pearson's correlation coefficient, *r*, between carbon emissionsand peat physical properties (*n*= 4) of pooled data from both ecosystems. The *r* values in bold with *, **, and *** are significant at 0.05, 0.01, and ≤ 0.001, respectively

Carbon emission	Physical property					
	Bulk density (g cm ⁻³)	Loss on ignition (%)	Total porosity (%)	Water-filled pore space (%)	Gravimetric water content (%)	Volumetric water content (%)
CO ₂	-0.447**	-0.151	0.449**	-0.419*	0.279	-0.381*
CH ₄	0.394*	-0.052	-0.433*	0.395*	-0.213	0.363*
Carbon emission	Chemical property					
	pH value (1:2.5) (Peat:water)	Total carbon (%)	Total nitrogen (%)	Hydrogen%	Sulphur%	C:N ratio
CO ₂	-0.047	-0.149	0.398*	-0.436*	0.173	-0.528**
CH ₄	0.047	0.213	-0.250	0.387*	0.138	0.444*

provided that the peat has a higher proportion of labile carbon at the early stages of the land opening period.

4 Conclusions

The current research marks an initial exploration into the mechanical compaction of tropical peatlands, which is particularly insightful for two distinctive ecosystems: PSF and OPP. Our findings reveal that mechanical compaction significantly affects the levels of greenhouse gases emitted from these ecosystems, driving a decrease in CO₂ and an increase in CH₄, specifically during periods of high-water tables as noted within the PSF sites. Despite these changes, other physiochemical aspects in both ecosystems remained unaffected, likely due to the slacking effect attributed to the high compressibility and low clay content of the peatlands. Noteworthy is the establishment of a strong correlation between water table depth and emission levels. This relation could prove invaluable for predicting future emissions from peat swamp forests. Although additional compaction in areas that have been naturally compacted due to drainage for agricultural conversion (like the OPP sites) showed no significant changes in physiochemical parameters or carbon emissions, it hints towards a potential bulk density threshold in compaction effects. While it is essential to distinguish the effects of compaction from those of groundwater levels, the undetectable *in-situ* effect of compaction in already converted areas is worth acknowledging. Further compaction is inadvisable when it has a tangible impact. More comprehensive research is required to delineate these effects and to explore the limitations of compaction's impact on tropical peatlands. These findings contribute to our understanding of the

environmental implications of peatland management and can guide future research and mitigation efforts in these critical ecosystems.

Author contributions All the authors contributed to the study conception and design. Material preparation, investigation, data collection, and analysis were performed by M.K.S. The first draft of the manuscript was written by M.K.S, and all authors (M.K.S and S.L.E) commented on the previous versions of the manuscript. All authors read and approved the final manuscript.

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Data availability All data generated or analysed during this study are included in this published article and its supplementary information files.

Declarations

Conflict of interest The authors have no competing interests to declare that are relevant to the content of this article.

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