

Assessing the hazard of saltwater intrusion in coastal regions of the Red-Thai Binh River Basin, Vietnam

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Abstract

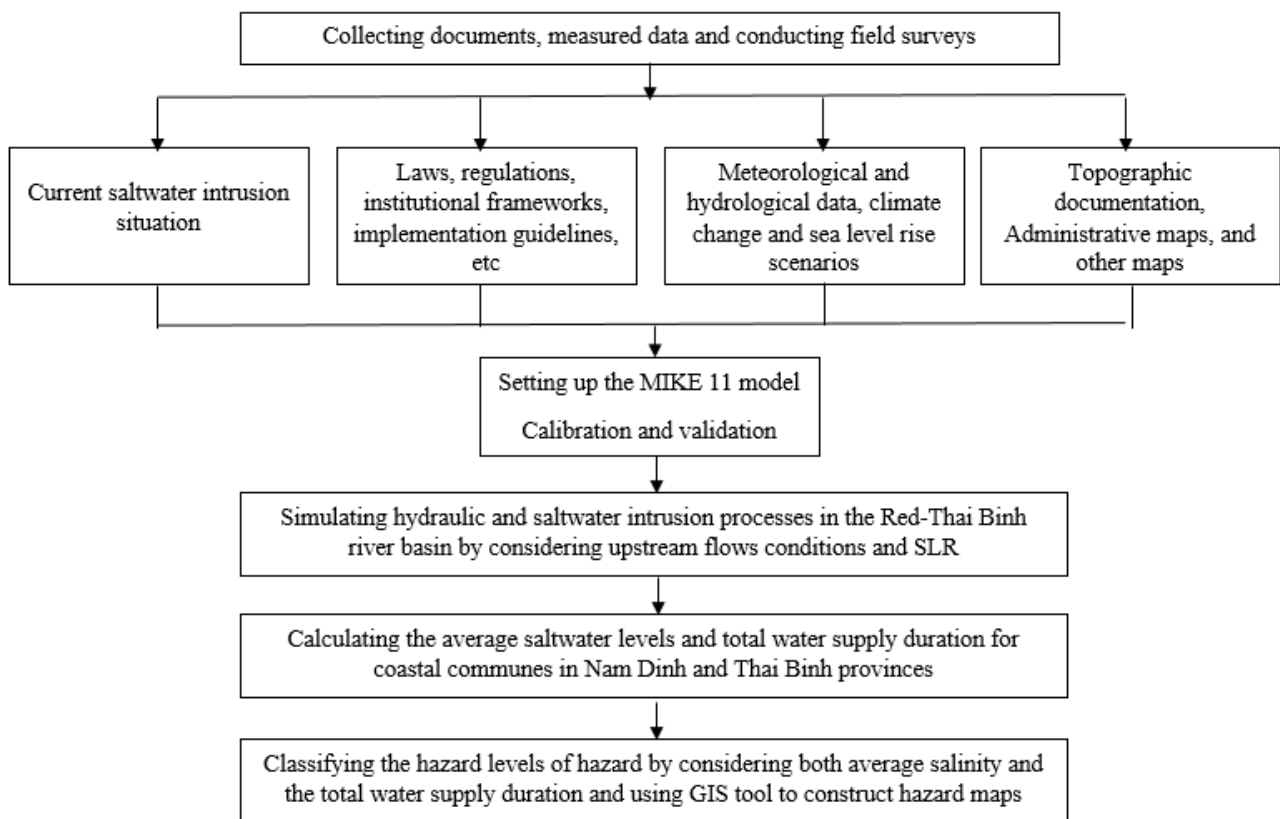
The Red–Thai Binh River Basin in Vietnam faces severe saltwater intrusion, exacerbated by climate change, upstream water use, and inadequate reservoir management. During the dry season, reduced upstream discharge leads to higher salinity at sluice gates, thereby limiting the duration of the water supply. This study aims to assess and classify saltwater intrusion hazards in the basin under varying MIKE11 model upstream-flow and sea-level-rise scenarios. To achieve this, a multi-step framework was applied. First, the model was used to simulate saltwater intrusion dynamics, with calibration results showing strong performance ($NSE = 0.72\text{--}0.92$) and salinity deviations of less than 2.18‰. Second, simulated outputs were used to calculate average salinity at the commune level, enabling localised analysis. Third, a hazard classification system was developed by integrating average salinity levels with water-supply duration at sluice gates, thereby capturing both water-quality and availability constraints. The approach was applied to 15 scenarios across 243 coastal communes in the Nam Dinh and Thai Binh provinces, accounting for the combined effects of upstream discharge conditions and projected sea-level rise. The results reveal distinct spatial variations in saltwater intrusion hazards. In Nam Dinh Province, approximately 40–50% of communes were classified as having moderate-to-high hazard levels, indicating significant vulnerability. In contrast, in Thai Binh Province, most communes were classified below moderate hazard levels, suggesting comparatively lower risk. Overall, the findings identify priority areas where saltwater intrusion poses the greatest challenges and provide a basis for targeted adaptive management and improved water resource planning.

Keywords: saltwater intrusion, average salinity, water supply duration, hazard level mapping, Red-Thai Binh River Basin.

Highlight:

- Saltwater intrusion is increasingly intense and frequent in coastal regions due to climate change, rising sea levels, and upstream water resource use.
- This study proposed a method to evaluate and classify hazard levels based on average salinity and water supply duration.
- The study evaluated 15 scenarios accounting for the combined effects of upstream flow conditions and projected sea-level rise.
- The results reveal distinct spatial variations in saltwater intrusion hazards in Nam Dinh and Thai Binh Province, Vietnam.
- The findings highlight areas where saltwater intrusion poses the most significant challenges and where adaptive management measures should be prioritised.

Graphical Abstract



1. Introduction

Saltwater intrusion is a complex process occurring in coastal deltas and estuaries (Bear *et al.*, 1999), involving the replacement or mixing of freshwater with saltwater (Hebert, 2019). This phenomenon affects both surface water and groundwater systems (Tu *et al.*, 2021), and typically develops gradually over long periods. The interaction between hydrodynamic and human activities governs saltwater intrusion. Major controlling factors include upstream freshwater withdrawal, regulation of coastal flow, and Sea Level Rise (SLR) driven by climate change. (Liu *et al.*, 2019; Wang *et al.*, 2019; Zhang *et al.*, 2010). Tidal dynamics, storm surges, wind, wave action, estuarine geomorphology, and topography further influence the extent and variability of saltwater intrusion (Gong *et al.*, 2013; Liu *et al.*, 2017; Shen *et al.*, 2018; Perales-Valdivia *et al.*, 2018; Tian, 2019).

Recent advances in computational methods have enhanced the use of mathematical and machine learning models for analysing and predicting water quality dynamics (Maruthai *et al.*, 2025; Venkatraman *et al.*, 2025). Salinity is a key indicator for evaluating coastal hydrology and environmental conditions, particularly in the context of saltwater intrusion. To simulate saltwater intrusion processes, hydraulic interaction models of varying dimensionality have been developed, including one-dimensional (1D) (Lafta, 2022; Zhang *et al.*, 2013), two-dimensional (2D) (Haddout *et al.*, 2018; Lee *et al.*, 2016) and three-dimensional (3D) (Jeong *et al.*, 2010; Veerapaga *et al.*, 2019; Chen *et al.*, 2016) models. However, numerous studies have shown that the accuracy and reliability of complex models, such as 2D or 3D simulations, depend heavily on the availability of comprehensive observational datasets and, in many cases, require additional field surveys to ensure robust results (Kaddi *et al.*, 2022). Consequently, implementing such models in data-scarce regions remains a significant challenge (Lindenschmidt, 2006).

While higher-dimensional models offer detailed representations of hydrodynamic and solute transport processes, their computational demands often limit their practical application. To address these

27 challenges, 1D models are frequently employed as efficient alternatives, incorporating simplifying
28 assumptions to balance model complexity with predictive accuracy (Doulgeris *et al.*, 2012; Ferdous,
29 2014; Haque *et al.*, 2021; Miller, 2023; Vuong Tai *et al.*, 2021). When implemented with high
30 temporal resolution, 1D models have demonstrated strong capability in reproducing the temporal
31 dynamics of saltwater intrusion (Gaiolini *et al.*, 2023).

32

33 Simulation models have been widely employed to evaluate how flow regimes and tidal dynamics
34 influence the spatial extent of saltwater intrusion (Becker *et al.*, 2010). In coastal estuaries, these
35 processes are becoming increasingly complex due to the combined pressures of climate change and
36 SLR (Bricheno *et al.*, 2021; Ferguson & Gleeson, 2012; Hong *et al.*, 2020; Oude Essink, 2001). This
37 growing complexity poses substantial socio-economic risks for vulnerable regions (Khong *et al.*,
38 2018; Meisner, 2007). A broad body of research further demonstrates that SLR plays a critical role
39 in reshaping salinity regimes and contributing to the deterioration of coastal water quality (Erdbrügger
40 *et al.*, 2023; Rice *et al.*, 2012; Thai & Thuc, 2011; Thanh *et al.*, 2023; Yuan *et al.*, 2015).

41

42 Vietnam has a coastline of over 3,200 km and is one of the countries most affected by natural
43 disasters, including saltwater intrusion. In the context of climate change, saltwater intrusion is
44 becoming more severe in the country (UNDP Vietnam, 2016). Recent studies have examined
45 saltwater intrusion in Vietnam's large river basins, focusing on the impact of freshwater resources
46 from upstream and SLR, notably in the Red-Thai Binh (RTB) River Basin (Hien *et al.*, 2018; Hien *et*
47 *al.*, 2023). The RTB is an international river basin that flows through three countries: Vietnam, China,
48 and Laos. It encompasses an area of approximately 169,000 km² (**Figure 1**). In Vietnam, the basin
49 covers 86,680 km², accounting for 51.3% of the total area, with the delta region spanning
50 approximately 17,000 km².



Figure 1. The RTB River basin in Vietnam (green area), and the boundary of Nam Dinh and Thai Binh provinces (red line).

The RTB River system discharges into the Gulf of Tonkin, located in the northwestern part of the South China Sea. The gulf is characterised by a diurnal tidal regime that produces some of the largest tidal amplitudes in Vietnam, reaching up to 4 metres. These tidal fluctuations have a strong influence on flow dynamics and saltwater intrusion processes across the eight estuaries of the RTB, particularly during the dry season from October to May, when river discharge is significantly reduced (Devienne, 2006).

The terrain of the RTB Delta is characterised by a gentle slope from northwest to southeast, with ground elevations ranging approximately from 0.4 to 9 meters above sea level (NARBO, 2006). This low-lying topography plays a crucial role in hydrodynamics, drainage patterns, and the delta's susceptibility to saltwater intrusion and flooding.

67 To supply water for agricultural activities, the rivers in the RTB Delta have about 432 sluices, whose
68 operation depends on water demand, river level, and salinity. The system of sluices and canals creates
69 a high-density water-supply network, ensuring water for communes and agricultural areas.

70

71 The salinity monitoring network in the RTB Delta comprises 18 stations, some of which have been
72 active since 1963. However, these observations are not uniform; while some stations have recorded
73 data for 15 to 20 years, others have only monitored salinity for 2 to 3 years. The data indicate that
74 river salinity levels change seasonally. Salinity levels are low during the rainy season and high during
75 the dry season, primarily influenced by the amount of freshwater released from upstream and the
76 strength of tidal movements. The highest salinity levels are observed in January.

77

78 Many studies on saltwater intrusion in the RTB Delta have employed 1D models. A limitation of
79 these studies is their focus on salinity processes in the main rivers only, while neglecting saltwater
80 intrusion in the canal network. The findings indicate that saltwater can penetrate as far as 30 to 40
81 kilometres inland, with salinity levels exceeding 1‰ in some areas. Moreover, although salinity
82 boundary maps have been created, they have been generated through linear interpolation of salinity
83 levels along the main rivers, without considering the salinity in the internal canal network (Thai &
84 Thuc, 2011). This methodology has its limitations and is open to varied interpretations. Additionally,
85 limited research has examined the total duration of water supply from sluices, constraining a
86 comprehensive assessment of saltwater intrusion impacts.

87

88 Previous investigations in the RTB Delta have been limited by the complexity of its river and canal
89 networks and by the restricted spatial coverage of salinity monitoring stations. Furthermore, available
90 salinity records are often discontinuous and lack temporal synchronisation among stations, hindering
91 data interpretation and the construction of comprehensive salinity boundary maps. In most cases,
92 salinity is measured at a single point along the river cross-section, which inadequately represents both

horizontal and vertical variations in salinity distribution. To overcome these limitations, there is a need for more integrated and adaptive research approaches. For example, when evaluating the hazards posed by salinity, it is crucial to consider both average salinity concentrations and the duration of water availability. Hence, this study proposes a novel methodological framework for estimating average salinity levels in irrigated rice fields in deltaic and coastal regions. This framework is beneficial in situations where cross-sectional measurements are unavailable, and hydrodynamic models cannot effectively simulate salinity dynamics.

This study aims to develop a hazard map for saltwater intrusion based on average salinity levels and the total duration of water supply. The specific objectives are as follows: (1) to simulate the hydraulic and saltwater intrusion processes in the RTB River Basin during the peak dry season (January-March), considering the influence of upstream flow and SLR; (2) to calculate the average salinity levels and total water supply duration for 243 coastal communes in Nam Dinh and Thai Binh provinces that are most affected by saltwater intrusion in the RTB Delta; and (3) to create saltwater intrusion hazard maps that illustrate the impact levels on the communes under different SLR scenarios.

2. Methodology

Figure 2 illustrates the methodological approach used in this study.

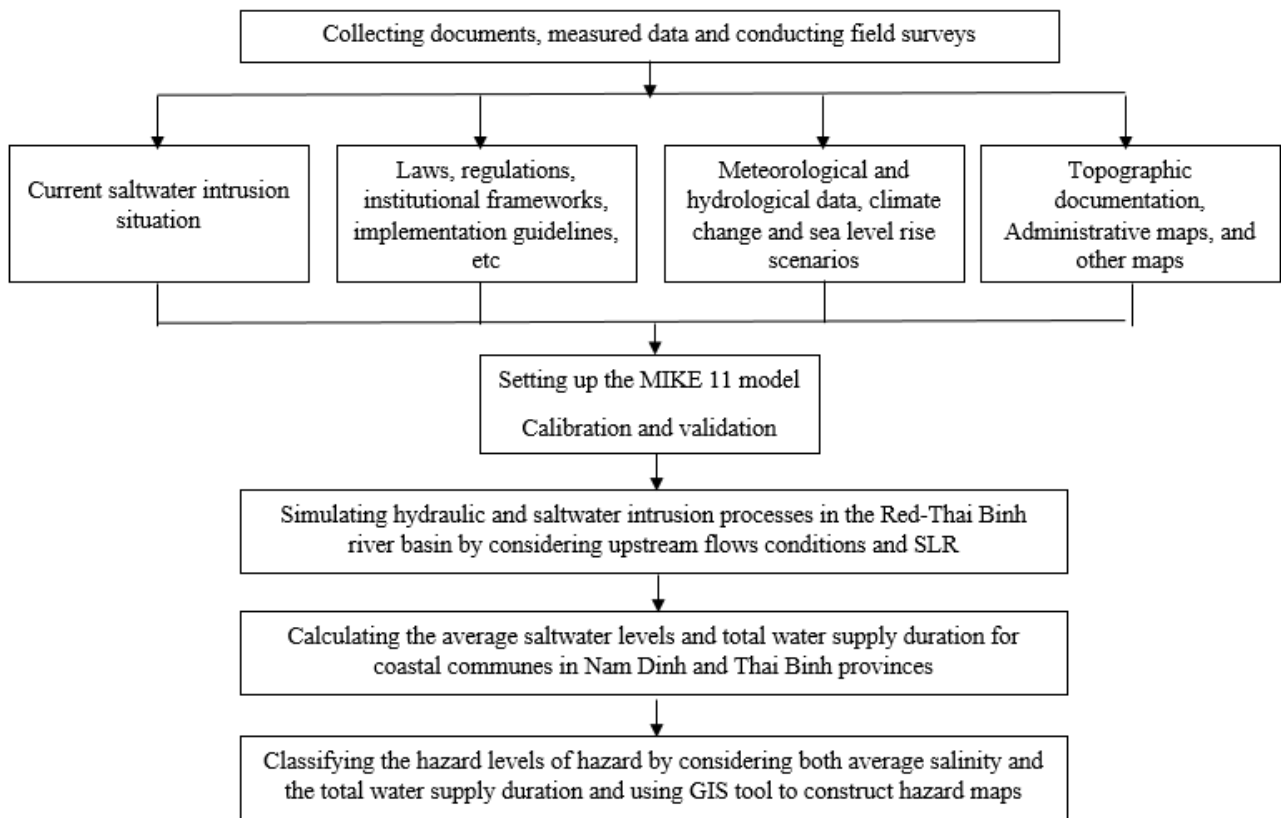


Figure 2. Flowchart of the research methodology.

2.1 Setting up the simulation model

1-D models utilised to simulate saltwater intrusion in river estuaries operate under the basic assumption that flow and salinity characteristics are uniform across the river's cross-section. Although this assumption may not be valid, the simulation results remain acceptable even with limited input data. Among the various 1-D models used to address this issue, the MIKE 11 model is particularly prominent. MIKE 11 simulates flow, sediment transport, and water quality in estuaries, rivers, irrigation systems, and canals. Additionally, it is recognised for its user-friendly dynamic modelling capabilities. Consequently, this study employed the MIKE 11 model to simulate saltwater intrusion in the RTB River Basin.

The model had nine upstream and eight downstream model boundaries (**Table 1**), with the latter located 15-25 km from the river mouth. The salinity at the upstream boundaries was set at 0‰, while

127 the salinity at the offshore locations was 32‰, based on data recorded from 2000 to 2016 at Hon Dau
128 station. For calibration and validation, this study used water-level, discharge, and salinity
129 measurements from several downstream stations. **Figure 3** illustrates the river network, boundaries,
130 and monitoring stations.

131 **Table 1.** List of upstream and downstream boundaries

No.	River	Station	Location	
		name	LAT (°)	LONG (°)
Upstream boundaries				
1	Da	Hoa Binh	20.82	105.33
2	Thao	Yen Bai	21.70	104.86
3	Chay	Thac Ba	21.75	105.05
4	Lo	Ham Yen	22.05	105.09
5	Gam	Chiem Hoa	22.09	105.28
6	Cau	Gia Bay	21.60	105.83
7	Luc Nam	Chu	21.37	106.61
8	Thuong	Cau Son	21.45	106.32
9	Day	Ba Tha	19.92	105.58
Downstream boundaries				
1	Day		19.82	106.03
2	Ninh Co		19.89	106.25
3	Red		20.16	106.64
4	Tra Ly		20.44	106.70
5	Thai Binh		20.57	106.72
6	Van Uc		20.62	106.78
7	Lach Tray		20.76	106.84
8	Cam		20.77	106.88

134 The data used in this study reflect conditions from 2019 (**Table 2**). The modelled cross-section spans
 135 approximately 10 km in width and ranges in depth from 20 to 30 m, accurately representing the coastal
 136 topography. Simulations were run from January to March, during the peak of the dry season,
 137 characterised by minimal rainfall and lateral inflows. Initial conditions for flow discharge, water
 138 level, and salinity were based on field measurements to simulate intrusion processes and to calculate
 139 average salinity and water-supply duration, which were used to create hazard maps for affected
 140 coastal areas.

141

142 **Table 2.** The number of cross-sections on the simulation network's rivers

No.	River	Number of cross-sections	No.	River	Number of cross-sections
1	Cam	54	16	Lo	68
2	Cau	68	17	Luc Nam	23
3	Chay	50	18	Luoc	120
4	Da Bac	11	19	Mia	60
5	Dao	88	20	Moi	62
6	Day	78	21	Ninh Co	38
7	Duong	42	22	Pho Day	30
8	Gam	32	23	Quan Lieu	38
9	Hoa	48	24	Tam Bac	24
10	Hoang Long	18	25	Thai Binh	105
11	Red	63	26	Thao	78
12	Kinh Mon	80	27	Thuong	63
13	Kinh Thay	65	28	Tich	33
14	Lach Tray	75	29	Tra Ly	96
15	Lai Vu	57	30	Van Uc	53

143 The hydraulic (HD) and advection-dispersion (AD) modules in the MIKE 11 model were used
 144 together to simulate the salinity process. To calibrate and validate the parameter set, we followed a
 145 two-step trial-and-error method: (1) calibrating the HD parameter, specifically the Manning
 146 roughness coefficient (n), the Manning roughness coefficient is not fixed for the entire network but
 147 is zoned based on the morphological characteristics of each river section, and (2) calibrating the AD
 148 parameters, which include the dispersion factor (a), dispersion exponent (b), and dispersion
 149 coefficient (D), D is a function of velocity (u) as ($D = a \cdot |u|^b$). We evaluated the model's performance
 150 using the Nash-Sutcliffe Efficiency (NSE) coefficient and the absolute error (ΔS). The NSE and ΔS
 151 are calculated using **Equation 1** and **Equation 2**, respectively:

$$152 \quad NSE = \frac{\sum_{i=1}^n (x_i - \bar{x})^2 - \sum_{i=1}^n (y_i - x_i)^2}{\sum_{i=1}^n (x_i - \bar{x})^2} \quad (1)$$

$$153 \quad \Delta S = |X_{\max} - Y_{\max}| \quad (2)$$

154 where x_i and y_i are the measured and simulated values at each time i , $\bar{x}$ is the mean value of the
 155 measured time-series data; $X_{\max}$ and $Y_{\max}$ are the maximum measured and simulated values.

156

157 2.2 Development of salinity intrusion and climate change scenarios

158 To assess the influence of upstream flow and tidal levels on salinity, this study considered multiple
 159 scenarios: (1) three upstream flow conditions corresponding to the dry seasons of 2003, 1995 and
 160 2015, which have return periods of approximately 20, 4 and 2 years, respectively, based on a
 161 frequency analysis of observed data from January to March for 1990-2020; and (2) tidal conditions
 162 represented by the 2011 tide level, used as an approximation of the average tide between 1990 and
 163 2020 (scenario group A). In addition, SLR scenarios were examined for 2030, 2040, and 2050, under
 164 the RCP 4.5 pathway, and for 2050 under the RCP 8.5 pathway (scenario groups B-E). Combining
 165 upstream and downstream boundary conditions, a total of 15 simulation scenarios were considered
 166 (**Table 3**).

167 **Table 3.** List of simulated scenarios

No.	Scenario	Year of upstream discharge conditions	Downstream water level conditions for the year
1	1A	2003	2011
2	1B		2030 RCP 4.5
3	1C		2040 RCP 4.5
4	1D		2050 RCP 4.5
5	1E		2050 RCP 8.5
6	2A	1995	2011
7	2B		2030 RCP 4.5
8	2C		2040 RCP 4.5
9	2D		2050 RCP 4.5
10	2E		2050 RCP 8.5
11	3A	2015	2011
12	3B		2030 RCP 4.5
13	3C		2040 RCP 4.5
14	3D		2050 RCP 4.5
15	3E		2050 RCP 8.5

168

169 **Figure 4** illustrates the discharge hydrograph, while **Figure 5** presents the average discharge and
170 water levels recorded at various stations during the simulation period.

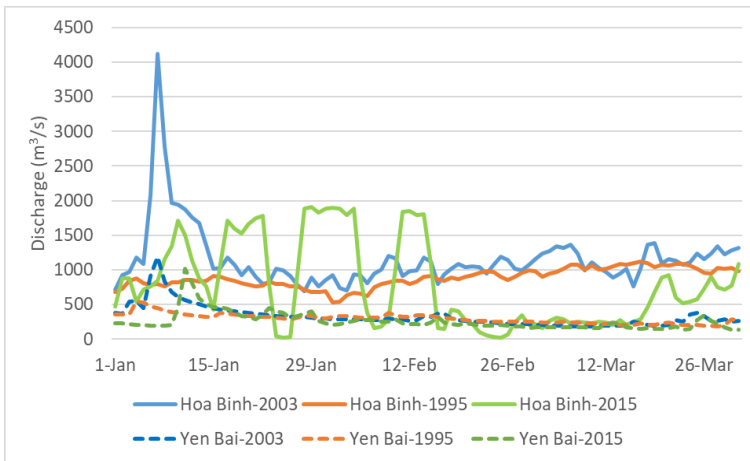


Figure 4. The discharge hydrograph from January 1 to March 31 of various years at selected stations.

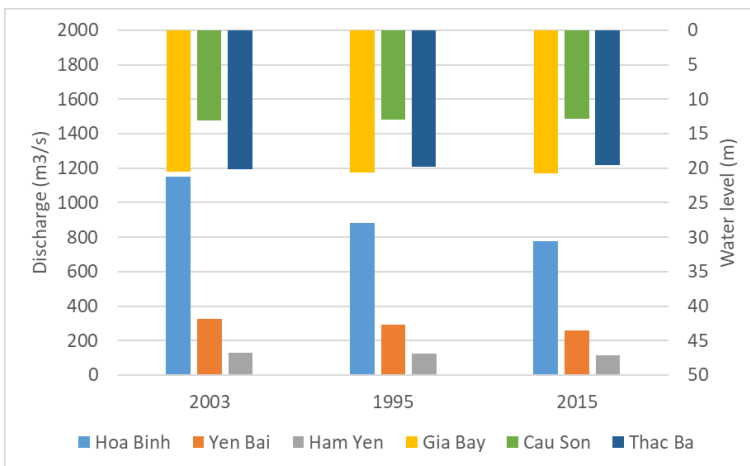


Figure 5. The average discharge and water level from January 1 to March 31 of various years at selected stations.

According to the national report on climate change and SLR scenarios released by the Ministry of Natural Resources and Environment of Vietnam (MONRE, 2021), sea levels are projected to rise by 12 cm, 17 cm, and 22 cm under the RCP4.5 scenario by the years 2030, 2040, and 2050, respectively. Additionally, a rise of 26 cm is expected by 2050 under the RCP8.5 scenario. This study calculated tide-level time series for sea level rise scenarios by averaging tide levels from 1990 to 2020 (using the 2011 water level value) and adding the projected increases for each corresponding year. We did not consider SLR in 2030 and 2040 under the RCP 8.5 scenario because the sea level difference between RCP 4.5 and RCP 8.5 in those decades is minimal. The time-series data of tide levels at the

simulation network's downstream boundaries can be extracted from MIKE 21's Global Tide Model for 2011 (DHI, 2014). The data were adjusted to account for the time zone and altitude of the Vietnamese mainland (Figure 6).

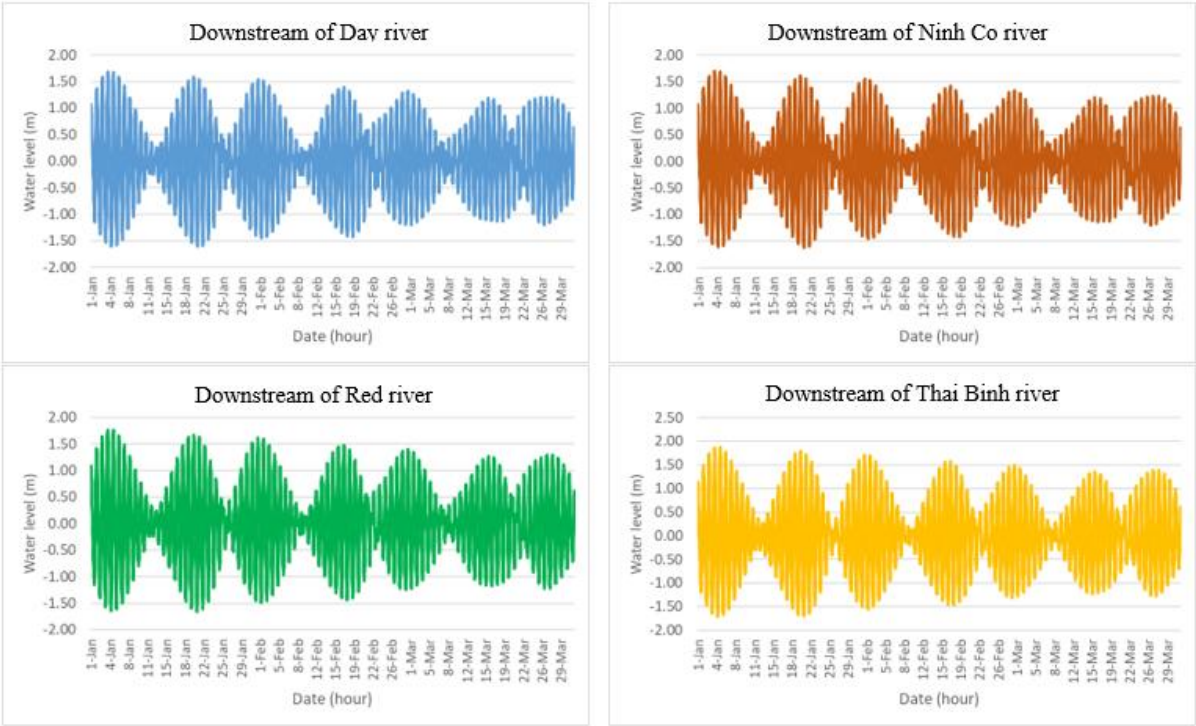


Figure 6. The tide level graph in 2011, as extracted from the MIKE model for different rivers

2.3 Assessing water supply duration and average salinity

The MIKE 11 model provides time-series data on water levels and salinity at specific locations and sluice gates along rivers for a given scenario. The total duration of water supply from each sluice across different scenarios k is determined based on two conditions: (1) the salinity in the main river for scenario (S_k) must be below the permitted level (S_p) for water to be supplied to the field, and (2) the water level in the river for scenario (Z_k) must be higher than the sluice bed elevation (Z_s). Thus, the total duration of the water supply for each sluice can be determined using this approach.

The spatial distribution of salinity in the study area is based on the following assumptions: (1) only saltwater intrusion through surface water systems is considered, excluding any intrusion into

203 groundwater layers; (2) the salinity concentration within a commune is calculated based on the
 204 salinity levels measured at sluices, and (3) the dilution of water in canals and fields is not considered
 205 because the flows and salinity in those areas were not observed.

206

207 Based on the assumptions outlined above, as well as an analysis of salinity diffusion principles,
 208 mathematical logic, and the study area's conditions, this research proposes **Equation 3** to calculate
 209 the average salinity for a commune under various scenarios. The advantage of this formula is that it
 210 calculates the average salinity based on the ratio of the total water supply from each sluice. This
 211 method simplifies calculations and ensures reliable results.

$$212 \quad S_{i,k} = \frac{\sum_{j=1}^n (V_{i,k,j} \times S_{i,k,j})}{\sum_{j=1}^n V_{i,k,j}} \quad (3)$$

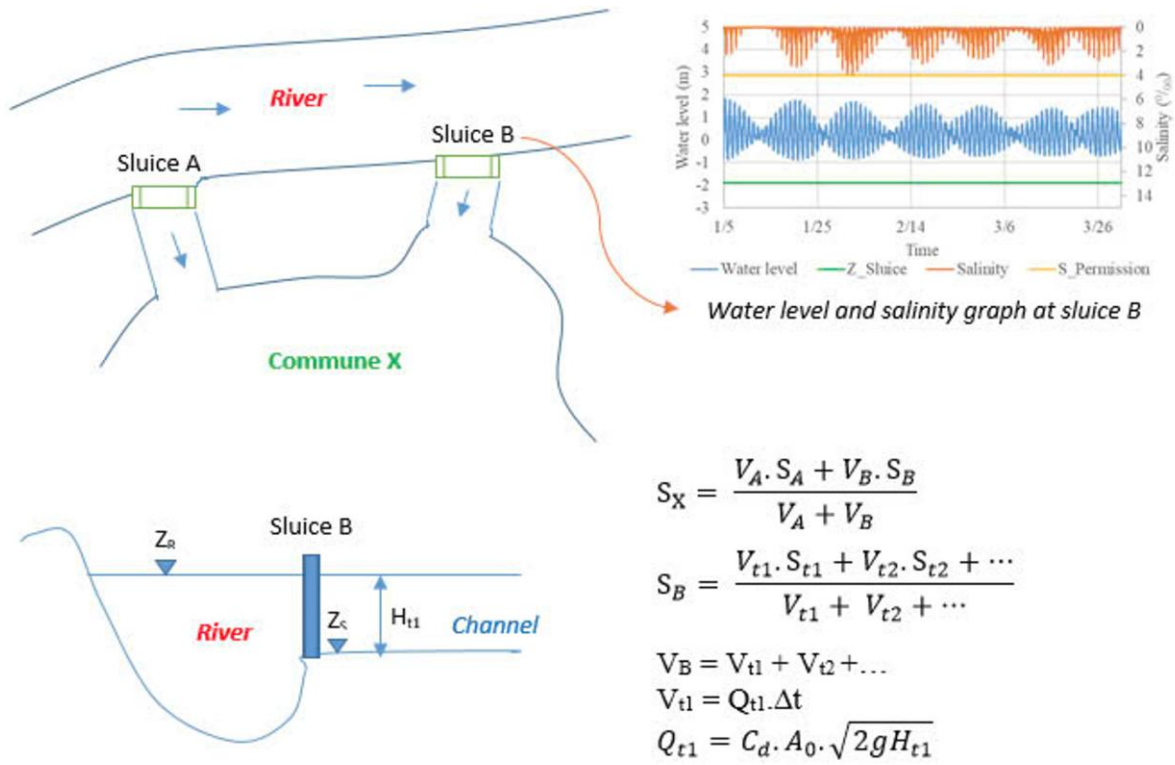
213 where $S_{i,k}$ is average salinity at a commune (i) corresponding to scenario (k), i ranges from 1 to 243
 214 (total commune in the study area is 243), $V_{i,k,j}$ is total water supply (V_j) of the sluice (j) for a commune
 215 (i) corresponding to scenario (k), j ranges from 1 to n , n is the total number of sluice supply water for
 216 a commune (i), the total water supply of each sluice is determined based on its size (exploited
 217 discharge) and the supply duration corresponding to different scenarios. $S_{i,k,j}$ is the average salinity at
 218 the sluice (j) corresponding to scenario (k) supply water for a commune (i), see **Equation 4**.

$$219 \quad S_{i,k,j} = \frac{\sum_{t=1}^m (V_{i,k,j,t} \times S_{i,k,j,t})}{\sum_{t=1}^m V_{i,k,j,t}} \quad (4)$$

220 where $S_{i,k,j,t}$ is salinity and $V_{i,k,j,t}$ is total water supply at the sluice (j) corresponding to scenario (k),
 221 supply water for a commune (i) has a salinity below the threshold in each supply duration (t), m
 222 denotes the total supply duration.

223

224 This example illustrates how to calculate the average salinity of commune X using **Equations 3** and
 225 **4**. The calculation involves taking water from the river to supply commune X through sluices A and
 226 B, as depicted in **Figure 7**.



In which: Q_{t1} is intake discharge at time $t1$; Δt is time step (1 hour); C_d is discharge coefficient; A_0 is area of flow in channel; g is acceleration due to gravity; H_{t1} is water depth at gate, equal to difference of water level in river, Z_R and channel, Z_S .

Figure 7. Diagram describing the method for calculating average salinity for a commune X according to **Equation 3** (as S_X) and **Equation 4** (as S_B).

2.4 Assessing salinity intrusion hazard

Following the IPCC (2014) and IPCC (2022), hazard is defined as the occurrence of a natural or man-made phenomenon, a physical impact, or a trend that is likely to cause loss of life, injury, or other health effects; as well as damage and loss to property, infrastructure, livelihoods, service provision, ecosystems, and environmental resources. Hazard level quantifies the intensity, magnitude, or probability of a particular hazard occurring at a specific location and time. Vulnerability is the tendency or inclination to be adversely affected. Vulnerability encompasses several concepts, including susceptibility to harm and a lack of capacity to cope and adapt. Risk is the potential for loss resulting from the dynamic interaction among the salinity hazard, exposure, and vulnerability. In this paper, hazard refers to the physical characteristics of salinity intrusion (concentration and duration)

241 that pose a threat to the water supply. Hazard level is a quantified classification of the hazard's
242 severity, categorised in this study into five levels (1–5) based on salinity thresholds and total water
243 supply duration.

244

245 There are no specific regulations regarding the salinity threshold for water supply in Vietnam.
246 However, to ensure safety, most Irrigation Operation Companies in the Nam Dinh and Thai Binh
247 provinces only extract water from rivers when the salinity level is below 1‰.

248

249 Salinity levels between 1‰ and 4‰ may be suitable for specific crops (Kotuby-Amacher *et al.*, 2000).
250 Therefore, to assess the hazards and impacts of saltwater intrusion in each commune within the study
251 area, this study adopts a salinity threshold of 4‰. Still, with this threshold, the total duration of the
252 water supply is limited. This results in significant water scarcity for agricultural activities in the
253 affected area. From a hazard perspective, the total water supply duration is also important; a longer
254 duration corresponds to a lower hazard level.

255 Researching on the effects of salinity on crops, Grismer & Bali, 2015; Kotuby-Amacher *et al.*, 2000)
256 mentions a method for determining the crop yield reduction due to salinity as:

257
$$Y = 100 - B(EC_e - A)$$

258 In which Y is the relative yield (%), A is the threshold value (dS/m), B is the percentage of yield
259 reduction, and EC_e is the average salinity in the field. The values of A and B are determined for
260 different crop types.

261 In the study area, the salinity simulation period from January to March is when rice is primarily
262 grown. Therefore, if the water is saline, crop yield may be reduced. According to Grieve *et al.* (2014),
263 the coefficients A and B were determined to be 1.9 and 9.1 % per (dS/m), respectively. Therefore,

264 given the yield reduction in rice plants under saline conditions, the study has classified the hazard
 265 levels associated with average salinity (H_s) into five distinct categories, as shown in **Table 4**.

266

267 **Table 4.** The hazard levels due to average salinity - H_s

Average salinity (‰)	Yield reduction (%)	Hazard level – H_s	Notes
≤ 1	0	0	No hazard
1.01 - 1.50	< 4	1	Very low
1.51 - 2.00	4.1÷11.1	2	Low
2.01 - 2.50	11.2÷18.3	3	Medium
2.51 - 3.20	18.4÷28.2	4	High
3.21 - 4.00	28.3÷39.6	5	Very high

268

269 Considering the total water supply duration, the simulation period from January 1 to March 31 was
 270 determined to be 2163 hours. This is also the dry season in the study area; therefore, the water supply
 271 for water management activities depends mainly on river water. In this study, we assumed that the
 272 sluice gate is fully opened when salinity and water level conditions permit. It can be seen that the
 273 total amount of water extracted through the sluice gate depends on the sluice gate size and the total
 274 supply duration. Therefore, the thresholds in Table 5 are calculated assuming that if the total supply
 275 duration exceeds 2051 hours, it meets over 95% of the water needs, which can be considered a no-
 276 hazard case. Conversely, if the total supply duration is less than 431 hours, it only meets about 20%
 277 of the water needs. Therefore, the hazard levels associated with the total water supply duration (H_T)
 278 have been classified into five distinct levels, as shown in **Table 5**.

279 **Table 5.** The hazard levels due to total water supply duration - H_T

Total supply duration (hours)	Water demand (%)	Hazard level - H_T	Notes
-------------------------------	------------------	----------------------	-------

≥ 2051	95	0	No hazard
1728 - 2050	80÷94.9	1	Very low
1296 – 1727	60÷79.9	2	Low
864 – 1296	40÷59.9	3	Medium
432 – 863	20÷39.9	4	High
≤ 431	19.9	5	Very high

280

281 **Equation 5** evaluates the hazard level at commune (i) for scenario (k) by considering the combined
282 effects of average salinity and total water-supply duration.

$$283 \quad H_{i,k} = w_S \cdot H_{S_{i,k}} + w_T \cdot H_{T_{i,k}} \quad (1)$$

284 where $H_{i,k}$ is combined hazard level at a commune (i) corresponding to scenario (k), $H_{S_{i,k}}$ is hazard
285 level due to the average salinity at a commune (i) corresponding to scenario (k), $H_{T_{i,k}}$ is hazard level
286 due to the total supply water duration at a commune (i) corresponding to scenario (k), w_S and w_T is
287 the weighting factor of the hazard due to average salinity and total water supply duration, $w_S + w_T =$
288 1.

289

290 When salinity levels fall below the allowable threshold, the duration of the water supply becomes a
291 critical factor because it directly determines the total volume delivered. For instance, if two drainage
292 points of equal capacity extract water only when the salinity is below 4‰, the point with a longer
293 extraction period will achieve a greater total intake. In this study, the weighting factors in Equation 5
294 were set at 0.4 for average salinity and 0.6 for total supply duration. The weighting factors were
295 determined using the Analytic Hierarchy Process (AHP). Based on consultation with experts and
296 local water management officials, the results showed that the importance of total supply water
297 duration is approximately 1.5 times that of average salinity. Therefore, these values were chosen to
298 reflect the greater practical importance of maintaining an adequate water supply over time, while still
299 accounting for salinity in water quality. In other words, while salinity determines whether water is

300 usable, the cumulative supply duration governs how much water can be effectively delivered to meet
 301 irrigation and domestic needs. The resulting hazard assessments under this weighting scheme are
 302 summarised in Table 6, with the corresponding hazard levels presented in Table 7.

303

304 **Table 6.** Combined hazard (H) of average salinity hazard (H_s) and total supply water duration hazard
 305 (H_T)

H		H _T					
		0	1	2	3	4	5
H _s	0	0	0.6	1.2	1.8	2.4	3.0
	1	0.4	1.0	1.6	2.2	2.8	3.4
	2	0.8	1.4	2.0	2.6	3.2	3.8
	3	1.2	1.8	2.4	3.0	3.6	4.2
	4	1.6	2.2	2.8	3.4	4.0	4.6
	5	2.0	2.6	3.2	3.8	4.4	5.0

306

307

308 **Table 7.** The classification of the combined hazard (H)

Hazard levels	Combined hazard	Category
1	< 1.50	Very low
2	1.50 – 2.50	Low
3	2.51 – 3.50	Medium
4	3.51 – 4.50	High
5	> 4.50	Very high

309

310 To create a saltwater intrusion hazard map for the study area, we first assessed the hazard value for
311 each commune. Because of the complexity of the canal network, accurately determining the salinity
312 spread within specific hazard zones is difficult. Therefore, we used administrative boundaries to
313 delineate these areas.

314

315 **3. Results and discussion**

316 *3.1 Calibration and validation of the simulation model*

317 The calibration and verification results of the simulation model are presented in **Table 8**. The
318 calibration of model parameters was conducted using data from January 1 to March 31, 2012. The
319 findings show that the highest NSE coefficient, 0.92, was observed at the Son Tay station, where
320 discharge measurements were taken. In contrast, the lowest NSE coefficient, 0.72, was recorded at
321 the Truc Phuong water-level station. Calibrating parameters, such as roughness, in areas affected by
322 tides poses significant challenges, often resulting in lower NSE coefficients. For example, the stations
323 Truc Phuong, Quyet Chien, Cat Khe, Pha Lai, Quang Phuc, and Ben Trieu show less favourable
324 results. In contrast, upstream areas that are not influenced by tidal effects yield better calibration
325 results and higher NSE values. Successful stations in these upstream regions include Son Tay, Hanoi,
326 Thuong Cat, and Hung Yen. Although some stations located in coastal estuaries exhibit poorer results,
327 these outcomes still fall within an acceptable range.

328

329 The validation was conducted from January 1 to March 31, 2011. The model validation results show
330 a pattern like that observed during model calibration. Specifically, the validation results for
331 hydrological stations in the upper region, which is not affected by tides, show higher NSE coefficients
332 than those from coastal plain stations that are significantly influenced by tides. As indicated in **Table**
333 **8**, the NSE efficiency coefficients for the Son Tay, Hanoi, Thuong Cat, and Hung Yen stations are all
334 higher than those of the other stations, where the tidal process has a more substantial impact on the
335 hydrological regime. This difference can be attributed to the greater complexity of hydrological

336 processes (water levels and flows) in the tidal-influence area compared to the upstream region. In
 337 coastal areas, flow processes are governed by upstream flow and tidal influences at river mouths.
 338 Furthermore, as mentioned above, due to limitations in upstream input data, such as daily discharge,
 339 which lack sufficient smoothness compared to the hourly fluctuations in downstream tidal water
 340 levels, discrepancies may occur between calculated and measured values at high/low tides. In
 341 addition, the limited availability of riverbed topography data, which is subject to continuous
 342 sedimentation and erosion, also contributes to this problem. This dual influence leads to irregular,
 343 unpredictable fluctuations, making the model's calibration and verification more challenging and
 344 resulting in lower NSE values at some stations, such as Quyet Chien, which achieves only around
 345 0.70.

346

347 **Table 8.** The NSE coefficient during the calibration and validation of the HD model

No	Hydrological station	Measurement factor	NSE	
			Calibration	Validation
1	Truc Phuong	Water level	0.72	0.72
2	Quyet Chien	Water level	0.88	0.70
3	Cat Khe	Water level	0.77	0.75
4	Pha Lai	Water level	0.80	0.78
5	Son Tay	Discharge	0.92	0.90
6	Hanoi	Discharge	0.92	0.90
7	Thuong Cat	Water level	0.88	0.87
8	Hung Yen	Water level	0.95	0.85
9	Quang Phuc	Water level	0.81	0.81
10	Ben Trieu	Water level	0.82	0.82

348

Salinity observations at the monitoring stations are not recorded continuously but are instead limited to measurements taken during high and low tides each month. This intermittent sampling presents challenges for model calibration and validation. To address these limitations, the study assessed the performance of the AD model parameters using a restricted dataset and evaluated accuracy through the absolute error (ΔS) indicator. The difference between the maximum measured and simulated salinity values ranges from 0 to 2.18‰. Most stations exhibit a difference of less than 0.5‰ in the calibration (**Table 9**) and validation steps (**Table 10**).

Table 9. Absolute errors associated with the calibration of the AD model parameters at selected monitoring stations

No.	Monitoring station / Sluice	Smax (‰)		ΔS (‰)
		Simulated	Measured	
1	Tai	16.57	16.26	0.02
2	Ngo Dong	21.32	21.70	0.38
3	Xe	8.82	11.00	2.18
4	Con Nhat	16.16	16.00	0.16
5	Duong Lieu	0.33	0.14	0.19
6	Phuc Khe	0.41	0.36	0.05
7	Kinh Khe	1.72	1.40	0.32
8	Tien Tien	0.63	0.70	0.07
9	Cao Kenh	0.63	0.30	0.33
10	An Phu	0.05	0.05	0.00

360 **Table 10.** Absolute errors associated with the validation of the AD model parameters at selected
 361 monitoring stations

No	Monitoring station / Sluice	Smax (%)		ΔS (‰)
		Simulated	Measured	
1	Phu Le	24.5	25	0.5
2	Dong Quy	26	24	2.0
3	Trung Trang	1.1	1.03	0.07
4	Dong Xuyen	11.7	12.1	0.4
5	Cao Kenh	1.3	1.29	0.01
6	Ba Lat	27.5	29.5	2.0
7	Don Son	2.2	2.3	0.1

362

363 The calibration and validation results show good agreement between the measured and observed
 364 values of the HD and AD modules, with reasonable trends in peak timing and magnitude. Despite
 365 some limitations in the salinity data and the MIKE 11 model's simplified advection and dispersion,
 366 the model performs adequately for simulating various scenarios in the study area.

367

368 *3.2 Results of the salinity intrusion simulation scenarios*

369 Although the study simulated hydraulic and water-quality dynamics across the entire RTB river
 370 system, the subsequent analysis focuses specifically on the Day, Ninh Co, Red, Tra Ly, and Thai Binh
 371 rivers. These rivers are regulated by sluice systems that supply water to all coastal communes in Nam
 372 Dinh and Thai Binh provinces.

373

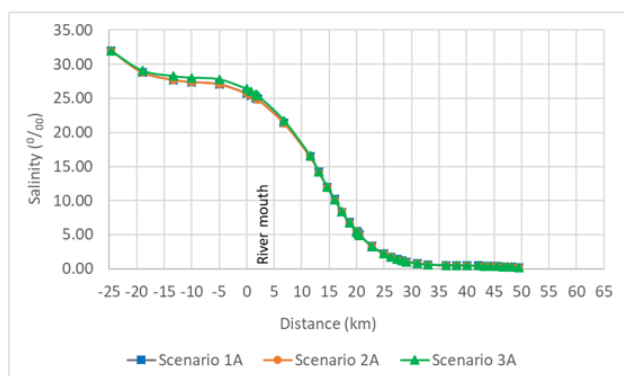
374 *3.2.1 The process of saltwater intrusion along the rivers*

375 (1) Variations in upstream flow (Scenario group A)

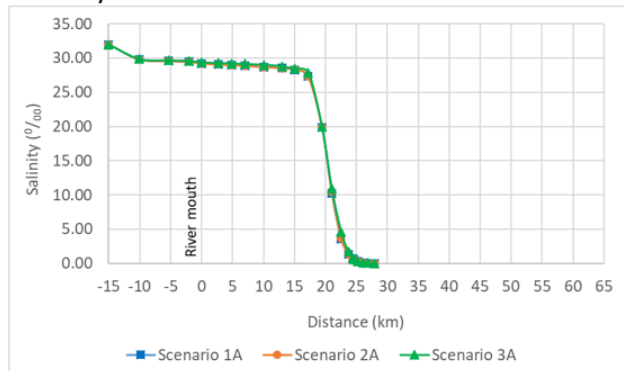
376 Generally, if the tide level remains constant, an increase in upstream flow leads to a decrease in the
377 maximum downstream salinity (**Table 11**) and a reduction in saltwater intrusion inland (**Figure 8**).

378 **Table 11.** The maximum salinity at some sluices on the rivers

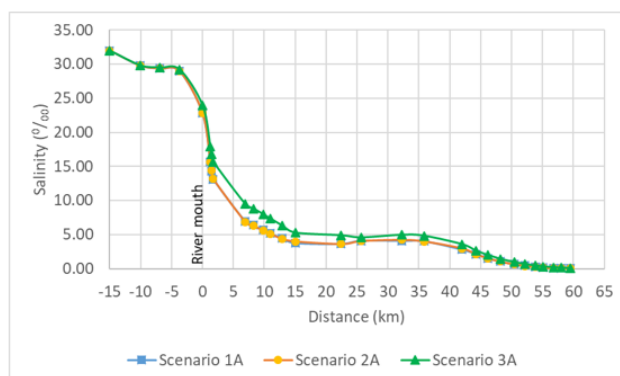
River	Sluice	Maximum salinity (‰)		
		Scenario 1A	Scenario 2A	Scenario 3A
Day	Quy Nhat	11.98	12.00	12.02
	Thuan Hau	8.32	8.36	8.39
	Dong Lieu	4.90	4.96	4.98
	Tam Toa	0.60	0.61	0.63
Red	Ngo Dong	4.10	4.24	4.95
	Mo Dao	3.77	4.01	4.28
	Co Le	2.80	2.98	3.56
Ninh Co	Muc I	12.12	12.17	12.19
	Muc II	8.42	8.45	8.47
Tra Ly	Thai Phuc	5.05	5.20	6.25
	Nam Cuong	0.04	0.04	0.05
Thai Binh	Van Am	1.15	1.16	1.30
	Thu Cuc	0.61	0.63	0.73



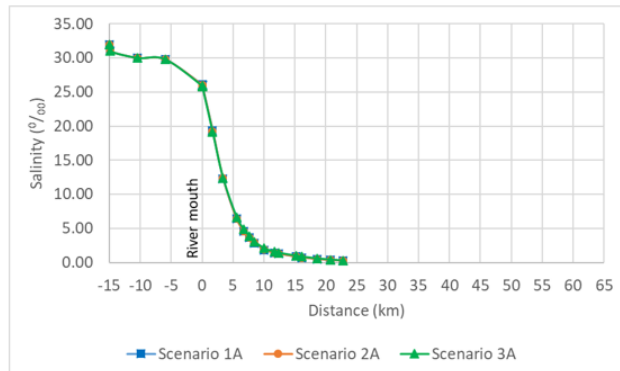
The Day river



The Tra Ly river



The Red river



The Thai Binh river

Figure 8. Salinity profiles of the studied rivers according to group A scenarios.

On the Day River, all simulation scenarios show similar trends in changes in maximum salinity along the river. Salinity decreases sharply from the river mouth toward the mainland. Specifically, between KM0 (the estuary) and KM22 (the mainland), salinity drops significantly from about 25‰ to less than 4‰. Beyond KM22, the maximum salinity along the river continues to decline steadily. The differences in maximum salinity along the river under different scenarios are relatively small, ranging from 0.2‰ to 5‰. There is no significant reduction in the extent of saltwater intrusion into the mainland across the scenarios. In Scenario 3A, the distance from the river mouth to the point where salinity fell below 4‰ was 22 km, whereas in Scenario 1A it decreased slightly to 21 km. Likewise, the distance to where salinity falls below 1‰ decreased from 31 km in Scenario 3A to 29 km in Scenario 1A.

The Thinh Long Estuary on the Ninh Co River and the Ba Lat Estuary on the Red River are significantly affected by tidal influences. Additionally, the distribution of upstream flow from the Red River impacts salinity levels. As a result, most locations along the Ninh Co River experience maximum salinity levels

395 exceeding 4‰ in all scenarios. However, the maximum salinity along the river quickly decreases from
396 26‰ (KM0) to 4‰ (KM45). The difference in maximum salinity at sluices between scenarios 1A, 2A,
397 and 3A is minimal.

398

399 The Red River is the main river in the study region and has the strongest flow. Because of this, the
400 highest salinity drops quickly from 23‰ at KM0 to about 5‰ at KM15. Comparing scenarios 1A and
401 2A, the change in maximum salinity at some locations is less than 6%, whereas between scenarios 1A
402 and 3A, the change can be as significant as 40%. Comparing the results between scenarios 1A and 3A
403 reveals a significant difference in the maximum salinity at several locations along the Red River. This
404 phenomenon can be attributed to the non-linear interaction between upstream discharge and tidal
405 energy. The Red River (main branch) has a higher discharge than other rivers regulated by upstream
406 reservoirs. The small increase in discharge in scenario 1A compared to 3A creates a greater "pushing
407 force" leading to lower salinity. Furthermore, the Red River's estuary has a more concentrated, deeper
408 channel than the funnel-shaped estuaries of the Day River. When upstream flow increases, the
409 velocity in the deep channel increases rapidly, directly neutralising the tidal energy entering the river.
410 The distance from the river mouth to the position with salinity below 4‰ decreases from 37 km
411 (scenario 3A) to 35 km (scenario 1A), and the position with salinity below 1‰ decreases from 50 km
412 (scenario 3A) to 48 km (scenario 1A).

413

414 On the Tra Ly River, the salinity decreases slightly from the estuary up to around KM15, dropping from
415 29‰ to 28‰. However, between KM15 and KM22, the maximum salinity falls below 4‰. When
416 comparing different scenarios, the maximum salinity levels indicate a slight decrease in the extent to
417 which saltwater intrudes into areas. Specifically, the distance from the river mouth to locations with
418 salinity below 4‰ decreases from 23 km (in scenario 3A) to 22 km (in scenario 1A). For locations with
419 salinity below 1‰, it also decreases from 23 km (in scenario 3A) to 22 km (in scenario 1A).

420

421 The salinity of the Thai Binh River decreases significantly from the estuary to KM9, ranging from
422 26‰ to less than 3‰. The decrease is relatively steady upstream of the river. Across scenarios, the
423 maximum salinity and the distance of saltwater intrusion remain relatively unchanged. The distance
424 of saltwater intrusion decreases from 14 km (scenario 3A) to 12 km (scenario 1A).

425

426 The results indicate that salinity intrusion in the Day, Ninh Co, and Red rivers in Nam Dinh province
427 is more severe than in the Tra Ly and Thai Binh rivers in Thai Binh province. This difference
428 primarily stems from the distribution of water throughout the system during the dry season. The flow
429 regulation from upstream to the Tra Ly and Thai Binh rivers is more abundant and stable compared
430 to that of the Ninh Co and Day rivers, which have weaker and less stable flows.

431

432 Due to the stronger flow in Thai Binh province, rivers such as the Tra Ly create an effective "wall of
433 fresh water" that helps prevent salinisation. Additionally, the riverbed terrain plays a role; for
434 instance, the mouth of the Day River is wide and relatively shallow, allowing salinity to penetrate
435 easily during high tides. In contrast, the Tra Ly river mouth is narrow and deep, with strong freshwater
436 flow, which minimises salinity intrusion into this river.

437

438 This study evaluates the maximum salinity levels and the extent of river saltwater intrusion. It also
439 determines the total duration of water supply at the sluices in the study area. The simulation period
440 spans January to March, during which the total water supply duration could reach a maximum of
441 2,160 hours if salinity remains below the allowable threshold and the river level exceeds the sluice
442 bed elevation. Generally, the total water supply duration at most sluices increases slightly between
443 scenarios, ranging from 1 to 3 hours. However, at some sluices, such as the Muc I on the Ninh Co
444 River, it experiences a more significant increase, rising from 70 to 93 hours (**Table 12**).

445

446 **Table 12.** Duration of total water supply at key sluices along the rivers

River	Sluice	Duration (hours) with salinity under 4‰		
		Scenario 1A	Scenario 2A	Scenario 3A
Day	Quy Nhat	1848	1837	1834
	Thuan Hau	1945	1930	1930
	Dong Lieu	1988	1986	1987
Red	Ngo Dong	1049	1046	1027
	Mo Dao	1997	1992	1984
Ninh Co	Muc I	716	646	623
	Muc II	1690	1658	1636

447

448 Our analysis of water supply duration across different salinity levels revealed varying changes. When
449 salinity is below 1‰, the total water supply duration typically decreases due to lower upstream flow.
450 For example, at the Mo Dao sluice, it drops from 1,728 hours in scenario 1A to 1,586 hours in scenario
451 3A. Conversely, when salinity is between 3.2 and 4‰, the duration increases, such as at the Quy Nhat
452 sluice, where it rises from 34 to 43 hours, and at Muc II, from 83 to 119 hours.

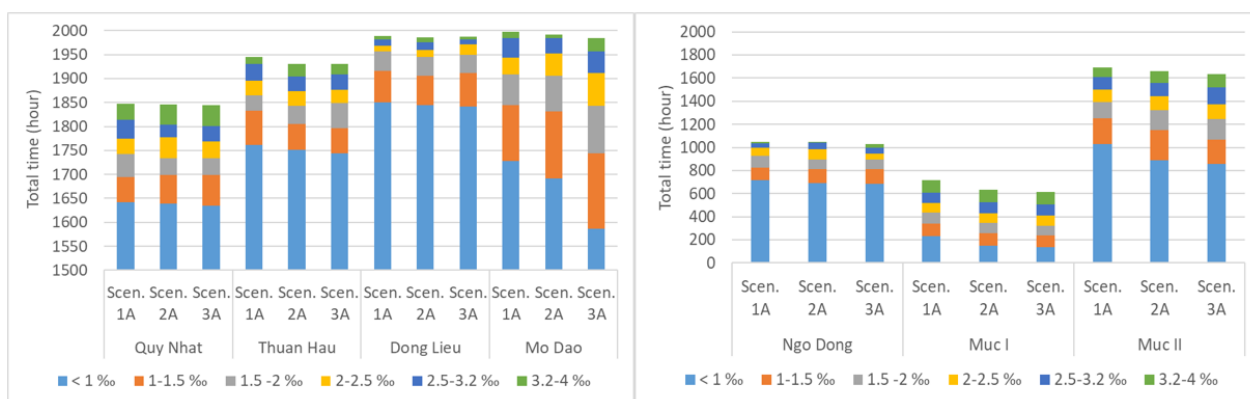
453 **Figure 9** shows how the total water supply duration varies across salinity levels under different
454 scenarios.

455

456

457

458



459

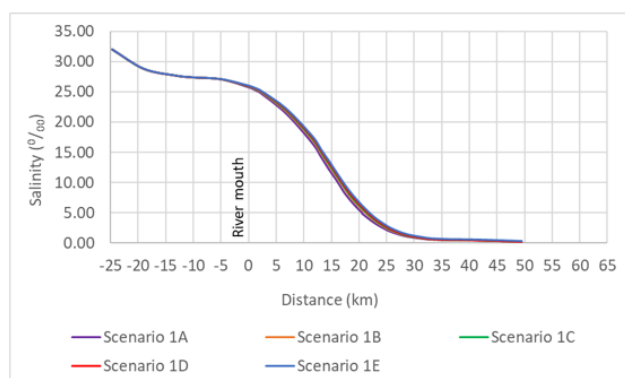
460 **Figure 9.** The total duration of water supply at various sluices for different salinity levels, according
 461 to different scenarios

462

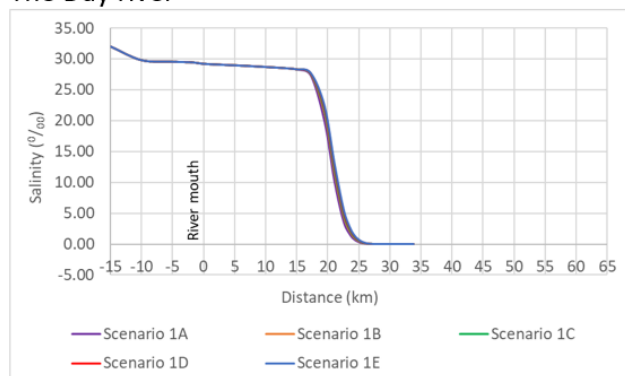
463 (2) SLR scenarios

464 SLR is also a cause of salinity intrusion into rivers, along with upstream flow and topographic
 465 conditions. As SLR increases, both the maximum salinity at sluices and the extent of saltwater
 466 intrusion into the continent also rise (**Figure 10**). Furthermore, when there is reduced upstream flow
 467 coupled with elevated tide levels, the maximum salinity at the sluices increases, as illustrated in

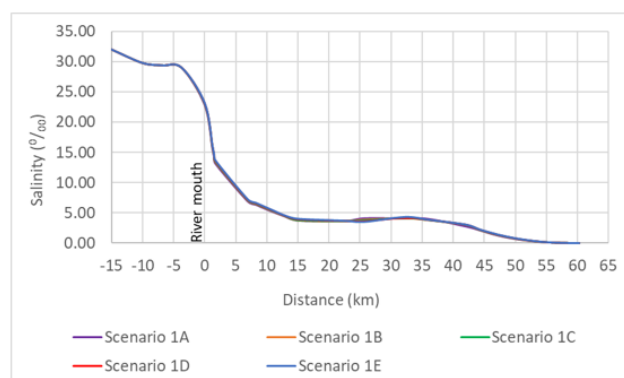
468 **Figure 11.**



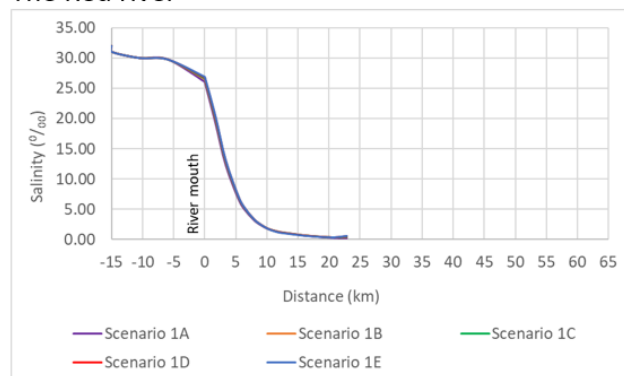
The Day river



The Tra Ly river



The Red river



The Thai Binh river

Figure 10. Salinity profiles of the studied rivers under different SLR scenarios.

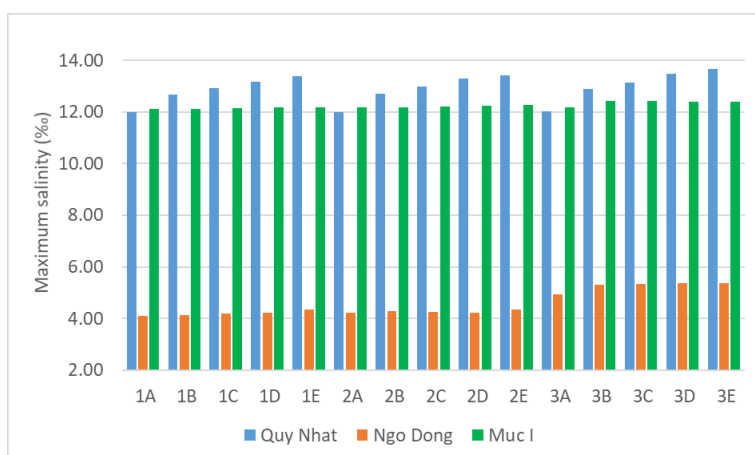
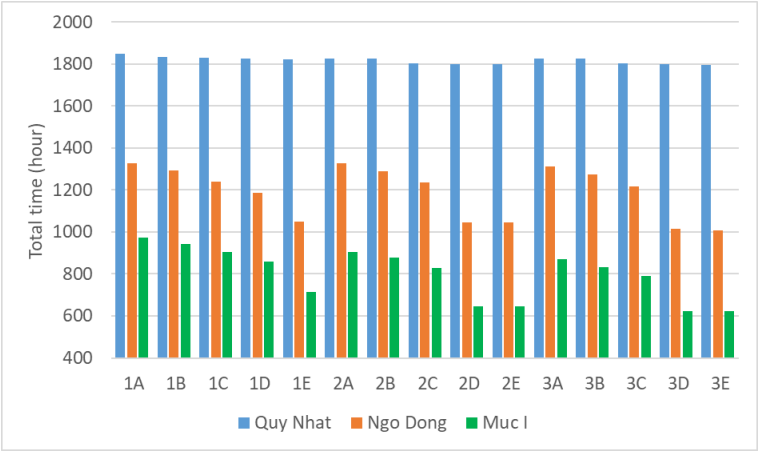


Figure 11. Maximum salinity at various sluice levels across different scenarios.

Compared to scenario group A, the maximum increase in salinity at the sluices on the Day River due to SLR is significantly greater, ranging from 6% to 39%. The saline increase on the Thai Binh River rose from 6% to 26%. Other rivers experienced much smaller increases, ranging from 1% to 10%. Additionally, the distance of saltwater intrusion into the land has increased by 1-2 kilometres. When

479 salinity levels fall below the permitted threshold, the total duration of the water supply varies across
480 different scenarios. For some sluices, such as Dong Lieu, Mo Dao, and Muc II, the decrease in supply
481 is minimal, ranging from 1 to 8 hours. However, other sluices, such as Ngo Dong and Muc I,
482 experience significant reductions, with decreases ranging from 32 to 305 hours (**Figure 12**).

483



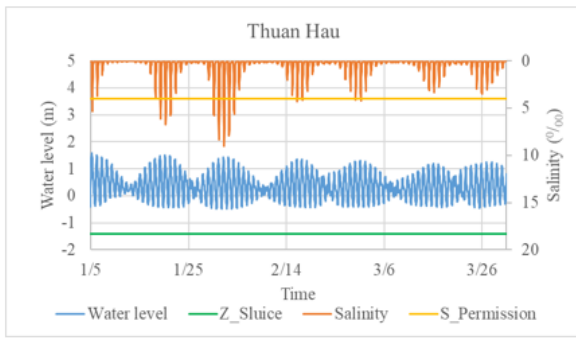
484

485 **Figure 12.** Total water-supply duration at various sluices across different scenarios.

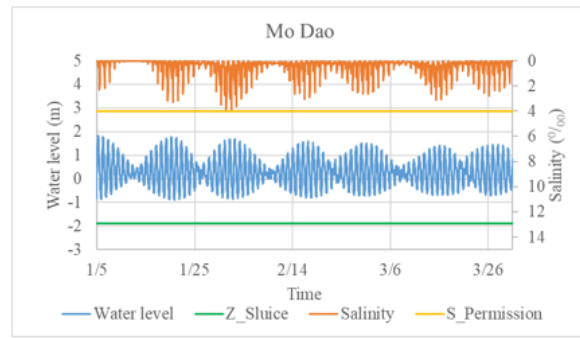
486

487 *3.2.2 The average salinity and hazard levels in communes*

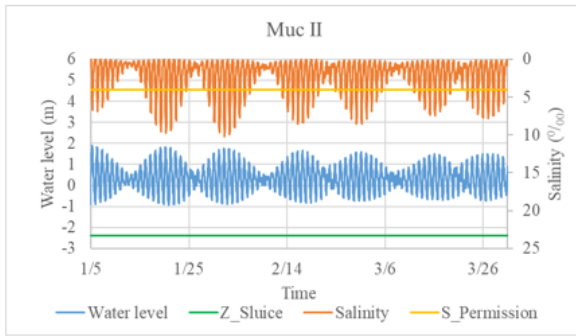
488 **Figure 13** presents water-level and salinity measurements at various sluice gates throughout the
489 simulation period across different scenarios. Using this data, we can calculate the average salinity for
490 each commune by considering the total amount of water supplied from each sluice, as detailed in
491 section 2.3.



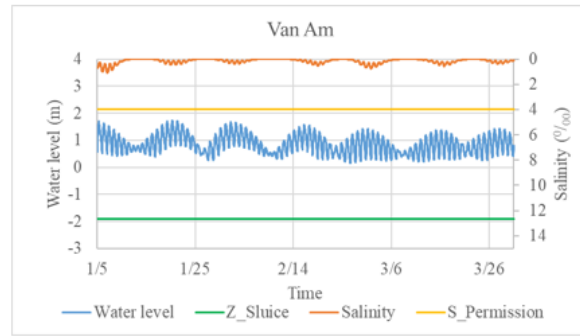
Thuan Hau sluice, Day River, scenario 1B.



Mo Dao sluice, Red river, scenario 1C.



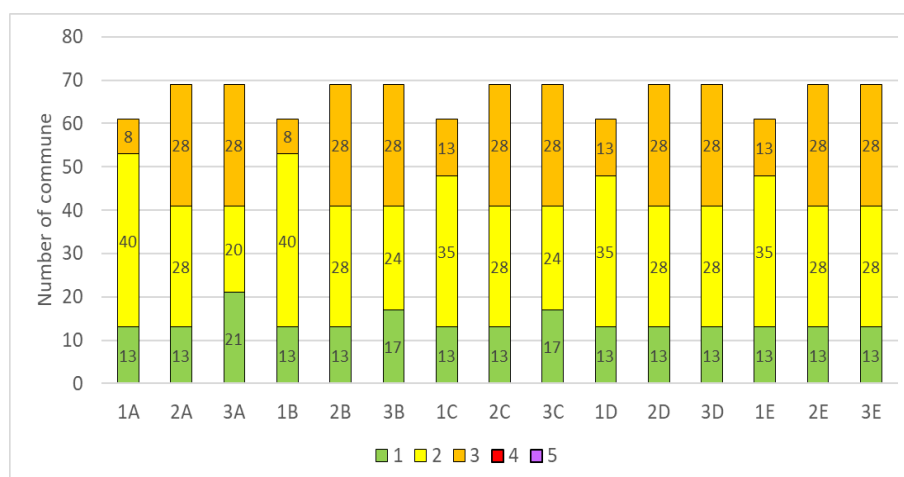
Muc II sluice, Ninh Co river, scenario 2D.



Van Am sluice, Thai Binh river, scenario 2E.

Figure 13. Water level and salinity graphs at selected sluices for various scenarios

An assessment of 123 communes in Nam Dinh province reveals that the number of communes exposed to salinity-related hazards has risen from 61 in scenario group 1 to 69 in scenario groups 2 and 3. Among these communes exposed to hazards at levels 1 to 3, the highest proportion, 60%, is at hazard level 2. A detailed analysis reveals that the number of communes at each hazard level varies with fluctuations in upstream flow and tide levels. For example, when comparing scenarios 1D, 2D, and 3D, the total number of communes exposed to hazard level 3 rises from 13 in scenario 1D to 28 in scenario 2D and 29 in scenario 3D (**Figure 14**).



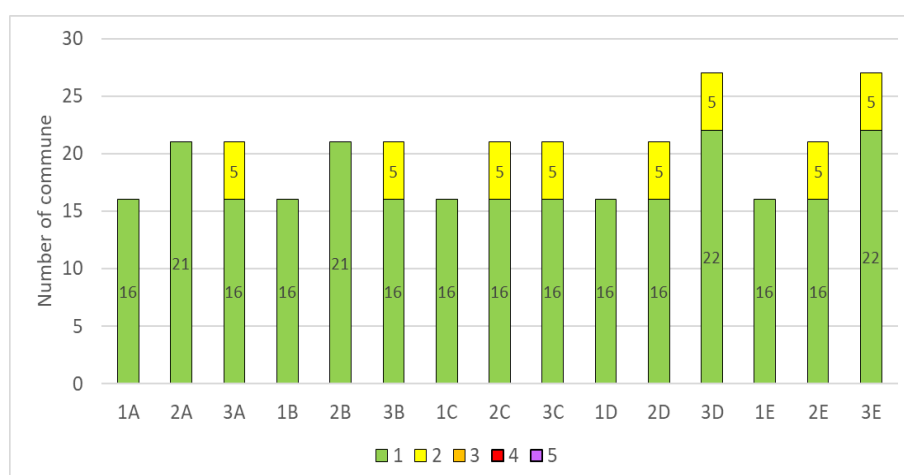
502

503 **Figure 14.** The total number of communes in Nam Dinh province affected by salinity hazards across
 504 various scenarios.

505

506 An assessment of 120 communes in Thai Binh province reveals that the total number of communes
 507 exposed to hazards due to the salinity factor is significantly lower than in Nam Dinh province. Only
 508 16 to 21 communes, most at level 1, are exposed to hazards. Due to upstream flow and SLR, some
 509 communes changed from no hazard to level 1, while others changed from level 1 to level 2 (**Figure**
 510 **15**). This is understandable because the salinity of the Thai Binh River is lower than that of other
 511 rivers, and there is no water supply issue if the salinity remains below the allowable threshold.

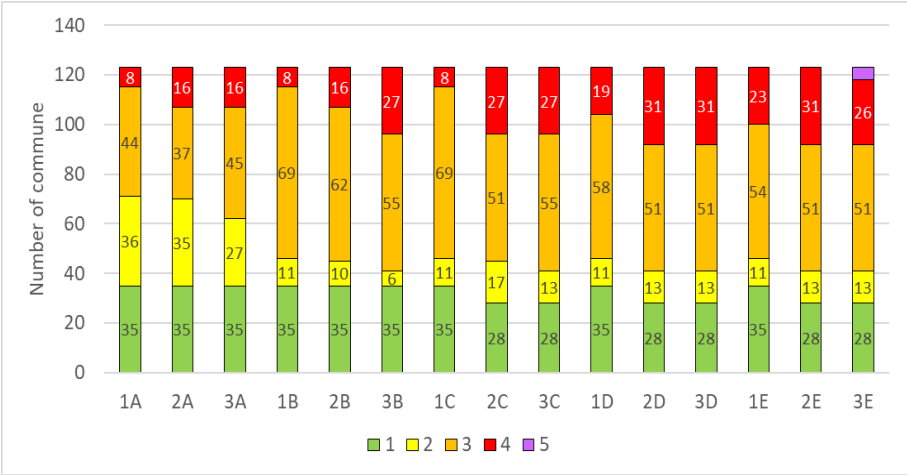
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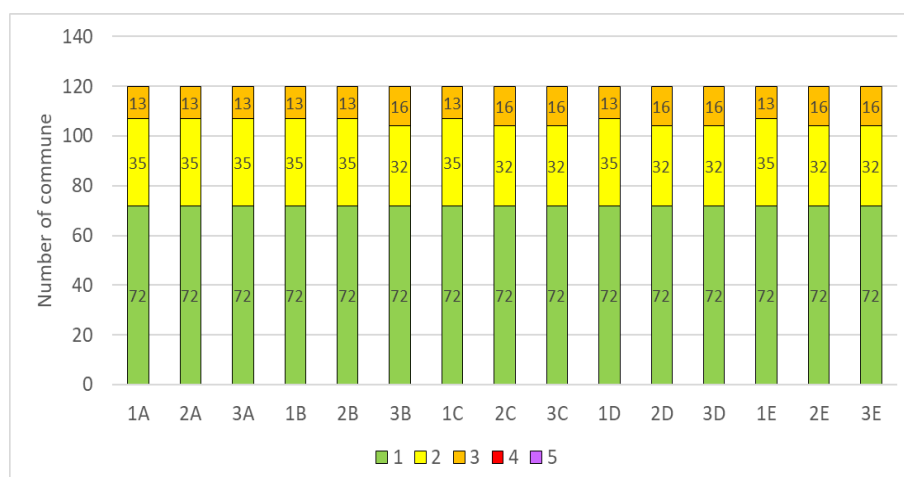
514 **Figure 15.** The number of communes in Thai Binh province facing various salinity hazard levels
 515 across scenarios.

516 Assessment of the hazard level due to water supply duration: all communes in Nam Dinh province
517 are exposed to hazard levels of 1-5. The percentage of communes at level 3 is the highest, ranging
518 from 30% to 56% across scenarios. Additionally, the number of communes at various levels increases
519 as upstream flow decreases and sea levels rise (**Figure 16**).



520
521 **Figure 16.** The number of communes in Nam Dinh province facing various hazard levels due to
522 different water supply durations

523
524 Unlike Nam Dinh province, Thai Binh province's water supply duration is longer due to the lower
525 impact of saline intrusion. All 120 communes in the study area are exposed to hazard levels ranging
526 from 1 to 3, with 60% at level 1. While the number of communes at each level varies across scenarios,
527 the change is minimal. Only three communes shift from levels 2 to 3 in scenario groups 2 and 3
528 (**Figure 17**).

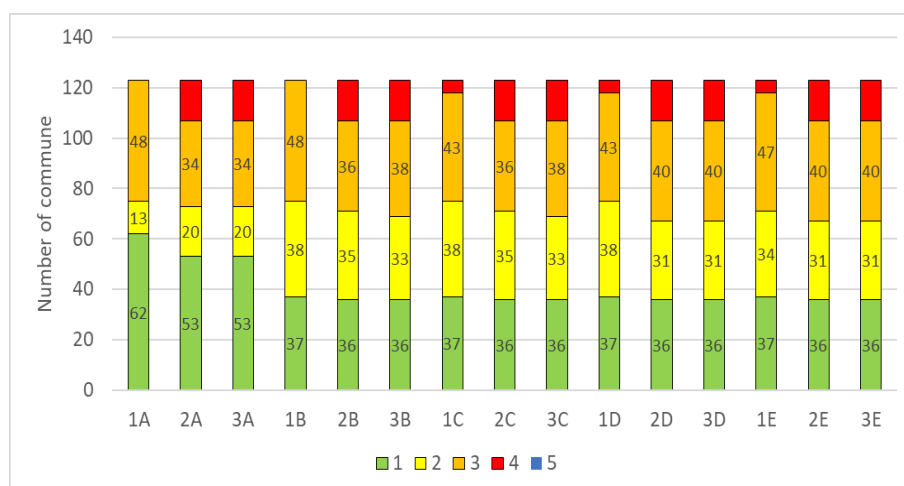


531

532 **Figure 17.** The total number of communes in Thai Binh province affected by varying hazard levels
 533 related to the duration of water supply

534

535 The interaction between salinity levels and water-supply duration can influence the hazard posed by
 536 saltwater intrusion. All communes in the study area of Nam Dinh province are exposed to hazard
 537 levels 1-4. The total number of communes at levels 2 and 3 accounts for approximately 40-60%,
 538 depending on the scenario. Due to reduced upstream flow, many communes have moved from low to
 539 higher hazard levels. For instance, between scenarios 1A, 2A, and 3A, seven communes have moved
 540 from level 1 to level 2, and 16 communes have moved from level 3 to level 4. Additionally, the
 541 number of communes vulnerable to high hazards has increased due to SLR. For example, between
 542 scenarios 1A and 1B, the number of communes exposed to hazard level 1 has decreased from 62 to
 543 37, while the number of communes at hazard level 2 has increased from 13 to 38 (**Figure 18**). In this
 544 province, some communes identified as having the highest hazard levels are Hai Chau, Hai Hoa, Hai
 545 Ninh, Hai Trieu, and Thinh Long, all at level 4.

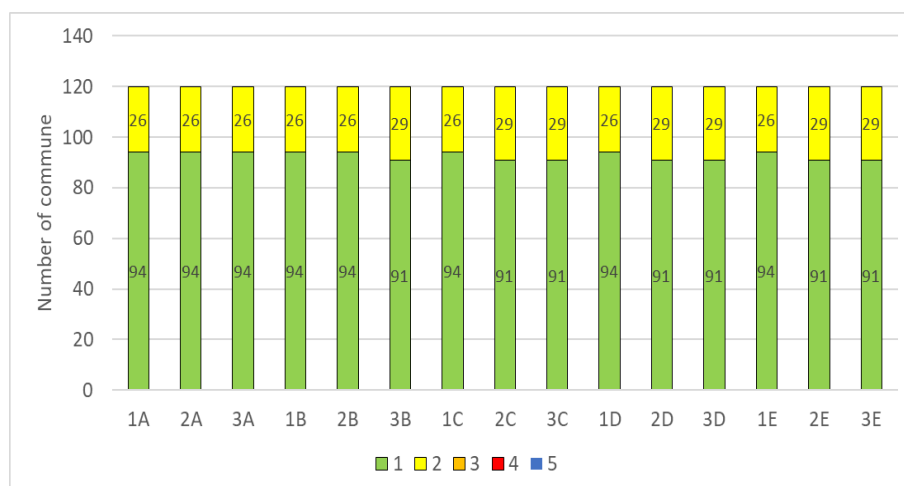


546

547 **Figure 18.** The total number of communes in Nam Dinh province exposed to different hazard levels
 548 under various scenarios

549

550 In Thai Binh province, all communes are only exposed to hazards at levels 1 and 2. Approximately
 551 75% of them are exposed to level 1 hazards (**Figure 19**). The change in the number of communes
 552 exposed to hazards at different levels remains almost constant across scenarios. Some communes
 553 exposed to the highest hazard levels include Nam Chinh, Nam Hong, and Tay Phong, all of which
 554 are categorised as level 2.



555

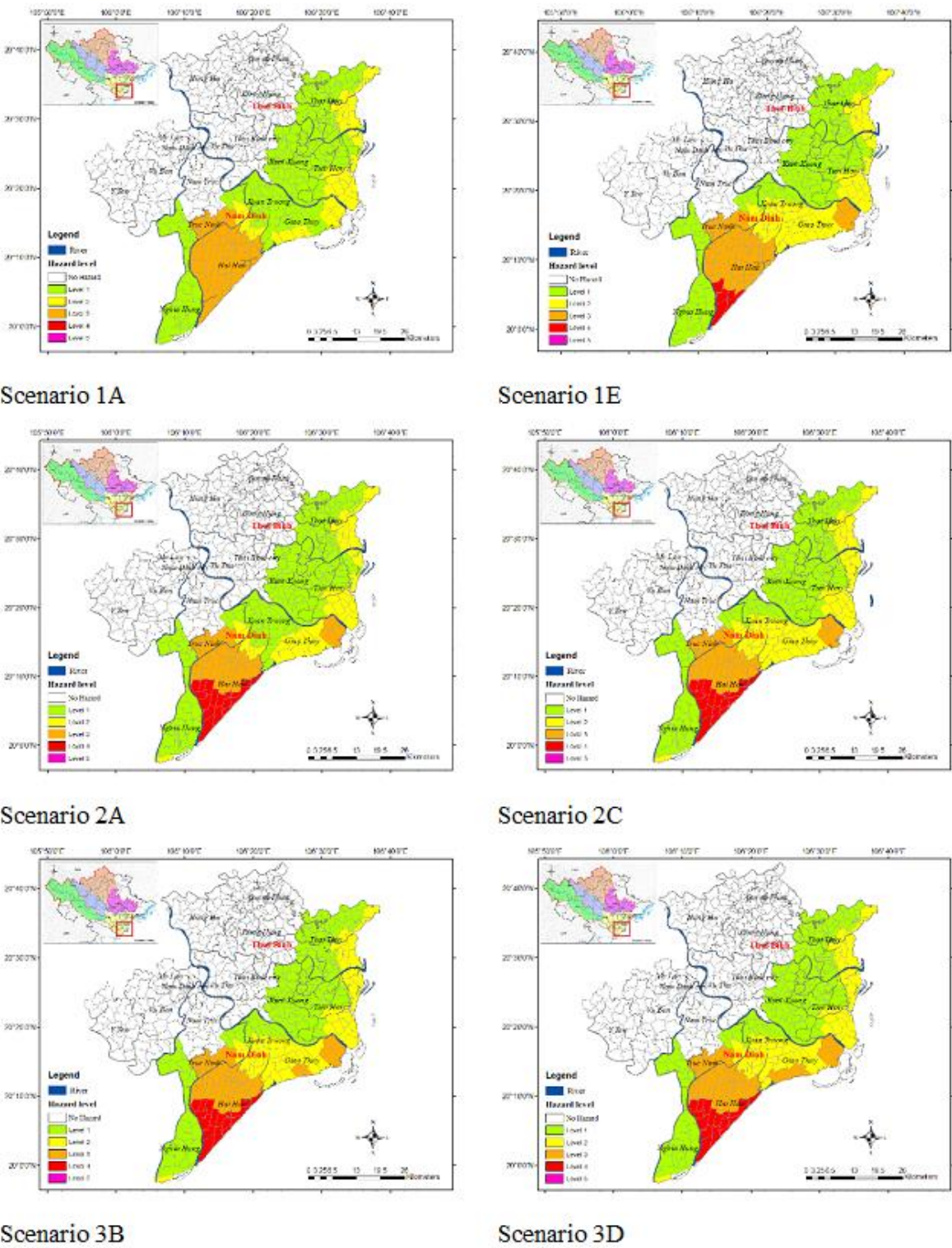
556 **Figure 19.** The number of communes in Thai Binh province exposed to various hazard levels under
 557 different scenarios

558

559

561 3.2.3 Saltwater intrusion hazard maps

562 Based on analyses of average salinity and water-supply duration, a saltwater intrusion hazard map
563 was constructed for the 243 communes of the study area for the different scenarios. **Figure 20** shows
564 the changes in hazard level across communes due to upstream flow conditions and SLR.



566 **Figure 20.** Saltwater intrusion hazard maps of the study area for different scenarios

567 Overall, communes in downstream areas are subject to greater saltwater intrusion hazards than those
568 upstream. Districts such as Hai Hau, Giao Thuy, Tien Hai, and Thai Thuy face elevated risk because
569 their water is drawn directly from river sluice gates, where salinity tends to be higher. In contrast,
570 communes with well-developed canal networks and effective distribution of low-salinity water from
571 upstream sources, such as those in the Nghia Hung district, exhibit comparatively lower hazard levels.

572

573 This study takes a more comprehensive approach to mapping salinity intrusion boundaries than
574 previous research, which focused solely on the maximum simulated salinity along major rivers.
575 Earlier methods often led to misrepresentations of salinity distribution in the RTB River Basin by
576 neglecting inland variations and secondary channels. In contrast, the saltwater intrusion hazard map
577 developed in this study offers a more accurate depiction of the spatial extent and intensity of the
578 affected areas. This improved representation enhances our understanding of salinity dynamics. It
579 allows for the identification of priority zones for management interventions, thereby supporting the
580 development of targeted strategies to mitigate the impacts of saltwater intrusion.

581

582 **4. Conclusions**

583 Saltwater intrusion is a complex phenomenon influenced by factors such as upstream flow patterns,
584 tidal dynamics, and other environmental conditions. Understanding how these elements interact is
585 essential for assessing the impacts of saltwater encroachment in coastal and estuarine areas. This
586 study used the MIKE 11 model to simulate saltwater intrusion in the RTB river basin by examining
587 15 scenarios that account for variations in upstream flow during the dry season and the effects of
588 SLR. The findings suggest that maximum salinity levels in the river and the extent of saltwater
589 intrusion are likely to increase due to reduced upstream flow and/or rising sea levels.

590

591 When analysing the impact of upstream flow changes on salinity levels at sluice gates, maximum
592 salinity remains relatively stable across most rivers, with variations typically under 6%. However, the
593 Red River shows more significant fluctuations, with salinity changes reaching up to 40%. In terms of
594 SLR, both the Day River and the Thai Binh River experience more substantial salinity changes
595 compared to other rivers. In all scenarios, saltwater intrusion extends 1 to 2 kilometres inland.

596

597 This paper presents a new approach to estimating the average salinity during the water supply period
598 when salinity levels fall below the allowable threshold. The study determined the average salinity and
599 the total duration of water supply for 123 coastal communes in Nam Dinh province and 120 in Thai
600 Binh province. A comparison of results across scenarios reveals significant variations in average
601 salinity and total water-supply duration, reflecting the different hazard levels experienced by each
602 commune. All communes in Nam Dinh province face higher hazard levels compared to those in Thai
603 Binh province. Specifically, the hazard levels in Thai Binh province range only from 1 to 2, while in
604 Nam Dinh province, they range from 1 to 4.

605

606 Unlike previous studies that created saltwater intrusion boundary maps for the RTB River Basin, the
607 hazard maps developed in this study provide a more accurate representation of the extent of saltwater
608 intrusion and the associated hazard levels, despite some limitations in the available observational
609 network. This improved precision is achieved by considering both the average salinity and the total
610 duration of the water supply for the canal system. The findings of this study provide valuable insights
611 for assessing risks for the development of adaptation strategies to mitigate the impacts of saltwater
612 intrusion in the study region.

613

614 Further research is recommended in the following areas: (1) measuring cross-sections and salinity
615 levels in the rivers and canals of the irrigation system. This data would be used to simulate
616 hydrodynamics and salinity transmission throughout the entire system. Comparing these findings

617 with the current results will provide a scientific basis for model calibration and validation, ensuring
618 that the results align more closely with real-world conditions. (2) Investigating the operation of
619 upstream reservoirs to enhance freshwater discharge, which could help reduce salinity levels during
620 periods of strong tides in the downstream areas.

621

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627

628 **References**

- 629 Bear, J., Cheng, A., Sorek, S., Ouazar, D., & Herrera, I. (1999). *Seawater intrusion in coastal*
630 *aquifers: concepts, methods, and practices*.
631 <https://api.semanticscholar.org/CorpusID:127213695>
- 632 Becker, M. L., Luettich, R. A., & Mallin, M. A. (2010). Hydrodynamic behavior of the Cape Fear
633 River and estuarine system: A synthesis and observational investigation of discharge-salinity
634 intrusion relationships. *Estuarine, Coastal and Shelf Science*, 88(3), 407–418.
635 <https://doi.org/10.1016/j.ecss.2010.04.022>
- 636 Bricheno, L. M., Wolf, J., & Sun, Y. (2021). Saline intrusion in the Ganges-Brahmaputra-Meghna
637 megadelta. *Estuarine, Coastal and Shelf Science*, 252, 107246.
638 <https://doi.org/https://doi.org/10.1016/j.ecss.2021.107246>
- 639 Chen, W., Chen, K., Kuang, C., Zhu, D. Z., He, L., Mao, X., Liang, H., & Song, H. (2016). Influence
640 of sea level rise on saline water intrusion in the Yangtze River Estuary, China. *Applied Ocean*
641 *Research*, 54, 12–25. <https://doi.org/10.1016/j.apor.2015.11.002>
- 642 Devienne, S. (2006). Red River Delta: Fifty Years of Change. *Moussons*, 9–10, 255–280.

643 <https://doi.org/https://doi.org/10.4000/moussons.2042>

644 DHI. (2014). *MIKE 21 Toolbox - Global Tide Model - Tidal prediction* (p. 20). DHI.

645 Doulgeris, C., Georgiou, P., Papadimos, D., & Papamichail, D. (2012). Ecosystem approach to water
646 resources management using the MIKE 11 modeling system in the Strymonas River and Lake
647 Kerkini. *Journal of Environmental Management*, 94(1), 132–143.
648 <https://doi.org/https://doi.org/10.1016/j.jenvman.2011.06.023>

649 Erdbrügger, J., van Meerveld, I., Seibert, J., & Bishop, K. (2023). Temporal and spatial variation in
650 shallow groundwater gradients in a boreal headwater catchment. *Journal of Hydrology*, 626(PB),
651 130301. <https://doi.org/10.1016/j.jhydrol.2023.130301>

652 Ferdous, A. (2014). Cumulative Hydrologic Impact of Wetland Loss: Numerical Modeling Study of
653 the Rideau River Watershed, Canada. *Journal of Hydrologic Engineering*, 19(3), 593–606.
654 [https://doi.org/10.1061/\(ASCE\)HE.1943-5584.0000817](https://doi.org/10.1061/(ASCE)HE.1943-5584.0000817)

655 Ferguson, G., & Gleeson, T. (2012). Vulnerability of coastal aquifers to groundwater use and climate
656 change. *Nature Climate Change*, 2(5), 342–345. <https://doi.org/10.1038/nclimate1413>

657 Gaiolini, M., Colombani, N., Mastrocicco, M., & Postacchini, M. (2023). Seawater intrusion
658 assessment along the Volturno River (Italy) via numerical modeling and spectral analysis.
659 *Journal of Hydrology*, 626(October). <https://doi.org/10.1016/j.jhydrol.2023.130289>

660 Gong, W., Shen, J., & Jia, L. (2013). Salt intrusion during the dry season in the Huangmaohai Estuary,
661 Pearl River Delta, China. *Journal of Marine Systems*, 111–112, 235–252.
662 <https://doi.org/https://doi.org/10.1016/j.jmarsys.2012.11.006>

663 Grismer, M. E., & Bali, K. M. (2015). Drought Tip: Use of Saline Drain Water for Crop Production.
664 *Drought Tip: Use of Saline Drain Water for Crop Production*, October.
665 <https://doi.org/10.3733/ucanr.8554>

666 Haddout, S., Maslouhi, A., & Magrane, B. (2018). A 2D fluid motion model of the estuarine water
667 circulation: Physical analysis of the salinity stratification in the Sebou estuary. *The European
668 Physical Journal Plus*, 133(2), 79. <https://doi.org/10.1140/epjp/i2018-11890-6>

669 Haque, A., Hossain, K. M. A., & Shadia, N. (2021). Coupled 1D Hydrodynamic-Water Quality Model
 670 of Kobadak-Sibsa River System for Salinity. *International Journal of Environmental Science
 671 and Development*, 12(10), 311–315.

672 Hebert, D. L. (2019). Intrusions. *Encyclopedia of Ocean Sciences*, 555–559.
 673 <https://doi.org/10.1016/B978-0-12-813081-0.00148-8>

674 Hien, N. T., Ranzi, R., & Cat, V. M. (2018). Potential Use of Reservoirs for Mitigating Saline
 675 Intrusion in the Coastal Areas of Red River Delta. *Proceedings - 2018 IEEE International
 676 Conference on Environment and Electrical Engineering and 2018 IEEE Industrial and
 677 Commercial Power Systems Europe, IEEEIC/I and CPS Europe 2018, November*.
 678 <https://doi.org/10.1109/IEEEIC.2018.8493799>

679 Hien, N. T., Yen, N. H., Balistocchi, M., Peli, M., Cat, V. M., & Ranzi, R. (2023). Salinity dynamics
 680 under different water management plans coupled with sea level rise scenarios in the Red River
 681 Delta, Vietnam. *Journal of Hydro-Environment Research*, 51(October), 67–81.
 682 <https://doi.org/10.1016/j.jher.2023.10.003>

683 Hong, B., Liu, Z., Shen, J., Wu, H., Gong, W., Xu, H., & Wang, D. (2020). Potential physical impacts
 684 of sea-level rise on the Pearl River Estuary, China. *Journal of Marine Systems*, 201, 103245.
 685 <https://doi.org/https://doi.org/10.1016/j.jmarsys.2019.103245>

686 Jeong, S., Yeon, K., Hur, Y., & Oh, K. (2010). Salinity intrusion characteristics analysis using EFDC
 687 model in the downstream of Geum River. *Journal of Environmental Sciences*, 22(6), 934–939.
 688 [https://doi.org/10.1016/S1001-0742\(09\)60201-1](https://doi.org/10.1016/S1001-0742(09)60201-1)

689 Kaddi, Y., Cierco, F. X., Faure, J. B., & Proust, S. (2022). New Developments in a 1D+ ISM Model
 690 for Operational Purposes. In *Springer Water*. Springer Nature Singapore.
 691 https://doi.org/10.1007/978-981-19-1600-7_5

692 Khong, T. D., Young, M. D., Loch, A., & Thennakoon, J. (2018). Mekong River Delta farm-
 693 household willingness to pay for salinity intrusion risk reduction. *Agricultural Water
 694 Management*, 200, 80–89. <https://doi.org/https://doi.org/10.1016/j.agwat.2017.12.010>

695 Kotuby-Amacher, J., Koenig, R., & Kitchen, B. (2000). Salinity and Plant Tolerance.
696 *UtahStateUniversity, January 2000*, 1–8.

697 Lafta, A. A. (2022). Numerical assessment of Karun river influence on salinity intrusion in the Shatt
698 Al-Arab river estuary, northwest of Arabian Gulf. *Applied Water Science*, 12(6), 124.
699 <https://doi.org/10.1007/s13201-022-01640-4>

700 Lee, H. L., Tangang, F., Hamid, M. R., Benson, Y., & Razali, M. R. (2016). Modeling the Influence
701 of River Flow and Salt Water Intrusion in the Terengganu Estuary, Malaysia. *IOP Conference*
702 *Series: Materials Science and Engineering*, 136(1), 12076. [https://doi.org/10.1088/1757-](https://doi.org/10.1088/1757-899X/136/1/012076)
703 [899X/136/1/012076](https://doi.org/10.1088/1757-899X/136/1/012076)

704 Lindenschmidt, K.-E. (2006). The effect of complexity on parameter sensitivity and model
705 uncertainty in river water quality modelling. *Ecological Modelling*, 190(1), 72–86.
706 <https://doi.org/https://doi.org/10.1016/j.ecolmodel.2005.04.016>

707 Liu, B., Liao, Y., Yan, S., & Yan, H. (2017). Dynamic characteristics of saltwater intrusion in the
708 Pearl River Estuary, China. *Natural Hazards*, 89(3), 1097–1117.
709 <https://doi.org/10.1007/s11069-017-3010-4>

710 Liu, B., Peng, S., Liao, Y., & Wang, H. (2019). The characteristics and causes of increasingly severe
711 saltwater intrusion in Pearl River Estuary. *Estuarine, Coastal and Shelf Science*, 220, 54–63.
712 <https://doi.org/https://doi.org/10.1016/j.ecss.2019.02.041>

713 Maruthai, S., Rajendran, S., Selvanarayanan, R., & Gowri, S. (2025). Wastewater recycling
714 integration with IoT sensor vision for real-time monitoring and transforming polluted ponds into
715 clean ponds using HG-RNN. *Global Nest Journal*, 27(4). <https://doi.org/10.30955/gnj.06758>

716 Meisner, C. (2007). WPS4136 The Impact of Sea Level Rise on Developing Countries. *Tropical*
717 *Agriculture*.

718 Miller, R. L. (2023). Forcing uncertainty and salinity response to dredging in a tidal freshwater river.
719 *International Journal of River Basin Management*, 21(2), 243–251.
720 <https://doi.org/10.1080/15715124.2021.1961795>

721 Ministry of Natural Resources and Environment of Vietnam. (2021). *Climate Change Scenario*
722 (2020th ed.). NARENCA.

723 NARBO (Network of Asian River Basin Organizations). (2006). *Flood management in the Red-Thai*
724 *Binh river basin*. The 2nd General Meeting (Indonesia).
725 https://www.narbo.jp/data/01_events/materials/gm02_2_2_05pa.pdf

726 Oude Essink, G. H. P. (2001). Improving fresh groundwater supply—problems and solutions. *Ocean*
727 *& Coastal Management*, 44(5), 429–449. [https://doi.org/https://doi.org/10.1016/S0964-](https://doi.org/https://doi.org/10.1016/S0964-5691(01)00057-6)
728 5691(01)00057-6

729 Perales-Valdivia, H., Sanay-González, R., & Valle-Levinson, A. (2018). Effects of tides, wind, and
730 river discharge on the salt intrusion in a microtidal tropical estuary. *Regional Studies in Marine*
731 *Science*, 24, 400–410. <https://doi.org/https://doi.org/10.1016/j.rsma.2018.10.001>

732 Rice, K. C., Hong, B., & Shen, J. (2012). Assessment of salinity intrusion in the James and
733 Chickahominy Rivers as a result of simulated sea-level rise in Chesapeake Bay, East Coast,
734 USA. *Journal of Environmental Management*, 111, 61–69.
735 <https://doi.org/10.1016/j.jenvman.2012.06.036>

736 Shen, Y., Jia, H., Li, C., & Tang, J. (2018). Numerical simulation of saltwater intrusion and storm
737 surge effects of reclamation in Pearl River Estuary, China. *Applied Ocean Research*, 79(July),
738 101–112. <https://doi.org/10.1016/j.apor.2018.07.013>

739 Thai, T. H., & Thuc, T. (2011). Impacts of climate change on the flow in Hong-Thai Binh and Dong
740 Nai river basins. *VNU Journal of Science, Earth Sciences*, 27(December), 98–106.

741 Thanh, T. N., Huynh Van, H., Vo Minh, H., & Tri, V. P. D. (2023). Salinity Intrusion Trends under
742 the Impacts of Upstream Discharge and Sea Level Rise along the Co Chien River and Hau River
743 in the Vietnamese Mekong Delta. *Climate*, 11(3). <https://doi.org/10.3390/cli11030066>

744 Tian, R. (2019). Factors controlling saltwater intrusion across multi-time scales in estuaries, Chester
745 River, Chesapeake Bay. *Estuarine, Coastal and Shelf Science*, 223, 61–73.
746 <https://doi.org/https://doi.org/10.1016/j.ecss.2019.04.041>

747 Tu, V. T., Thuy, N. T., Hien, L. T. T., Lan, P. T. H., & Thao, N. H. P. (2021). *Preventing drought*
 748 *and saltwater intrusion by structural and non-structural solutions* (In Vietnam). Construction
 749 Publishing House, Vietnam.

750 UNDP Vietnam. (2016). *Vietnam Drought and Saltwater Intrusion: Transitioning from Emergency*
 751 *to Recovery*. July 17.

752 Veerapaga, N., Azhikodan, G., Shintani, T., Iwamoto, N., & Yokoyama, K. (2019). A three-
 753 dimensional environmental hydrodynamic model, Fantom-Refined: Validation and application
 754 for saltwater intrusion in a meso-macrotidal estuary. *Ocean Modelling*, 141, 101425.
 755 [https://doi.org/https://doi.org/10.1016/j.ocemod.2019.101425](https://doi.org/10.1016/j.ocemod.2019.101425)

756 Venkatraman, M., Surendran, R., Srinivasulu, S., & Vijayakumar, K. (2025). Water quality prediction
 757 and classification using attention based deep differential recurflownet with logistic giant
 758 armadillo optimization. *Global Nest Journal*, 27(1), 1–13. <https://doi.org/10.30955/gnj.06799>

759 Vuong Tai, C., Nguyen Khanh, L., Vu Huong, N., Dinh Nhat, Q., & Truong Van, A. (2021).
 760 Investigation of saltwater intrusion in thach han river system by mike hydro river package.
 761 *Transport and Communications Science Journal*, 72(1), 46–57.
 762 <https://doi.org/10.47869/tcsj.72.1.6>

763 Wang, J., Li, L., He, Z., Kalhoro, N. A., & Xu, D. (2019). Numerical modelling study of seawater
 764 intrusion in Indus River Estuary, Pakistan. *Ocean Engineering*, 184(May), 74–84.
 765 <https://doi.org/10.1016/j.oceaneng.2019.05.029>

766 Yuan, R., Zhu, J., & Wang, B. (2015). Impact of Sea-Level Rise on Saltwater Intrusion in the Pearl
 767 River Estuary. *Journal of Coastal Research*, 31(2), 477–487.
 768 <https://doi.org/10.2112/JCOASTRES-D-13-00063.1>

769 Zhang, W., Feng, H., Zheng, J., Hoitink, A. J. F., van der Vegt, M., Zhu, Y., & Cai, H. (2013).
 770 Numerical Simulation and Analysis of Saltwater Intrusion Lengths in the Pearl River Delta,
 771 China. *Journal of Coastal Research*, 29(2), 372–382. [https://doi.org/10.2112/JCOASTRES-D-](https://doi.org/10.2112/JCOASTRES-D-12-00068.1)
 772 [12-00068.1](https://doi.org/10.2112/JCOASTRES-D-12-00068.1)

773 Zhang, W., Ruan, X., Zheng, J., Zhu, Y., & Wu, H. (2010). Long-term change in tidal dynamics and
774 its cause in the Pearl River Delta, China. *Geomorphology*, 120(3), 209–223.
775 <https://doi.org/https://doi.org/10.1016/j.geomorph.2010.03.031>