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1 **Oil palm 'slash-and-burn' practice increases post-fire greenhouse gas emissions and**
2 **nutrient concentrations in burnt regions of an agricultural tropical peatland**

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21 **Abstract**

22 Fire is one of the major issues facing Southeast Asian peatlands causing socio-economic,
23 human health and climate crises. Many of these fires in the region are associated with land
24 clearing or management practices for oil palm plantations. Here we study the direct post-fire
25 impacts of slash-and-burn oil palm agriculture on greenhouse gas emissions, peat physico-
26 chemical properties and nutrient concentrations. Greenhouse gas (GHG) emissions were
27 measured using Los Gatos ultraportable greenhouse gas analyser one month after a fire in dry
28 season and five months after the fire event, in wet season. Surface soil samples were collected
29 from each individual GHG measurement points, along with 50 cm cores from both burnt and
30 non-burnt control areas for lab analyses. As an immediate post-fire impact, carbon dioxide
31 (CO₂) and methane (CH₄) emissions, pH, electrical conductivity, and all macronutrient
32 concentrations except nitrogen (N) were increased multi-fold, while the redox potential, carbon
33 (C) and N content were greatly reduced in the burnt region. While some of the properties such
34 as CO₂ emissions, and electrical conductivity reverted to normal after five months, other
35 properties such as CH₄ emissions, pH and nutrient concentrations remained high in the burnt
36 region. This study also found very high loss of surface peat C content in the burnt region post
37 fire, which is irreversible. The results also show that surface peat layers up to 20 cm depth were
38 affected the most by slash-and-burn activity in oil palm agriculture, however the intensity of
39 fire can vary widely between different oil palm management and needs further research to fully
40 understand the long term and regional impacts of such slash-and-burn activity in tropical
41 peatlands.

42 **Keywords:** tropical peat fire; oil palm; carbon dioxide; methane; land use change; burnt
43 peatlands.

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46 **1. Introduction**

47 One of the most fundamental challenges of the 21st century is to meet the global food and
48 energy needs of the growing population, whilst conserving nature and soil carbon (C) storage
49 to mitigate climate change (Godfray et al., 2010). Oil palm is considered as a crop for the future
50 to serve these global needs, owing to its high productivity and versatility in usage (Abdullah et
51 al., 2009; Corley, 2009; Tan et al., 2009). Oil palm has become the most consumed vegetable
52 oil in the world over the past decade, and already represents 30% of the global usage (Lam et
53 al., 2019). However, as oil palm is an equatorial crop, this has come at the expense of some of
54 the world's most biodiverse and C rich ecosystems (Corley, 2009; Koh and Wilcove, 2008;
55 Murdiyarso et al., 2010; Murdiyarso et al., 2019). For the same reason, oil palm has a higher C
56 footprint than most other vegetable oil crops (Schmidt, 2015), with its expansion into tropical
57 peatlands further increasing its C footprint (Reijnders and Huijbregts, 2008).

58 Oil palm is native to West Africa, and was first introduced to Malaysia as an ornamental plant
59 in 1875 (Abdullah et al., 2009) and then commercially as a crop of choice for agricultural
60 diversification in 1917 (Rashid et al., 2013). The initial expansion of oil palm plantations was
61 in mineral soil at the expense of other agricultural plantations such as cocoa, rubber and coconut
62 (Basiron, 2007; Dhandapani, 2015). Since then, oil palm has rapidly expanded in Southeast
63 Asia, greatly expanding into peatlands from early 1990s (Miettinen et al., 2012; Shevade and
64 Loboda, 2019). As a consequence, Malaysia had the highest 21st century deforestation rate
65 globally (Hansen et al., 2013). A forested land should be completely cleared of vegetation,
66 drained and levelled to establish an oil palm plantation (Luskin and Potts, 2011). The lowering
67 of water table in peatlands make them highly susceptible for fire, as the dry peat itself is highly
68 flammable (Posa et al., 2011). Unlike contexts involving mineral soils, peat fires can sustain
69 both above and below ground, profoundly affecting surface and subsurface biota and seed bank,
70 damaging the vegetation structure (Posa et al., 2011; Turetsky et al., 2015; Uda et al., 2017).

71 The forest fires in Southeast Asia are often associated with land clearing for oil palm
72 plantations (Chazdon, 1998; Dennis et al., 2005). An estimated 60,000 peatland fires have
73 occurred in Indonesia between 1997 and 2007, causing harmful haze (Tan et al., 2009). About
74 2.6 million hectares of land was burned in Indonesia between June and October of 2015, which
75 led Indonesian president to cancel existing concession on agricultural expansion and make
76 conversion of peatlands to agricultural land illegal (World Bank, 2016). However, despite this,
77 fires and the use of fire in peatland agriculture persists in both Malaysia and Indonesia.
78 Furthermore, the use of fire is not just confined to the initial conversion and establishment
79 stage, it is also commonly used in 'slash-and-burn' to burn the waste from previous generation
80 oil palms to clear land for the next generation of crops. Slash-and-burn practice has been used
81 all round the world as a quick short term activity to shift agriculture or for land clearing
82 (Myllyntaus et al., 2002). Even though under certain circumstances in some ecosystems, slash-
83 and-burn practice can be sustainable (Kleinman et al., 1995; Myllyntaus et al., 2002), it is
84 certainly not sustainable in tropical peatlands in any circumstance, as dried peat itself is highly
85 flammable, and fundamentally any agriculture involving drainage itself in tropical peatlands is
86 not sustainable (Evers et al., 2017).

87 Fires in such tropical peatlands can convert large amount of C stored in peat for millennia into
88 carbon dioxide (CO₂) and cause socio-economic, human health, pollution and climate crises
89 (Cheong et al., 2019; Page et al., 2002; Turetsky et al., 2015; Wiggins et al., 2018). Tropical
90 peatlands in their natural undisturbed state are fire resistant, and historically fire had not played
91 any significant role in tropical peatland ecology (Turetsky et al., 2015). Most of these fires
92 have anthropogenic origin such as clearing of forest land for oil palm or similar agricultural
93 use, and slash-and-burn agricultural practices (Page et al., 2002). As tropical peatlands are
94 naturally acidic, this deliberate use of fire for agricultural practices brings short-term
95 advantages such as increase in pH and reduced cost for land conversion (Islam et al., 2016).

96 Though there were some research and estimates of direct emission from fire as smoke, haze or
97 CO₂ (Cheong et al., 2019; Hu et al., 2019; Hu et al., 2018; Islam et al., 2016; Smith et al., 2018;
98 Varkkey, 2013; Varkkey, 2016; Wiggins et al., 2018), the subsequent impact of fire on peat
99 physico-chemical properties and long term peat surface GHG emissions are not well
100 documented.

101 Peat fires can continue for a long time, lasting for months in the form of smouldering, persistent
102 flameless low-heat combustion of organic matter (Ohlemiller, 1985; Rein, 2013). Smouldering
103 in peatlands can be a complex process with interactions between aerobic and anaerobic
104 conditions, where pyrolysis of porous fuel occurs in high temperature at anaerobic conditions,
105 producing char (Ohlemiller, 1985; Rein, 2013) and the produced char is then oxidised when it
106 comes in contact with atmospheric oxygen, simultaneously producing and burning char during
107 the process (Rein, 2013). These complex interactions of aerobic and anaerobic conditions
108 along with other variations in intensity of fire and peat moisture levels can greatly impact peat
109 physico-chemical properties such as pH, electrical conductivity and redox potential, yet these
110 properties are not well documented for fire affected tropical peatlands.

111 The limited number of studies that do exist, have shown that peat fire significantly reduces the
112 organic matter content of the peat soil (Sazawa et al., 2018), as that very organic matter is lost
113 as C gases during the fire (Wiggins et al., 2018). Sazawa et al. (2018) also found that fire and
114 resultant heat dehydrates and denatures organic matter, which leaves fire affected peat more
115 susceptible to greater C losses, either through oxidation or repeated fire. Fire is generally found
116 to increase both total and available nutrient concentrations in mineral soils due to disintegration
117 of complex forms held in plants, and ash addition (Giardina et al., 2000; Van Reuler and
118 Janssen, 1993). This increase in available nutrients make soils more fertile, which is another
119 motive for the practice of slash-and-burn agriculture (Giardina et al., 2000). There are also
120 instances when total nutrient concentrations increased immediately after fire and then levelled

121 off after few months (Kutiel and Naveh, 1987). The decrease in nutrient concentrations after
122 few months may be because some of the nutrients are highly prone to leaching (Beliveau et al.,
123 2015; Ulery et al., 1993). Similar increase immediately after a fire was also observed for other
124 physico-chemical characteristics such as pH and electrical conductivity because of the addition
125 of ash (Bang-Andreasen et al., 2017; Gay-Des-Combes et al., 2017), before reverting back to
126 normal over time (Kutiel and Naveh, 1987). Fire may also effect increased mineralisation
127 indirectly through soil microbial processes (Gay-Des-Combes et al., 2017). This is especially
128 important in nutrient poor and acidic tropical peatlands, where fire is seen as a short-term and
129 cost-effective solution to make peat more cultivable. Naturally, tropical peat is a very complex
130 soil system and many of the environmental and microbiological interactions are understudied
131 and not well known (Dhandapani et al., 2019c). The interactions of fire in such a complex soil
132 system that is, in itself flammable (Rein, 2013; Uda et al., 2017), needs a more intensive and
133 greater number of research with varying environmental parameters to comprehend and infer a
134 cause and effect pattern.

135 In this study, we aim to evaluate the direct post-fire impact of slash-and-burn agriculture on
136 peat physico-chemical properties, nutrient concentrations and GHG emissions. We hypothesise
137 that fire increases both CO₂ and methane (CH₄) emission as an immediate post-fire effect. We
138 also hypothesise that fire significantly affects peat physico-chemical properties and nutrient
139 concentrations at surface layers, driven by high heat denaturing organic matter and the addition
140 of ash increasing pH and nutrient concentrations. We postulate that changes in peat physico-
141 chemical properties exhibit significant correlation with GHG emissions.

142 **2. Materials and Methods**

143 **2.1. Study site and sampling strategy**

144 The study site ($3^{\circ}25'17.7''\text{N}$ $101^{\circ}18'44.4''\text{E}$) is located at Kampong Raja Musa village in North
145 Selangor peatlands, Malaysia (Fig 1). The climatic conditions of the site is tropical, and the
146 soil classification comes under Histosols. The North Selangor peatland complex is the second
147 largest area of peatland in Peninsular Malaysia, and is located adjacent to Thennamaram region,
148 where oil palm was first commercially planted in Malaysia. The site is bordered with oil palm
149 monocropping, polyculture consisting of oil palm, yam and pineapple, pineapple
150 monocropping under an electric pylon trail, and a gravel road on the fourth side. The study site
151 and all the neighbouring land-use type blocks are roughly 2 ha in size each, consistent with
152 other small-holdings in the village (Dhandapani et al., 2019a,b). Two drainage ditches run as
153 two parallel borders with the oil palm monoculture, and the oil palm, yam and pineapple
154 polyculture. There is no drainage ditch within the site itself (Fig 2). The previous generation of
155 oil palm monoculture was cleared and the waste such as dead wood and fronds were stacked
156 in two parallel lines within the site. The stacked piles were burnt on June 2018, but with the
157 fire front extending into the peat itself, and smoke still visible from smouldering of the peat
158 surface for few weeks. The depth of burn was uneven throughout the burnt region, with surface
159 reaching mineral layer in some parts, and some other parts still containing peat to a depth of
160 atleast 50 cm from surface. New young oil palm rows were planted in non-burnt area in July
161 2018. Alongside this, a pineapple crop was planted in all open area in the non-burnt region and
162 a banana crop was planted in rows in the burnt area, in between dry season (July 2018) and wet
163 season (December 2018) sampling. Complete random sampling was carried out with 20 GHG
164 measurements and surface peat (0-5 cm) collection each for burnt and non-burnt region during
165 August 2018 dry season.

166 All sampling points were at least 1 metre away from each other. The same method was used
167 for December 2018 wet season sampling with reduced number of measurements to 10 each for
168 burnt and non-burnt region. The dry season measurement points were not the exact

169 measurement points used for the wet season, however they are in the same general area of burnt
170 and non-burnt locations used for the wet season measurements. The measurement points from
171 the first sampling were not marked with collars in order to minimise the disturbances from
172 disturbance from our field measurements, and hence a different set of random measurement
173 points in the same region were used for wet season sampling. Weather information for each
174 sampling period is given in Table 1. During the dry season sampling, 3 peat cores to a depth of
175 50 cm were collected for each burnt and non-burnt region. In some areas of burnt area peat was
176 completely burnt and mineral layers was reached at the surface. Selective sampling was carried
177 out for core collection, unlike complete random sampling for GHG measurements and
178 associated surface peat collection. The 3 peat cores for burnt region were collected from burnt
179 areas with peat remaining and not the areas with surface mineral exposure, to characterise the
180 impact on leftover peat through the peat depth profile.

181 **2.2. Greenhouse gas emissions**

182 CO₂ and CH₄ emissions from the peat surface were measured using a Los Gatos (San Jose,
183 California, USA) ultraportable greenhouse gas analyser as described in Dhandapani et al.
184 (2019a,b,c). The gas measurements were made using closed dynamic chamber method with an
185 inlet and outlet connecting to the gas analyser. The dynamic chamber was inserted up to 1 cm
186 into the peat during each measurement, to seal the chamber to the surface. The chamber was
187 15 cm high with 27 cm diameter. The Los Gatos gas analyser was set to record the gas flux
188 changes within the chamber every 20 seconds, and measurement was made at each sampling
189 point for 3 minutes, where the measurements from the first minute were ignored allowing the
190 gas concentrations inside the chamber to settle down after initial disturbance. The gas
191 concentrations were then converted to mg CO₂ m⁻² hr⁻¹ and µg CH₄ m⁻² hr⁻¹ for CO₂ and CH₄
192 respectively, as described in Dhandapani et al. (2019a,b,c).

193 **2.3. Peat analyses**

194 The procedure used for peat analyses were based on Dhandapani et al., (2019a,b,c). Peat
195 temperature was measured *in situ*, using a digital thermometer Cosmark PDT300 (Norwich,
196 UK). Peat samples were collected for measuring gravimetric moisture. For this, fresh peat was
197 dried in an oven at 105°C for 48 hours. The gravimetric moisture was calculated as follows:

198 Gravimetric moisture (%) = Mass of the water lost in oven drying/ mass of oven dried peat

199 For pH, redox and electric conductivity measurements, 5 mL volume of peat sample was
200 diluted in 10 mL deionised water in a centrifuge tube and shaken on a table shaker for 2 hours.
201 The pH of the supernatant was then measured using a Eutech pH700 pH meter supplied by
202 Thermo scientific (Loughborough, UK). The redox potential and electrical conductivity were
203 measured using Eutech Ion 2100 (Thermo scientific, Loughborough, UK) and Groline
204 HI98331 probe (Hanna, Leighton Buzzard, UK), respectively.

205 For analysing total C and nitrogen (N) content, all samples were oven dried (105°C for 48 h)
206 and finely ground using a Fritsch mortar grinder pulveristte 2 (Brackley, UK). Approximately
207 70 mg of sample was weighed into a Skalar ceramic crucible and the exact weight was
208 recorded. The samples were then transferred to an auto sampler in Skalar primacs series
209 SNC100 TC TN analyser (Breda, The Netherlands) and analysed for C and N content.

210 **2.4. Nutrient analyses**

211 Total nutrient concentrations of phosphorus (P), calcium (Ca), magnesium (Mg), sulphur (S)
212 and potassium (K) in peat were analysed using inductively coupled plasma mass spectroscopy
213 (ICP-MS). For this, approximately 0.15 g of oven dried (105°C for 48 h) and ground peat were
214 weighed in microwave digestion tubes (MARSXpress vessels, CEM Microwave Technology
215 Ltd., Buckingham, UK.). The digestion tubes are sealed with a stopper and a screw lid, after
216 adding 10 mL of nitric acid to each sample. The digestion tubes were then placed in a

217 MARSXpress microwave (CEM Microwave Technology Ltd., Buckingham, UK.) and run at
218 1600 W & 100% power with a ramp for 20 minutes and held for 20 minutes at 170°C. The
219 tubes were left overnight in the microwave to cool down. The digested samples were then
220 filtered and made up to 30 mL using milliQ water. Then, 1 mL of each sample were transferred
221 in to 10 mL tube and further diluted with 9 mL of milliQ water. The samples were then analysed
222 using 'Agilent Technologies (Milton Keynes, UK) 7900 ICP-MS' fitted with 'SPS 4'
223 autosampler.

224 **2.5. Statistical analyses**

225 All the statistical analyses were carried out using Genstat[®] 17th edition (VSN international,
226 2017). The significance of differences between sites for greenhouse gas emissions, nutrient
227 concentrations and other physico-chemical properties were evaluated using linear mixed
228 models with restricted maximum likelihood (REML) incorporating conditions (burnt or non-
229 burnt) and season (August and December 2018) as fixed affects. Similar REML was also
230 performed incorporating condition and depth as fixed effects, to identify the changes with
231 depth. For the data sets that were not normally distributed, the data were log transformed. For
232 data that did not meet normality assumption after log transformation, Boxcox transformation
233 was used. Normality was assessed by visual examination of 4 different residual plots, namely
234 histogram of residuals, fitted-value plot, normal plot and half-normal plot. Backward stepwise
235 multiple regression was performed with CO₂ and CH₄ as response variables and nutrient
236 concentration and other physico-chemical properties as fitted terms. Statistical significance
237 was assessed at $p < 0.05$ for all analyses.

238 **3. Results**

239 **3.1. Greenhouse gas emissions**

240 CO₂ emissions in the burnt region were twice as high as the CO₂ emissions from the non-burnt
241 region in the dry season one month after the fire incident (Fig 3a; Table 2). In wet season, five
242 months after the fire incident, CO₂ emissions in the burnt region levelled off to the same level
243 as non-burnt region. The difference between the seasons were significant, while the interactions
244 between regions and season were not significant (Table 2).

245 CH₄ emissions also varied significantly between the burnt and non-burnt regions (Fig 3b; Table
246 2). CH₄ emissions were multi-fold higher in the burnt region than in the non-burnt region for
247 both seasons, while the seasonal variations and interaction terms were not significant (Table
248 2).

249 **3.2. Surface peat/soil properties**

250 Peat surface temperature did not significantly differ between burnt and non-burnt regions, with
251 significantly higher temperature in dry season than in wet season (Fig 4a; Table 2). The wet
252 season temperature was slightly yet significantly different between the regions, resulting in
253 significant interactions terms. Gravimetric moisture did not significantly vary between the
254 burnt and non-burnt regions, while both regions had significantly higher gravimetric moisture
255 content in wet season than in dry season with no significant interaction between region and
256 season (Fig 4b; Table 2). During the dry season, electrical conductivity was more than 3 times
257 higher in the burnt region than in the non-burnt region resulting in significant difference
258 between the regions. Electrical conductivity was greatly reduced in wet season compared to
259 that of the dry season, resulting in significant seasonal variation (Fig 4c; Table 2). During the
260 wet season there was no significant difference in electrical conductivity between the two
261 regions while the difference was significant in dry season, resulting in significant interactions
262 (Fig 4c; Table 2). Both redox potential and pH were significantly higher in the burnt region

263 than in the non-burnt region, with no significant difference between seasons nor significant
264 interactions (Fig 4d,e; Table 2).

265 **3.3. Surface peat/soil C and nutrient content**

266 C content in the non-burnt region was more than double the content in the burnt region, with
267 no significant difference between seasons or interactions (Fig 5a; Table 2). N content was also
268 significantly higher in the non-burnt region than in the burnt region, and both regions had
269 significantly higher N content in wet season than in the dry season, with no significant
270 interactions between region and season (Fig 5b; Table 2). The C:N ratio was higher in the non-
271 burnt region during the dry season and the ratio was higher in the non-burnt region for wet
272 season, resulting in significant seasonal variations and significant interaction terms (Fig 5c;
273 Table 2).

274 All the rest of the macronutrients except N and S were greater in the burnt region than in the
275 non-burnt region and only S and K varied significantly between seasons with lower
276 concentrations for the burnt region in wet season (Fig 5d-h; Table 2). All macronutrients except
277 K exhibited significant interactions between seasons, as most nutrient concentrations did not
278 significantly differ between seasons in the non-burnt region (Fig 5d-h; Table 2).

279 **3.4. Changes with peat depth**

280 Electrical conductivity and pH were significantly higher in the burnt than in non-burnt region,
281 while the variations with depth and the interactions were not significant (Fig 6a,c; Table 3).
282 Gravimetric moisture in the non-burnt region was significantly higher than that of burnt region
283 in all depths (fig 6b,c; Table 3). The gravimetric moisture increased up to 20-30 cm and 30-40
284 cm for burnt and non-burnt region respectively (fig 6b; Table 3) and then showed a slight
285 decrease in deepest layers. Redox potential was significantly lower in the burnt region than in
286 the non-burnt region in the top 3 surface layers upto the depth of 30 cm (Fig 6d; Table 3).

287 Redox potential increased with depth in the burnt region while the redox potential in the surface
288 was slightly lowered in 10-20 cm and stayed at the same level in the deeper layers, resulting in
289 significant seasonal variations and interaction terms (Fig 6d; Table 3).

290 C content was significantly greater in the non-burnt region than in the burnt region with no
291 significant variations with depth nor interactions (Fig 7a; Table 3). N content and C:N ratio did
292 not significantly vary between regions or depth, with no significant interactions between region
293 and depth (fig 7b,c; Table 3).

294 All macronutrient concentrations except S and N were significantly greater in the burnt than in
295 the non-burnt region (Fig 7d-h; Table 3). Mg, P and Ca concentrations significantly decreased
296 with depth, which is more pronounced in the burnt region (Table 3). Mg and Ca exhibited
297 significant interactions between region and depth, as the concentration in the burnt region
298 decreased with depth whilst the concentrations in the non-burnt region stayed at the same level
299 throughout the depths. K and S did not show any significant variations with depth, nor any
300 significant interaction between region and depth (Fig 7f,g; Table 3).

301 **3.5. GHG emissions and environmental controls**

302 Backward stepwise multiple regression showed that CO₂ emissions were positively correlated
303 with pH and electrical conductivity and negatively correlated with P and K concentrations
304 (Table 4).

305 Similar regression for CH₄ emissions showed that the CH₄ emissions were negatively
306 correlated with pH and P concentrations, and positively correlated with Ca concentrations
307 (Table 4).

308 **4. Discussion**

309 Drainage for agriculture in peatlands were known to increase the CO₂ emissions with exposure
310 of peat to aerobic decomposition (Couwenberg et al., 2010; Dhandapani et al., 2019a; 2019c),
311 this study shows that the use of fire in management practices results in further increase in CO₂
312 emissions post fire (Fig 3a). There are several factors that may have played a part in this
313 increase in emissions in fire affected region, such as increased pH and increased concentrations
314 of macronutrients. It should be noted that the CO₂ emissions decreased and levelled off with
315 that of non-burnt region after few months in the wet season, while the CO₂ emissions from the
316 non-burnt region did not significantly vary between the seasons, suggesting that this artefact
317 was not just seasonal variations of CO₂ emissions, but rather a reverting of emissions along a
318 temporal gradient since the fire. However, while CO₂ emissions of pristine, secondary peat
319 forests and first generation oil palm showing no significant seasonal variations, seasonal
320 variation have been observed in second generation agricultural systems in the area (Dhandapani
321 et al., unpublished; Dhandapani et al., 2019a; 2019c; 2020).

322 The observed increase in CO₂ emissions post fire may be due to intense disturbance during fire
323 and subsequent changes in peat physicochemical properties such as increased pH, electrical
324 conductivity and nutrient concentrations (Fig 4c-e; Fig 5d-h), along with breakdown of
325 complex organic matter due to fire and new exposure of deeper peat layers to aerobic conditions
326 (Sazawa et al., 2018). However, this increase is short lived as the surface conditions are
327 stabilised over time and also possibly because of greater loss of newly acquired nutrients over
328 time (Fig 4c; Fig 5d-h). This trend of initial increase and consequent decline in CO₂ emissions
329 post fire was previously observed by Astiani et al. (2018) in West Kalimantan peatlands. It
330 should be noted that in Astiani et al. (2018)'s study, the peak in emission was observed in the
331 9th month post fire, and emissions stabilising at the 11th month post fire. The stabilisation of
332 emissions in 5th month post fire (Fig 3a) in our study may be due to a lower intensity localised
333 fire in our agricultural study site, compared to high intensity fire burning much greater forest

334 biomass in Astiani et al. (2018)'s study that would likely result in greater disturbance, greater
335 nutrient addition from resultant ash, and thus would require greater time to recover and
336 stabilise.

337 CH₄ emissions were also significantly higher in the burnt region than in the non-burnt region
338 and stayed at similar higher levels even into the wet-season, five months after the fire event
339 (Fig 3b). It was previously observed that fire would make the peat dry and effect lower CH₄
340 emissions from the burnt areas because of the lack of moisture (Davidson et al., 2019), but this
341 study has shown that at similar moisture levels, burnt peat has a potential for higher CH₄
342 emissions than the non-burnt peat. This higher CH₄ emissions in the fire-affected region may
343 be due to the availability of more labile C due to the denaturation of peat by fire and resultant
344 heat (Sazawa et al., 2018). Once the moisture level for anaerobic conditions are met, CH₄
345 emissions are dependent on the availability of labile C in peatlands (Couwenberg, 2009). Fire
346 in peatlands are also known to reduce methanotrophic activity (Danilova et al., 2015), further
347 helping the increase in CH₄ concentration in the burnt areas. However, it should be noted that
348 CH₄ emissions are very low overall, with mean values under 0.1 mg m⁻²hr⁻¹, which is in line
349 with the lower near zero fluxes observed in different oil palm agricultural plantations in the
350 region. Southeast Asian peatlands are naturally low methane emitting landscapes (Dhandapani
351 et al., 2019c; Hatano et al., 2016) compared to peatlands in other regions such as neotropics
352 (Wright et al., 2013), boreal (Kettunen et al., 1996) or temperate regions (Abdalla et al., 2016).
353 This validates our first hypothesis that greenhouse gas emissions are significantly increased in
354 the fire affected burnt region in relation to the non-burnt region.

355 The surface peat temperature and moisture did not significantly vary between the burnt and
356 non-burnt region, possibly because both regions are open with no ground cover to provide
357 shade and cool down surface peat (Dhandapani et al., 2019a). The significant seasonal changes
358 in both these properties were as expected (Dhandapani et al., 2019a), with higher moisture and

359 lower temperature in both regions during wet season, because of increased rainfall in wet
360 season compared to the dry season (Global Environmental Centre, 2014). Similarly, higher
361 moisture with increasing depth were also as expected, as the top layers in these agricultural
362 peatlands are actively drained, and the surface layers are exposed to the sun, further facilitating
363 increased temperature and evaporation in the surface layers. Most other physico-chemical
364 properties and nutrient concentrations showed that the surface layers such as 0-10cm and 10-
365 20cm showed significant difference between burnt and non-burnt regions before gradually
366 narrowing in on the difference, to the same level in deeper layers starting from 20-30cm layer.

367 Redox potential is an important soil property which shows the electron exchange capacity for
368 reduction and oxidation reactions that have great impact on nutrient availability and dynamics
369 in soil by changing their electric charges (Niedermeier and Robinson, 2007; Søndergaard,
370 2009). The reduction in redox potential in the burnt region shows the possible lack of oxidants
371 in the burnt area (Fiedler, 2000). This is likely to be caused by fire, as this difference is observed
372 only in the surface layers, with deeper layers having same level of redox potential in the deepest
373 layers for both burnt and non-burnt regions. Fire were known to accumulate new particulate C
374 forms in mineral soil humus layers that are highly resistant to redox reactions (González-Pérez
375 et al., 2004), a similar trend is observed here in peat soil where fire affected surface layers have
376 low redox potential.

377 Electrical conductivity and pH increased one month after fire, however electrical conductivity
378 levelled off with that of non-burnt region during wet season, five months after the fire activity.
379 The increase in pH after fire has been widely reported in different soil environments (Chungu
380 et al., 2019; Heydari et al., 2017; Kennard and Gholz, 2001; Scharenbroch et al., 2012; Zaccone
381 et al., 2014), that also holds true for highly acidic tropical peatlands where pH was almost twice
382 as high in the burnt region compared to non-burnt region (Fig 4e). This increase may be directly
383 related to the addition of ash from burning (Bang-Andreasen et al., 2017; Zaccone et al., 2014),

384 as shown by the significant increase only in the top 20 cm surface layers. Electrical conductivity
385 exhibited very similar trend of change with depth as pH, possibly because of the same influence
386 of ash addition in the top layers (Bang-Andreasen et al., 2017; Zacccone et al., 2014). Electrical
387 conductivity indicates the salt content and is considered a rough estimate of soil nutrients that
388 influence important soil processes such as GHG emissions (Visconti and De Paz, 2016). The
389 significant reduction of electrical conductivity and levelling off during the wet season possibly
390 shows that many of the nutrients that are made available after fire (Beest et al., 2019), were
391 leached off in the proceeding few months (Beliveau et al., 2015), aided by heavy rainfall during
392 the wet season (Global Environmental Centre, 2014). This is also evident in reduced
393 concentration of all macronutrients in the wet season (Fig 5).

394 All of the macronutrients except N increased in the burnt region one month after fire and
395 slightly decreased over time, yet remained higher than the concentrations in the non-burnt
396 region throughout the seasons, except for S that had lower concentration in the burnt region
397 than in the non-burnt region in wet season. The nutrients that are strongly basic and form
398 important salts such as Mg and Ca (Visconti and De Paz, 2016), explicablely exhibited the same
399 trend of change with depth as pH and electrical conductivity. P is also widely reported to
400 increase in concentration in burnt peat and soil, as fire causes the release of P from organic
401 materials (Beest et al., 2019; Wang et al., 2015). The increase in Ca, Mg and K concentrations
402 in the burnt region can also be attributed to high content of these nutrient in wood ash that
403 accumulated in the region as a result of burning previous generation oil palm stems (Kennard
404 and Gholz, 2001). These nutrients are also found to be highly susceptible to leaching (Kennard
405 and Gholz, 2001), explaining the lower concentrations in wet season sampling, few months
406 after the fire event. C content was reduced to half in the surface layers, because of the C lost
407 on the burning of organic peat on the surface (Wiggins et al., 2018). The peat cores were
408 selectively collected in areas where the peat is not completely burned off, thus the difference

409 in C content between the surface layers in the cores and high replicate surface samples
410 associated with GHG emission measurements (Fig 5a and 7a). However, it should be noted that
411 lower number of replicate peat cores may also be a factor limiting the usability of these results.
412 This validates our second hypothesis that peat physico-chemical properties and nutrient
413 concentrations were significantly affected by fire, and most of these properties showed
414 significant variations with depth.

415 These changes in physico-chemical properties were significantly correlated with CO₂ and CH₄
416 emissions, most visibly CO₂ emissions exhibited the same pattern of change as electrical
417 conductivity and pH of surface peat. Multiple regression analyses showed that peat physico-
418 chemical properties and total nutrient concentrations were strong predictors of change in CO₂
419 and CH₄ emissions (Table 4). CO₂ emissions from tropical peat of different land uses are
420 known to exhibit positive correlations with pH (Dhandapani et al., 2019a) and our results show
421 that the fire event resulting in sudden and steep increase in pH does not override this
422 relationship. Inversely, CH₄ was negatively related with pH, this negative correlation has also
423 been observed in northern peatlands (Weslien et al., 2009). Acidic condition is a prerequisite
424 to functionalise precursors such as H₂/CO₂, acetic acid, formic acid, methanol and
425 methylamine, for CH₄ production (Qing-Yu et al., 2019), thus explaining this functional
426 relationship between pH and CH₄ emissions in the field.

427 The positive relationship between electrical conductivity or salinity and CO₂ emissions has
428 been observed in other ecosystems (Capooci et al., 2019), however this is the first time such
429 relationship is explored and reported in tropical peatlands. Though very high or excess salinity
430 can be toxic and can negatively impact microbial activity and reduce CO₂ production (Setia et
431 al., 2013; Yan et al., 2015), tropical peatlands are naturally nutrient poor (Sjögersten et al.,
432 2011) and this increase in electrical conductivity due to fire boosted CO₂ production (Table 4).
433 P concentrations were negatively related with both CO₂ and CH₄ emissions. Additionally, CO₂

434 emissions were negatively related with K, and CH₄ emissions were positively related with Ca.
435 Nutrient dynamics (with the exception of N) and their interactions with other peat processes
436 such as C cycling were not well documented in agricultural tropical peatlands. Addition of
437 manures containing these macronutrients had different effects on GHG emissions in different
438 soil types (Ren et al., 2017), and their cause and effect pattern in tropical peatlands need more
439 *in situ* research coupled with controlled experiments to understand the interactions, especially
440 considering that they exhibit significant functional correlations with GHG emissions. However,
441 P is known to support methane oxidation reactions in soil (Veraart et al., 2015) and similarly
442 limited addition of Ca compounds were known to support methane production in varied set ups
443 (Khor et al., 2015; Lar et al., 2010), suggesting similar functional relationship in tropical
444 peatlands. In line with this, Murakami et al. (2005) also found CH₄ potential increased with
445 liming (CaCO₃ addition) in agricultural tropical peatlands. This validates our third hypothesis
446 that the changes in soil physico-chemical properties and nutrient concentrations correlate with
447 changes in GHG emissions.

448 As oil palm is a high biomass crop, there is a potential alternative for the biomass removed
449 from previous generation to be used for biochar production, which then could be used to enrich,
450 and return C back to agricultural peatlands, potentially retaining some of the benefits of slash-
451 and-burn practice such as increased pH and nutrient concentrations (Bista et al., 2019; Kong et
452 al., 2014; Liew et al., 2018). Such biochar can be produced using methods ranging from
453 constructing a conical soil pit in the ground, and simple conical or retort kilns to modern
454 microwave pyrolysis equipment (Liew et al., 2018; Arafat Hossain et al., 2017). However
455 socio-economic feasibility, and practicality of such alternatives for biomass management are
456 yet to be explored and researched. There is a need for increased research in understanding the
457 holistic environmental and socio-economic impacts of slash-and-burn practice in tropical
458 peatlands, and other alternative practices for land clearing and biomass management between

459 oil palm generations. This study is a step towards such understanding of the impacts of different
460 practices. Further research addressing these issues will help us make informed suggestions on
461 land use policies and management practices, for rapidly increasing area of agricultural tropical
462 peatlands.

463 **5. Conclusion**

464 The use of fire for slash-and-burn management practice severely alters peat physico-chemical
465 properties and increases post-fire greenhouse gas emissions and peat nutrient concentrations.
466 While some of the properties such as CO₂ emissions, and electrical conductivity reverted back
467 to normal level after few months, other properties such as CH₄ emissions, pH and nutrient
468 concentrations remained high in the burnt region, though nutrient concentrations were
469 significantly reduced over time, possibly due to leaching. This study has shown very high loss
470 of surface peat C content due to fire, which is irreversible. This has important implications if
471 such a loss occurred in all agricultural peatlands in the region through slash-and-burn activity.
472 The practice of intercropping was previously known to prolong defining peat properties and
473 ameliorate environmental impacts of oil palm agriculture in peat, but the current results show
474 that use of fire as a management practice overrides any such benefits from intercropping,
475 because of high C loss. The peat surface physico-chemical properties and nutrient
476 concentrations also exhibited functional correlations with GHG emissions, providing new
477 insights into complex interactions between different biogeochemical processes in tropical
478 peatlands. The results also show that peat layers from surface to 20 cm depth were most
479 affected by this particular slash-and-burn activity in oil palm agriculture. However, the
480 intensity of fire can vary widely between different oil palm management and needs further
481 research to fully understand the long term and regional impacts of such slash-and-burn activity
482 in tropical peatlands.

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490

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719

720 **Figure Captions**

721 **Figure 1:** Location of the study site.

722 **Figure 2:** Site pictures - a) picture of the study site during smouldering fire July 2018 b) picture
723 of the site during the dry season measurements, one month after the fire event, in August 2018
724 c) picture of the site during the wet season measurements, five months after the fire event, in
725 December 2018.

726 **Figure 3:** Effect of burnt and non-burnt region, and season upon a) CO₂ emissions b) CH₄
727 emissions, during dry and wet season. Bars denote mean values (for dry season n=20 each
728 region; for wet season n=10 each region). Whiskers denote standard errors.

729 **Figure 4:** Effect of burnt and non-burnt region, and season upon a) Peat surface temperature
730 b) gravimetric moisture c) electrical conductivity d) redox potential e) pH during dry and wet
731 season. Bars denote mean values (for dry season n=20 each region; for wet season n=10 each
732 region). Whiskers denote standard errors.

733 **Figure 5:** Effect of burnt and non-burnt region, and season upon peat a) C content b) N content
734 c) C:N ratio d) Mg concentrations e) P concentrations f) S concentrations f) K concentrations
735 g) Ca concentrations during dry and wet season. Bars denote mean values (for dry season n=20
736 each region; for wet season n=10 each region). Whiskers denote standard errors.

737 **Figure 6:** Effect of burnt and non-burnt region, and depth upon a) pH b) gravimetric moisture
738 c) electrical conductivity d) redox potential during dry season. Points denote mean values (n=3
739 for each region). Whiskers denote standard errors.

740 **Figure 7:** Effect of burnt and non-burnt region, and depth upon peat a) C content b) N content
741 c) C:N ratio d) Mg concentrations e) P concentrations f) S concentrations f) K concentrations
742 g) Ca concentrations during dry season. Bars denote mean values (n=3 for each region).
743 Whiskers denote standard errors.

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